SYNTHETIC FLUOR-MAGNESIO-RICHTERITE

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ABSTRACT

Chemical, physical, optical and x-ray data are provided for fluor-magnesio-richterite, a fluoramphibole synthesized by melt and by solid state reactions. Chemical analysis reveals a unit cell content of $2(\text{Na}_{1.97}\text{Mg}_{6.01}\cdot\text{Si}_{7.97}\text{O}_{22.04}\text{F}_{1.96})$. Physical and optical properties are: perfect cleavage (110) \wedge (1 $\overline{10}$) 55°33′; density 2.994 gm. cm.⁻³; biaxial(-), α =1.5756, β =1.5894 γ =1.5954, $c \wedge$ Z 11.4°. Unit cell dimensions are a 9.929 Å., b 17.914 Å., c 5.274 Å., β 108°16′, space group I2/m.

Introduction

Synthetic asbestos investigations at the U. S. Bureau of Mines, Norris, Tennessee, have resulted in the synthesis of more than 75 fluoramphiboles of varied composition, many being fluoride analogs of the natural hydroxyl amphiboles—tremolite, edenite, richterite, and eckermanite. The syntheses and physical properties of these fluoramphiboles have been reported by Eitel (1952), Comeforo and Kohn (1958), Kohn and Comeforo (1955), and Shell, Comeforo, and Eitel (1958). Described here is a fluoramphibole referred to as fluor-magnesio-richterite (Na₂Mg₆Si₈O₂₂F₂) which differs from fluor-richterite (Na₂CaMg₅Si₈O₂₂F₂) by the substitution of an additional Mg²⁺ for Ca²⁺. When synthesized by pneumatolytic action is formed as an aggregate of acicular crystals, some of which approached fiber dimensions.

CHEMICAL SYNTHESIS

Fluor-magnesio-richterite was synthesized both from a melt and by pneumatolysis. The theoretical formula is Na₂Mg₆Si₈O₂₂F₂, however, batch compositions containing more fluoride were actually used. The excess fluoride provided for loss of volatiles and promoted reactivity. The batch composition Na₂Mg₆Si₈O_{21.5}F₃, when melted and crystallized, yielded acicular fluoramphibole as the major phase. The batch compositions (1) Na₂Mg₆Si₈O_{21.5}F₄, (2) Na₂Mg₆Si₈O_{20.5}F₅, and (3) Na₂Mg₆Si₈O_{19.5}F₇ were reacted by pneumatolysis. Typically the pneumatolytic product consisted of a mass of short interpenetrating needles considerably less in diameter than those obtained from the melt.

Fluor-magnesio-richterite is the most acicular of any fluoramphibole synthesized by us at this laboratory. In fact, some of the crystals synthesized by pneumatolytic and solid state reaction approached fiber dimensions; they were commonly too small in diameter to be resolved optically at 450×.

One non-stoichiometric batch composition (for amphibole) gave excellent single crystals for the x-ray and optical study. It was Na_{6.4}Mg_{8.6}Si₈O_{20,14}F_{12.8}. From this melt fluor-magnesio-richterite and NaMgF₃ were the major products. This composition was reacted because of a report by Saito (1955) that an orthorhombic amphibole was obtained therefrom. We, however, found only the monoclinic amphibole described herein.

Batch materials were: analytical reagent grade NaF and Na₂CO₃, technical grade MgO and MgF₂, and -200 mesh quartz that was 99.8 per cent SiO₂.

The properly proportioned batch materials were thoroughly admixed. For synthesis by melting, the blended raw materials were fired in a sealed graphite crucible which was placed inside a fire-clay crucible to prevent oxidation. After 2 hours at 1250° C., the temperature was lowered to below the solidus at 10° C./hour. Either graphite or platinum crucibles prevented contamination of the melt; both types were used. For synthesis by pneumatolytic-solid state reaction, fire-clay crucibles alone were satisfactory.

CHEMISTRY

The chemical analysis (Table 1, col. 2) was obtained on crystals selected from three separate reaction products. The crystals were selected under the binocular microscope and purified by electrodialysis at 400 volts for 19 hours.

The empirical unit cell content of fluor-magnesio-richterite was found by determining the mass in a mole of unit cells and allocating this mass to the various oxides on the basis of their respective percentages reported in the chemical analysis. The density of the fluoramphibole, $\rho = 2.994$ gm. cm.⁻³ at 27° C., was measured on single crystals using the Roller-Smith precision balance with toluene as the suspension medium. The unit cell volume V was calculated from the x-ray powder diffraction data to be 890.8067A3. Substituting these values in the relation ZM $= \rho V/1.66020$, the mass in a mole of unit cells (ZM), was determined to be 1606.47 (Table 1). Assuming Z (the number of formula units for clinoamphibole) to be 2, the chemical formula is determined, except for the oxygen, by dividing the value in column 3 by the values in column 4 (except for Na₂O, the number for which was doubled to yield Na). The oxygen was, however, readily calculated according to the principle of electrostatic neutrality and the formula for fluor-magnesio-richterite becomes Na_{1.97}Mg_{6.01}Si_{7.97}O_{22.04}F_{1.96}. A normalization of the chemical analysis on the basis of 24 ions after Shell and Craig (1956) results in the formula Na_{1.98}Mg_{6.03}Si_{7.99}O_{22.04}F_{1.96}. The two unit formulae, each calculated on a different basis, show good agreement.

Table 1. Calculation of the Unit Cell Content Analyst, R. E. Hooper

	Analysis wt. %	Mass units per mole of unit cells	Molecular wt.	Number of ions in the unit cell
SiO ₂	59.30	957.33	60.09	15.95
${ m MgO^2}$	30.00	484.31	40.32	12.02
Na_2O	7.55	121.89	61.98	3.94
F	4.60	74.26	19.00	3.908
	101.45	1637.79		
Less $O = F$	-1.94	-31.32		
Total	99.51	1606.47		

GONIOMETRIC DATA

Crystals studied with the Eichner two circle goniometer are typically acicular in the c-axis direction but were rarely terminated. The angle $(110) \wedge (1\overline{10})$ is 55°33′ (average of four measurements from different crystals). Calculation of this angle from the x-ray data yields 55°30′. For the angle $(110) \wedge (100)$ the measured angle is 62°12′ and the calculated angle 62°15′. Perfect cleavage is parallel to the c-axis.

OPTICAL CONSTANTS

The refractive indices of fluor-magnesio-richterite were measured using a 5-axis universal stage and the single variation technique and were plotted on Hartmann dispersion nets. Straight lines were fitted to the plots by a non-statistical method for the purpose of averaging the measured values. The refractive indices given in Table 2 were obtained directly from the Hartmann nets by reading the intersection of the dispersion lines of the grains with the F, D, and C Frauenhofer lines. The accuracy of these results is believed to be ± 0.0005 .

TABLE 2. OPTICAL CONSTANTS FOR FLUOR-MAGNESIO-RICHTERITE

A	γ	β	α	$\gamma - \alpha$	$\gamma - \beta$	Optic Sign	2V* (calc.)
F (4861)	1.6007	1.6952	1.5850	-0157	.0055	(-)	72.1
D(5893)	1.5954	1.5894	1.5756	.0198	.0060	(-)	67.8
C (6563)	1.5933	1.5869	1.5718	.0215	.0064	(-)	65.8
$N_F - N_c$.0074	.0083	.0132				

^{*} By direct measurement 2V = 70.2 using a white light source; $c \triangle Z = 11.4^{\circ}$.

SINGLE CRYSTAL DATA

A clear, single crystal of fluor-magnesio-richterite was mounted on the precession instrument, and the first two or three reciprocal lattice layers along [100], [010] and [001] were photographed. All spectra recorded were consistent with the diffraction rules of the space group I2/m as reported for clinoamphibole by Warren (1930).

POWDER DIFFRACTION DATA

The x-ray powder diffraction data were obtained using a Norelco highangle counter diffractometer operated in conjunction with an automatic strip chart recorder. Crystals of the fluoramphibole were optically selected, ground under acetone, and the fraction passing the -400 screen was used in making the smear mounts to be x-rayed. The smear mounts were prepared by admixing the powder with a finely ground internal standard and adhesive lacquer, and spreading the resultant slurry as a uniformly thin, translucent film on a BT-cut quartz oscillator plate. The use of the quartz plate is believed to minimize the background radiation thereby permitting a more precise location of the 2θ maxima (Buerger and Kennedy, 1958). The lacquer consisted of about 10 parts ducocement dissolved in about 100 parts acetone. An internal standard, a synthetic spinel MgAl₂O₄ calibrated by fixed counting diffractometer techniques to transistor grade silicon a = 5.4306 Å, was used to correct the diffraction peaks of the fluor-magnesio-richterite and to eliminate, in general, the systematic errors inherent in the diffractometer method.

The settings of the diffractometer were: divergence and scatter slits, 1°; receiving slit, 0.006 inches; scanning speed, $\frac{1}{4}$ degree 2θ per minute; strip chart scale, $\frac{1}{2}$ degree per inch; time constant, 4; multiplier, 1; scale factor, 16; and Ni filtered Cu radiation, $K_{\alpha}=1.54178$ Å, $K_{\alpha 1}=1.54051$ Å.

The 2θ values recorded on the diffraction charts were located by measuring the peak centers at two-thirds height as proposed by Donnay and Donnay (1951). This procedure of peak location was used, however, only for peaks showing completely resolved α_1 and α_2 spectral components. For unresolved or partially resolved peaks, the peaks were located by measuring their centers at half height (Smith and Sahama, 1954). The diffraction peaks were then corrected to the internal standard, and comparable 2θ values from four separate strip charts were averaged and converted to quadratic values from the tables of Gibbs and Lewis (1961). The assignment of diffraction indices to the powder diffraction record was made by comparison with the Kohn and Comeforo (1955) data for synthetic fluor-richterite and by use of the nomogram of Bloss and Gibbs (1961). The final indexing of the record was confirmed by visually comparing the relative intensities of the indexed spectral components re-

corded on the precession films with those of comparable powder diffraction peaks. The dimensions of the unit cell were found by a least squares refinement of forty-four reflections of the powder pattern as a = 9.929 Å, b = 17.914 Å, c = 5.274 Å, $\beta = 108^{\circ}16'$, all ± 0.05 per cent, space group

TABLE 3. POWDER DIFFRACTION DATA

hkl I/I_0	Observed Values ¹			Calculated Values ¹			ΔQ×10 ⁵	
	d	2θ	Q×10 ⁵	d	2θ	Q×10 ⁵ ²	ΔQΧ10°	
020	4	8.960	9.870	1,245	8.956	9.875	1,246	- 1
110	58	8.346	10.597	1,435	8.343	10.600	1,436	- 1
011	14	4.825	18.385	4,295	4.824	18.390	4,297	- 2
200	21	4.720	18.798	4,488	4.716	18.818	4.497	- 9
040	24	4.481	19.809	4,979	4.478	19.824	4,986	-7
21T	16	4.033	22.036	6.147	4.034	22.035	6,146	+ 1
031	39	3.840	23.170	6,786	3.837	23.177	6,790	- 4
$23\overline{1}$	52	3.404	26.191	8,639	3.403	26.193	8,640	- 1
150	11	3.353	26.587	8,897	3.351	26.614	8,915	-18
240	75	3.249	27.453	9,475	3.247	27.466	9,483	- 8
310	100	3.099	28.813	10,416	3.097	28.834	10,431	-15
$32\overline{1}$	40	2.964	30.150	11,383	2.964	30.145	11,379	+4
051	26	2.916	30.663	11,764	2.915	30.679	11,776	-12
330	27	2.784	32.150	12,902	2.782	32.179	12,924	-22
$25\overline{1}$	92	2.711	33.048	13,612	2,709	33.065	13,626	-14
231	18	2.680	33.427	13,918	2,678	33.450	13,936	-18
$16\overline{1}$	30	2.581	$34.726\alpha 1$	15,011	2.582	34.718	15,004	$+^{10}$
002	27	2.505	35.856	15,944	2.505	35.864		+ 4
170	< 5	2.469	36.380	16,398	2.469	36.374	16,393	$+\frac{1}{5}$
350	21	2.469		17,036	2.469	38.049	17,910	$^{+3}_{+26}$
	27		$38.078\alpha 1$	17,936	2.303	39.185	18,922	-27
251		2.300	39.157	18,895 19,246		20 542	19,254	-27 - 8
071	22	2.279	39.533	19,240	2.279	39.542	20, 024	$^{-34}$
112	7	2.233	$40.359\alpha 1$	20,058	2.235	40.323	20,024	-28
332	< 5	2.188	41.264	20,893	2.186	41.257	20,921	$-28 \\ -36$
271	5	2.178	41.444	21,068	2.177	41.445	21,104	
361	30	2.163	$41.716\alpha 1$	21,368	2.164	41.699		+17
132	< 5	2.108	$42.869\alpha 1$	22,509	2.107	42.877	22,517	-8
202	< 5	1.972	$45.989\alpha 1$	25,722	1.971	46.001	25,734	-12
190	< 5	1.948	$46.583\alpha 1$	26,353	1.947	46.595		-12
062	7	1.916	$47.397\alpha 1$	27,229	1.919	47.332		+71
510	13	1.876	48.470a1	28,400	1.875	48.488		-21
091	8 7	1.849	$49.234\alpha 1$	29,246	1.850	49.215		+20
530	7	1.797	$50.755\alpha 1$	30,960	1.799	50.714	30,914	+46
291	< 5	1.659	$55.342\alpha 1$	36,350	1.658	55.360		-22
480	13	1.623	$56.647\alpha 1$	37,941 38,795	1.624	56.640	37,933	+ 8
1.11.0	6	1.605	$57.339\alpha1$	38,795	1.605	57.367	38,830	-35
$25\bar{3}$	9	1.577	58.494a1	40,235	1.577	58.475	40,211	+24
600	20	1.571	$58.710\alpha 1$	40,506	1.572	58.687		+29
$65\overline{1}$	7	1.503	$61.667\alpha 1$	44,279	1.502	61.691		-31
402	< 5	1.499	$61.859\alpha 1$	44,528	1.499	61.858	44,527	+1
0.12.0	17	1.493	$62.120\alpha 1$	44,867	1.493	62.125	44,873	- 6
3.11.0	< 5	1.446	$64.394 \alpha 1$	47,853	1.446	64.372	47.825	+28
561	16	1.415	$65.942\alpha 1$	49,921	1.415	65.951	49,932	-11
563	7	1.344	$69.922\alpha 1$	55,345	1.344	69.916	55,336	+ 9
0.12.2	5	1.282	$73.877\alpha 1$	60,873	1.282	73.835	60,813	+60

 $^{^1}$ Values for CuK α unless otherwise indicated; CuK α 1.54178 Å, CuK α_1 , 1.54050 Å. 2 Based on the following lattice constants: $a\!=\!9.929$ Å, $b\!=\!17.514$ Å, $c\!=\!5.274$ Å, $\beta\!=\!108^\circ14'$, space group I2/m.

I2/m. Upon transformation to the C2/m orientation using the matrix $101/010/00\overline{1}$ set forth by Zussman (1959), the unit cell dimensions become a=9.677 Å, b=17.914 Å, c=5.274 Å, $\beta=102^{\circ}57'$, all ±0.05 per cent.

DISCUSSION

The general formula for amphibole may be given as $W \cdot X_2 \cdot Y_5 \cdot Z_8 \cdot O_{22} \cdot (OH, F)_2$ where W refers to ions of radius 0.7 to 1.3 Å which can occupy the A site at $(\frac{1}{2}, O, \frac{1}{2})$ in the structure (Whittaker, 1960), chiefly K⁺, Na⁺, Ca²⁺, Mg²⁺ and Li⁺; X refers to ions of radius 0.7 to 1.1 Å which occupy the M₄ sites at the edges of the amphibole chains, chiefly Ca²⁺, Na⁺, Fe²⁺, Mg²⁺, Li⁺, and Mn²⁺; Y refers to ions of radius 0.5 to 0.9 Å sandwiched in six-fold coordination between the chains, chiefly Mg²⁺, Fe²⁺, Mn²⁺, Fe³⁺, Al³⁺, Li⁺, and Ti⁴⁺; and Z refers to small, high valence ions of four-fold coordination in the chains, principally Si⁴⁺ but to a lesser extent (up to 25 per cent) Al³⁺.

Crystal chemical principles permit a likely assignment of sodium and magnesium to sites in flour-magnesio-richterite. As Whittaker (1960) has shown, the average radius of the ions filling the M_4 sites in clinoamphibole controls in general the value of the crystallographic angle β ; that is, larger ions reduce the displacement between the neighboring (Si₄O₁₁) chains and thereby the magnitude of β . Whittaker demonstrated this relation in a regression plot of the β angles on the average radii of the ions filling the M_4 sites for five structurally determined amphiboles and for amosite (Fig. 1). Since Whittaker published this regression, structural

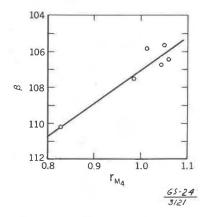


Fig. 1. Plot of β against the mean Goldschmidt radius of the ion at M_4 for the five clinoamphiboles whose structures have been published, and for amosite. The line is the regression line of β on r (After Whittaker).

analyses for grunerite and cummingtonite have been carried out by Ghose (1959, 1961). Employing the average radii of the ions filling the M_4 sites in these amphiboles, we read the β angles from Whittaker's regression lines as 109°59′ and 110°12′ which agree remarkably well with Ghose's experimental values of β , 109°50′, and 109°54′, respectively. To be consistent with the regression plot, it was necessary to transform the cell setting of the cummingtonite from a C-face centered cell as given by Ghose to a I-centered cell.

In fluor-magnesio-richterite, sodium and magnesium ions differ sufficiently in radii that a filling of the M4 sites by one rather than the other of these ions should cause an appreciable difference in β . Since Na⁺ tends to fill only the A and M4 sites in amphibole (Whittaker, 1949; Heritch, et al., 1957, 1960a, 1960b) one can give, in general, three possible distributions of Mg²⁺ and Na⁺ in fluor-magnesio-richterite: (1) Mg²⁺ filling the A site and 2Na+ the M₄ sites; (2) Na+ filling the A site and both Na+ and Mg²⁺ the M₄ sites; or (3) Mg²⁺ and Na⁺ randomly filling the A and M₄ sites. Depending upon whether (1), (2), or (3) pertains, the average radius for the ions filling the M₄ sites would be (1) 0.98, (2) 0.88 and (3) 0.91 Å, respectively. The β angles read from Whittaker's regression line for each of these possibilities are 107°24′, 109°08′ and 108°36′. The measured β angle for fluor-magnesio-richterite is 108°16′ which suggests that the third possibility is more correct, namely, Na+ and Mg²⁺ are randomly distributed in the A and M4 sites. If this is true, then the formula for fluor-magnesio-richterite should be given as Na2MgMg5Si8O22F2. However, this particular distribution of ions would be very difficult to verify by an x-ray structural analysis because of the similarity in electron density of Na+ and Mg2+.

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