

Fig. 1. Calcite with epitaxis of magnesite.

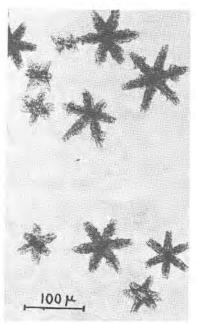


Fig. 2. Starshaped crystals of dolomite.

fraction patterns and by chemical analyses. These gave for the rhombohedral crystals: MgO 21.3%, CaO 29.8, organic residue insoluble in HCl 2n 2.2, loss on ignition 46.2, and for the starshaped crystals: MgO 22.4, CaO 31.2, residue 0.2, loss on ignition 46.4.

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THE UNIT CELL OF KRAUSITE1

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The mineral krausite, an anisodesmic oxysalt of potassium and iron, was discovered in association with alunite and coquimbite at Borate in

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the Calico Hills, San Bernardino County, California. It was described by Foshag (1931), who gave the composition as $K_2SO_4 \cdot Fe_2(SO_4)_3 \cdot 2H_2O$.

The sample of krausite available to the authors from the above locality consisted of an aggregate of pale-yellow short prismatic crystals less than 1.0 mm. in length. Brilliant crystals of approximately 0.5 mm diameter were chosen for single crystal studies.

Measurements of the unit-cell dimensions from single crystal photographs as well as from powder photographs (Fe $K\alpha$ radiation) led to the following lattice constants:

$$a = 7.910 \pm 0.01 \text{ Å}$$
 $\beta = 102^{\circ}45' \pm 5'$
 $b = 5.152 \pm 0.005 \text{ Å}$ $V = 357.2 \text{ Å}^3$
 $c = 8.990 \pm 0.01 \text{ Å}$ $a:b:c: = 1.535:1:1.745$

These values are in better agreement with the morphological constants reported by Foshag (1931) than with the recalculated values reported by LaForge (Palache; et al., 1951). For a unit cell containing two molecules KFe(SO₄)₂·H₂O, a density of 2.839 g.cm⁻³ was calculated; the measured density is 2.840 g.cm⁻³ (Foshag, 1931).

X-ray diffraction patterns were obtained of the hol, h1l, h2l levels by the Weissenberg method (Fe K α radiation) and the 0kl, h0l levels by the precession method (Mo K α radiation). The only systematic extinctions were of the type 0k0 absent for k odd, indicating that the space group is either $P2_1$ or $P2_1/m$.

The morphology is consistent with 2/m point symmetry. Etch pits produced by dilute hydrochloric acid on (001) cleavage surfaces indicate the presence of a mirror. Statistical distribution of observed diffraction intensities according to the method of Howells $et\ al.$ (1950) is in agreement with that of a centrosymmetric structure. Reflections of the type hol were omitted from this calculation because either space group would appear centric when projected onto the a-c plane. It appears that $P2_1/m$ is the correct space group.

There is an apparent relationship among the unit cell dimensions of yavapaiite (Hutton, 1959), goldichite (Rosenzweig and Gross, 1955), and krausite, particularly in the b axis lengths as shown in Table I. Hutton (1959) pointed out the similarity between these compounds in regard

TABLE I

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Name	Composition	S.G.	Z	a	b	G	β
Yavapaiite	KFe(SO ₄) ₂	C2/m (?)	2	8.12	5.14	7.82	94°24′
Krausite	KFe(SO ₄) ₂ ·H ₂ O	$P2_1/m$	2	7.908	5.152	8.988	102°45′
Goldichite	KFe(SO ₄) ₂ ·4H ₂ O	$P2_1/c$	4	10.45	10.53	9.15	101°49′

Table II. X-ray Powder Data for Krausite FeK α radiation, camera diameter 114.6 mm

d_{obs} Å	I_{obs}	d_{calc} Å	hkl
8.82	4	8.77	001
7.75	37	7.71	100
6.59	66	6.55	10∏
4.40	74	4.38	002
4.26	45	4.23	$10\overline{2}$
4.06	0	4.05	11 <u>T</u>
3.86	9 2	3.86	200
2 60			
3.69	64	3.68	111
3.51	22	3.49	102
3.35	2	3.34	012
3.28	18	3.27	201
3.09	100	3.09	21∏
		3.09	210
2.970	9	2.958	$10\overline{3}$
2.933	6	2.922	003
2.901	3	2.891	112
2.768	40	2.765	$21\overline{2}$
2		2.762	211
2.627	6	2.629	$30\overline{1}$
2.579	38	2.576	020
2.546	38	2.542	013
2.342	2 6	2.342	311
2.302	6	2.300	310
2.002	O .	2.286	113
2.251	2		
	$\frac{2}{4}$	2.242	104
2.199	4	2.221	022
		2.201	$12\overline{2}$
2.191	4	2.191	004
		2.185	303
2.146	7	2.142	221
		2.142	220
2.078	4	2.073	122
2.025	10	2.025	$\frac{122}{222}$
4.023	10		221
		2.024	
0.004	2	2.016	014
2.001	3	1.995	104
1.964	3		
1.928	2		
1.893	3 2 18 5 3 7 1		
1.840	18		
1.806	5		
1.773	3		
1.753	7		
	1		
1.731	1		
1.703	1		
1.665	2		
1.637	3		
1.570	2 3 7		
1.545	4		
1.524	2		
1.480	2		
	2 8 2 0		
1.449	20		
1.415	4		
1.288 1.275	4		
	10		

to the optical properties and cleavage. It is noteworthy that in krausite and yavapailte the b axis coincides with the optically slow direction Z; but in goldichite, where that axis is doubled, it coincides with the fast direction X.

X-ray powder diffraction data for Fe K α radiation is given in Table II; the intensities were measured by a double-beam recording microdensitometer. The crystal structures of krausite and its related hydrates are currently under investigation.

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FURTHER DATA ON THE UNIT CELL AND SPACE GROUP OF CHERALITE

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Bowie and Horne (1953) described cheralite as a green mineral belonging to the monazite group intermediate in composition between monazite and the compound CaTh(PO₄)₂. They presented chemical analyses from two supposedly different localities: Kuttakuzhi, Travancore, India and Cootykad Pothay, Vilavancode taluq, Travancore. These localities are now confirmed by the senior author to be one and the same. Bowie and Horne included approximate unit cell dimensions based on powder data and concluded that monazite and cheralite are isostructural.

Khubchandani (1956) determined the unit cell dimensions and space group of cheralite from rotation and equatorial Weissenberg photographs. He reported the space group as P $2_1/n$, C_2^5h , identical to that listed for monazite.

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