The refinement of zinnwaldite-1M in subgroup symmetry

STEPHEN GUGGENHEIM¹ AND S. W. BAILEY

Department of Geology and Geophysics, University of Wisconsin Madison, Wisconsin 53706

Abstract

The structure of a zinnwaldite-1M from the Sadisdorf Mine, D.D.R., has been refined in C2/m, C2, and C1 symmetries. The composition is $(K_{0.90}Na_{0.05})$ $(Al_{1.05}Fe_{0.16}^{2+}Ti_{0.01}Fe_{0.77}^{2+}Mn_{0.05})$ $Mg_{0.01}Li_{0.87}(Si_{3.09}Al_{0.91})O_{10}(OH)_{0.79}F_{1.21}$. The cell parameters are a = 5.296, b = 9.140, c =10.096 A, and $\beta = 100.83^{\circ}$. Refinement in triclinic symmetry C1 using the ordered-model approach gave atomic coordinates consistent with monoclinic C2 symmetry. In C2 symmetry, octahedra M(2) and M(3) related by the pseudo-mirror plane are significantly different in size (mean M—O,F = 1.882,2.131 A) and, to a lesser extent, in electron count (11.5 and 13.5). Octahedral Al completely occupies M(2), and the remaining Fe, Li, other cations, and vacancies are nearly randomly distributed over M(3) and the trans octahedron M(1). The (F,OH) atom has moved off the pseudo-mirror plane in order to coordinate more closely with Al in M(2). It is anticipated that all fluorine-rich zinnwaldites and lepidolites will have a similar ordering pattern. The cis-orientation of fluorine along the M(2): M(3) shared edge allows a closer F-F approach around a small cation in either site than is possible along any M(1) shared edge involving larger oxygen atoms. The larger M(1) and M(3) octahedra are flattened considerably ($\psi = 60.8^{\circ}$) in order to fit onto the smaller, more regular ($\psi = 56.5^{\circ}$) M(2) octahedron. The two nonequivalent tetrahedra differ slightly in size (mean T—O = 1.646,1.639 A), are elongate, and rotated by 5.8°.

Introduction

Zinnwaldite is a Li,Fe,Al trioctahedral mica commonly found in greisens, associated veins, and in some granites and pegmatites. It is defined by Foster (1960) as having a Li content of 1.00 \pm 0.25 atoms per formula unit ($\sim 2.50-4.50$ weight percent Li₂O), in contrast to protolithionite with 0.50 ± 0.25 atoms and lepidolite with >1.25 atoms. Rieder (1968, 1970a) and Rieder et al. (1971) have studied a suite of natural Li-Fe micas, from the Krušné hory Mountains (Erzgebirge) of Czechoslovakia and Germany, that lie close to the join between polylithionite and siderophyllite. As a result of this study Rieder (1970a) proposed broadening the composition of zinnwaldite so that it ranges from x = 1 to 3 in the series $K_2Fe_x^{2+}Li_{4-x}$ $(Al,Fe^{3+})_2(Al_xSi_{8-x})O_{20}(OH,F)_4$, where x = 0 for polylithionite and x = 4 for siderophyllite. Most of the zinnwaldites were of the 1M structural type (Rieder, 1970b).

Rieder (1968) agrees with Foster (1960) that there always is close to 1.0 small trivalent octahedral cation

(Al,Fe³⁺) per formula unit in zinnwaldite. Although the larger octahedral cations (Li,Fe²⁺,Mg,Mn) are somewhat variable in number, they never exceed 2.0. This ideal ratio of 1:2 led Rieder to postulate octahedral ordering for zinnwaldite-1M of the type found in clintonite-1M (= xanthophyllite), wherein each small M(1) octahedron on the mirror plane of the ideal C2/m space group is surrounded symmetrically by six larger M(2) octahedra in an overall 1:2 ratio. This ordering was believed to account for the observation that d(001) values for natural zinnwaldites are about 0.1 A smaller than for synthetic specimens, which are presumed to be disordered.

Approximately two dozen mica structures have been refined that show octahedral ordering, either of cations or, for dioctahedral compositions, of cations plus vacancies. Clintonite-1M (= xanthophyllite) is the only exception to the rule that the unique octahedron M(1) on the symmetry plane of each layer is larger than the two equivalent M(2) octahedra. The "normal" ordering pattern with M(1) larger than the average M(2) site has been shown to be adopted even in cases where the octahedral composition logically might suggest the reverse pattern. An example is

¹ Present address: Dept of Geological Sciences, University of Illinois at Chicago Circle, Box 4348, Chicago, Illinois 60680.

polylithionite-1M, in which there are two large octahedral Li ions and one smaller Al ion. However, the ordering pattern found by Takeda and Burnham (1969) consists of one large M(1) octahedral site on the mirror plane with composition $\text{Li}_{0.89}\text{Al}_{0.11}$ and two smaller symmetry-related M(2) sites of average composition $\text{Li}_{0.55}\text{Al}_{0.45}$.

This paper reports an X-ray refinement of one of Rieder's zinnwaldite specimens. This study was undertaken, first, to test the Rieder ordering hypothesis, because it would be of particular interest to find another exception to the prevalent ordering pattern. Second, even if the "normal" ordering pattern were to be found, the composition appears so favorable for ordering that the possibility of crystallographically-distinct M(2) sites should be investigated in subgroup symmetry. Refinement in subgroup symmetry has not been attempted often, but recently has been successful in showing tetrahedral ordering in margarite- $2M_1$ (Guggenheim and Bailey, 1975) and both tetrahedral and octahedral ordering in a dioctahedral 1M mica (Sidorenko *et al.*, 1975).

Experimental

Dr. Milan Rieder of Charles University kindly furnished a suite of analyzed zinnwaldite samples. A crystal from sample #40 from the Sadisdorf Mine dump, D.D.R., was selected as giving the sharpest reflections. The crystal is $0.4 \times 0.3 \times 0.05$ mm and light amber in color. The composition of the crystal used for the structural analysis was determined by electron microprobe analysis (Table 1) after the completion of data collection. The Li content and the ratio of Fe²⁺ to Fe³⁺ were taken from the wet-chemical analysis data for the bulk sample as reported by Rieder (1970a). The resulting formula unit is $(K_{0.90} Na_{0.05})(Al_{1.05} Fe_{0.16}^{3+} Ti_{0.01} Fe_{0.77}^{2+} Mn_{0.05} Mg_{0.01} Li_{0.67}$ $\bigcup_{0.28}$)(Si_{3.09}Al_{0.91})O₁₀(OH)_{0.79}F_{1.21}. This formula is similar to that given by the bulk analysis, except for a slightly lower Si content. Rieder et al. (1971) cite $2V_{\alpha} = 27^{\circ} \pm 1^{\circ}$ and the γ refractive index as 1.595 \pm 0.001 for this specimen.

Unit-cell parameters of a = 5.296(1), b = 9.140(2), c = 10.096(3) A, and $\beta = 100.83(2)^{\circ}$ were determined by least-squares refinement of 15 high-angle reflections measured on a Syntex P2₁ autodiffractometer. Two intensity data sets were collected. First, the intensities of 2,993 non-zero reflections were measured, using graphite-monochromatized MoK α radiation. The data were collected in the $2\theta:\theta$ variable-scan mode in four quadants of the limiting sphere from $2^{\circ} < 2\theta < 90^{\circ}$. Two standard reflections were mon-

Table 1. Chemical analysis of Sadisdorf zinnwaldite

Oxide	^a Wet		percent Best data	Cations per 22 positive charges				
SiO ₂ Al ₂ O ₃ TiO ₂	42.86 21.89 0.16	40.70 21.95 0.20	40.70 21.95 0.20	Si Al	3.090 { 0.910} { 1.054}	4.000 ^{IV}		
Fe ₂ O ₃ FeO MnO	2.78 11.84 0.73	15.04	{2.85 12.19 0.70	Ti Fe3+ Fe2+	0.011 0.163 0.774	2.721 ^{VI}		
Mg0 Li ₂ 0 Ca0	b _{tr} 2.19 0.51	0.04 c _{na} d _{bd}	0.04 2.19 0.00	Mm Mg Li	0.045 0.005 0.669			
Na ₂ 0 K ₂ 0	0.26 9.85	0.36 9.29	0.36 9.29	Na K	0.053 0.900	0.953 ^{XII}		
Р ₂ 05 Н ₂ 0 1 Н ₂ 0-	0.04 na 0.96	na na na	0.00 na 0.96					
H ₂ O- F Sum of	5.20 99.27	5.67 93.95	5.67 97.10					

^aWet chemical analysis by M. Huka and J. Obermajer, Geological Survey of Czechoslovakia, as reported by Rieder (1970a). ^bTrace. ^cNot analyzed. ^dBelow detection.

itored after every 50 reflections to check crystal and electronic stability. Reflections were considered as observed if $I > 2\sigma(I)$, where I was calculated from I = $[S - (B_1 + B_2)/B_r]T_r$, S being the scan count, B_1 and B_2 the background, B_r the ratio of background time to scan time, and T_r the 2θ scan rate in degrees per minute. $\sigma(I)$ was calculated from standard counting statistics. Integrated intensities were corrected for Lorentz and polarization effects, but not for absorption. This data set was used for initial refinement and for determining the best ordering models in both monoclinic and triclinic subgroup symmetry. Final refinement of the monoclinic models used a second data set of 1,550 independent non-zero intensities collected in the same manner from only two quadrants. These data were empirically corrected for absorption by comparing the data to complete ψ scans (10° increments in ϕ) for selected reflections spaced at 2θ intervals of 5°. Absorption is a major factor because of the platy nature of the crystal. A maximum intensity decrease of 49 percent was observed for some reflections during the ψ scans.

Refinement

Initial atomic coordinates in the ideal space group C2/m were obtained from the structure of fluor-polylithionite (Takeda and Burnham, 1969). After several cycles of least-squares refinement with program ORFLS using the first data set with reflections from $2^{\circ} < 2\theta < 55^{\circ}$, scattering-factor adjustments to M(1) and M(2) were made as a result of three-dimensional electron-density difference maps. Additional least-squares refinement reduced R_1 to 12.5 percent. At this stage M(1) was significantly larger than M(2), and

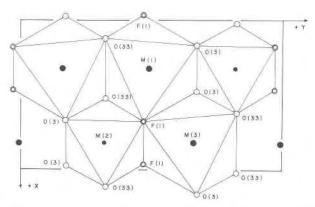


Fig. 1, Octahedral ordering pattern of Sadisdorf zinnwaldite-1M in subgroup C2.

the average (F,OH) atom had an unusually high temperature factor of 3.0 A². This is the "normal" ordering pattern, so subsequent refinement concentrated on subgroup ordering models.

Refinement in subgroup symmetry initially followed the method outlined by Guggenheim and Bailey (1975) for margarite- $2M_1$. Because of convergence problems, it is best not to start refinement in subgroup symmetry by using atomic coordinates of higher symmetry. Instead, atoms were moved away from their pseudosymmetrically-related positions by postulating an ordering scheme in subgroup symmetry that was still consistent with the parent spacegroup refinement. The atomic coordinates of each possible ordering model were determined from a distance-least-squares program, and those results were then refined by varying the parameters of pseudosymmetry-related positions independently. The latter precaution actually may not be necessary, since in the later stages of refinement it proved possible to vary all of the parameters together without significantly high correlations.

For subgroup C2, the M(2) site and its pseudo-symmetrically-related M(3) site become crystallographically independent (Fig. 1). Three-dimensional electron-density difference maps at this stage indicated M(3) had approximately four more electrons per site than M(2). Only two ordered models are compatible with the average structure in the parent space group C2/m, namely M(2) can be smaller or larger than M(3). Atomic coordinates for these two models were derived, using the DLS program OPT-DIS written by W. A. Dollase of the University of California at Los Angeles. Subsequent cycles of ORFLS refinement indicated the model with M(2) smaller than M(3) to be much better than the reverse model (wR = 8.9% versus 14.0%).

There is only one independent tetrahedral site in the parent space group C2/m. Although there are two independent tetrahedra in space group C2, mean T—O distances for both proved to differ by only 0.01 A, so that any ordering of Si, Al must be very small. The twofold axis passing through the 2:1 layer in space group C2 relates tetrahedra on one side of the octahedral sheet to tetrahedra on the other side. It is implicit in space group C2, therefore, that both tetrahedral sheets must be identical. To determine whether tetrahedral ordering is present in still lower symmetry or whether both tetrahedral sheets are the same in composition, four models representing all possible tetrahedral ordering patterns were postulated for triclinic C1 symmetry. The four models were tested with the four-quadrant intensity data of set one, and each converged back from the postulated ordered atomic coordinates to the more disordered structure of C2 symmetry with equivalent tetrahedral sheets. For all refinements involving noncentrosymmetric space groups, one atom was fixed in position.

Final refinement of the best C2 model and of the C2/m average structure was accomplished with the second data set in two stages, first using only 633 independent reflections with $2\theta \le 55^{\circ}$ from two quadrants and incorporating absorption corrections. Initially, pseudosymmetry-related atomic coordinates were varied independently in two sets until the final stages of refinement. Then a full matrix refinement was used. No significant differences in atomic positions resulted from this refinement relative to that using the first data set, so final refinement consisted primarily of adjustment of isotropic temperature factors (B) and of scattering factors (f) for the octahedral cations, as determined by alternating cycles of electron-density difference maps and leastsquares refinement. The f and B values proved interrelated in that an increase of one electron in octahedral scattering power would increase B for that same site by approximately 0.25, and the reverse for a decrease. Refinement was terminated for both models when f values that gave flat difference maps also satisfied the total octahedral-site composition for half-ionized atoms and were correlated with reasonable isotropic B values. Unweighted R for the isotropic C2 ordered model was 6.1 percent and for the average C2/m structure was 12.2 percent at this stage. In the second stage of final refinement, both the C2 ordered model and the C2/m average structure were refined with anisotropic temperature factors, using the second data set out to $2\theta \le 90^{\circ}$. Fifty-seven reflections known to be associated with strong white-

Table 2. Observed and calculated structure amplitudes

H E 10F0 10FC	H K 10F0 10FC	H & 10FO 10FC	н к 1050 1050	M K 10F0 10FC	H K 10F0 10FC	H E 10F0 10FC	H K 10FO 10FC
0 69 55 6 0 104 105 8 0 69 55 6 0 104 105 8 0 108 105 10 0 188 185 5 1 168 171 6 2 123 1 168 177 1 147 151 6 2 252 1 168 177 1 147 151 6 2 252 1 168 177 1 147 151 6 2 252 1 168 177 1 147 151 6 2 252 1 168 177 1 147 151 6 2 252 1 168 177 1 147 151 6 2 252 1 168 177 1 147 151 6 2 252 1 168 177 1 147 151 6 1 177 1 178 187 1 178 177 1 178 187 1 178 177 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178 178 1 178		-7 9 72 99 92 92 94 92 94 92 94 94 94 94 94 94 94 94 94 94 94 94 94	0 10 81 83 82 92 92 92 92 92 92 92 92 92 92 92 92 92	-1 11 102 103 3 11 107 113 3 11 109 113 3 11 109 113 5 11 74 57 5 11 74 57 5 11 70 179 113 5 11 70 179 5 11 204 196 6 12 204 233 6 12 246 233 6 12 246 236 6 12 13 10 2 6 13 10 2 6 13 10 2 6 13 10 2 6 14 144 6 17 10 5 6 14 70 6	4 14 181 201 4 14 181 201 4 14 181 201 4 14 181 201 4 18 198 66 83 5 13 18 16 66 83 5 13 18 16 18 5 18 18 18 5 18 18 18 5 18 18 5 18 18 5 18 18 18		-1 7 193 146 -3 7 180 183 -3 7 180 183 -3 7 180 183 -5 7 180 183 -7 7 95 79 -9 7 81 184 -7 7 95 79 -9 7 81 184 -8 180 183 -2 8 184 184 -8 180 1103 -1 19 184 -1 9 92 106 -8 8 110 110 -3 9 184 178 -1 9 92 106 -3 9 184 178 -1 9 92 106 -3 9 184 178 -1 9 92 106 -3 9 184 178 -1 9 92 106 -3 9 184 178 -1 9 92 106 -3 9 184 178 -1 9 92 106 -3 9 184 178 -1 9 92 106 -3 9 118 199 -7 9 129 129 -7 9 129 129 -7 9 129 129 -7 9 129 129 -7 9 129 129 -7 9 129 129 -7 9 129 129 -7 9 129 129 -7 9 129 129 -7 9 129 129 -7 9 129 129 -7 9 129 129 -7 9 129 129 -7 129 129
7 13 39 66 0 14 208 199 4 14 199 119 5 14 199 119 5 14 199 119 5 19 127 129 5 19 127 129 5 19 127 129 5 19 127 129 5 19 127 129 5 19 127 129 5 19 127 129 5 19 127 129 5 19 127 129 5 19 127 129 5 19 127 129 5 19 129 6 12	2 0 1047 1067 -4 0 2040 2040 -6 0 2040 2040 -6 0 2027 2732 -6 0 0 400 600 -6 0 722 732 732 -6 0 400 600 -1 0 0 70 62 -1 1 666 641 -1 1 666 641 -1 1 666 641 -1 1 100 195 -7 1 124 122 -6 2 163 147 -7 1 124 122 -6 2 163 147 -7 1 124 123 -8 2 164 135 -8 2 164 135 -8 2 164 135 -8 2 164 135 -8 2 164 135 -9 3 2040 251 -9 3 100 625 -9 3 100 625 -9 3 100 100 -1 1 100 105 -1 100 105	-1 1 849 953 -1 100 106 -3 1 100 106 -3 1 799 121 -3 1 799 121 -3 1 129 121 -5 1 129 121 -7 1 12		7 1 160 154 7 2 2 12 210 0 4 2 207 22 210 0 6 4 209 22 210 0 6 4 209 22 210 0 6 4 209 22 210 0 6 4 209 22 210 0 6 4 209 22 210 1 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	-3 5 227 249 -5 5 71 83 -7 5 277 84 -7 5 6 84 -7 5 6 84 -7 5 6 84 -7 6 84 -7 1 83 -7 5 6 84 -7 1 83 -7 1 83 -7 1 83 -7 1 83 -8 1 83 -8 1 83 -9 1 93 -9	-2 12 972 989 -4 12 79 91 -4 12 120 102 -4 12 120 103 -4 12 120 103 -4 12 120 103 -1 13 155 128 -1 13 156 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 112 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 13 168 12 -3 17 18 18 -3 17 18 18 -3 17 18 18 -3 17 18 18 -3 17 18 18 -3 17 18 18 -3 17 18 18 -3 18 18 -3 18 -3 18 -3	-7 1 205 200 -7 1 91 98 -2 2 74 98 -6 2 275 276 -6 2 275 277 -6 2 2 75 97 -6 2 2 75 97 -6 2 2 75 97 -7 1 3 661 709 -6 4 133 132 1382 -7 3 3 62 44 -8 12 13 661 709 -6 4 133 132 1382 -7 3 641 709 -6 4 133 132 1382 -7 3 6 7 13 13 129 -7 3 6 7 13 13 129 -7 3 7 13 13 129 -7 3 7 13 13 13 13 13 13 13 13 13 13 13 13 13

Table 2, continued.

н	·	10F0	10FC			torn	torr	н	ŧ	10F0	torc	н	ĸ	10F0	10F
3	13	78 113	103	-5	6	197	208	3	5	92	67.		١.	16	
1	15	222	244	6	6	119	107	-3 -5	5	174	156	7	0	93	9
~1 ~3	15 15 15	177	197	-1	7	165	157	2	6	148	142 98	-2	0	111	11
-2	16	86	75	-1	7	108	84	-2	6	8.5	80 138	-6 -1	0	128	13
	L	10		5	7	70	76 71	-4 -6	6	135 292 250	310	-3	1	116	11
				0	8	274	270	-6 -8	6	109	99	-5	1	61 87	7
-4	0	394 122	392	-2	8	179	158	-1	7	132	34 123	.0	2	106	9
-6 8	0	337	130 297 245	-8	8	139	51	-5	7	105	83	- 6	2	142	11
-8	0	120	113	-1	9	63	125 72	-7	7	103	89 107	-1	3	201	20
-1	1	86	84	-3	9	205	203	-4	8	95 189	88 187	-3	3	172 95	16
-3 -3	1	180	197	-5	9	279	274	-6	8 9	121	116	0	4	62	13
-5	1	112	117	-7	9	272	273	-1	9	376	338	44	4	66	5.
7	1	76	63 71	0 2	10	131	144	-3	9	111	117 93	-1 -3	5	81 95	10
-7 -9	1	137	124	-2	10	79 101	92	-6	10	106	88	~2	6	115	11
0	2	155 273	165 261	-3	10 11 11	101 192 71	186	-1	11	103	102 95 101	-6	6	81	9
-2	2	152	156	0	1.2	256	258	3	11	67	70	3	7	85	12
-4	2	81	119 67	-2	12	192	187 143	-5	11	109 129 74	95 130	-3	7	84	61
-8 1	2	91 195	88	-6	12	75 66	77	2	12	173	68 166	-1	9	170	
3 - 3	3	512	525 103	-1	13	66 76 108	63 70			1,,	100	-3	9	139	15
5	- 3	273	300	0	14	174	120		L.	14		-5	10	110	10
-5 -7	3	73 100	76 112	5	14	137	147	2	0	128	119	-4	10	101	70
-9 0	3	199	193		L =	12		-2	0	381 436	381 475		1	17	
-2	4	209 83	202	2	0	559	564	-6 1	0	68	93	2	0		
4	4	181	214	-2	0	152	151	-1	1	115	48	-2	0	180	114
6	4	184	71	-4 -4	0	351 56	365 47	-3	1	134	136 92	-6	0	180 251 115	10
-8 1	5	164	157	-6	0	131 356	134	-5	1	96 93	97 110	-5	1	81	7
-1	5 5	73	74	-6	0	242	226	-2	2	94	98	-7	1	83	6
5	5	79	8.8	-1	1	120	47	-2	2	104	97 116	-4	3	99 283	27
7	5	98	86	-3	1	120 106 121	98	-6 -8	2	150	141	-1	3	70	61
9	5	108	106	5	1	71	56	1	3	174	176	2		64	22
2	6	70	412 87	0	2	70	67 66	-3	3	174 153 159	144	-6	*	93 93	7
-2	6	389	374 94	-2	2 2	297 177	312	-7	3	185	137	0 2	6	139	130
6	6	165	94 455 180	-1 3	3	480 187	498	-2	*	130	106	-2	6	190	170
-6	6	246	256	-3 -5	3	486	514	4	+	6-8	110	-6	6 7	170 153	145
1	7	152 65	148	-7	3	228 60	216 36	-6 -8	2	91 84	89 88	-3 -9	7	106	87 72 99
3	7	101 93 76	101	2	4	70	219 56	3	5	129	114	-2		103	99
-5	7	76 80	79 77	-2	4	179	199 58	-3 5	5	83 73	85		•	10.	
0	8	74 175	65	-4	4	180	164	-7	5	75 150	74 149			18	
4	8	221	205	-6 -1	5	167	124 165	-2	6	388	366	-2	0	291	282
-6	8	60	122	-3	5	133	124 112	-4	6	294 129	297 134	-1	0	85 90	98 88
1	9	260	238	0	6	181 314	176 313	-3	7	104	76 90	0	5	75 78	53
3	9	370	336 125	-6	6	67	72 173	0	.0	60	51 116	-2	2	76	23
5	9	126 207	199	-8	6	257	255 75	1.	9	96	103	-1	3	74 110	63 98
-7	9	70 76	90 83	-1	7	80 78	65	-1 3	9	101	116	-3	3	68 191	26 168
2	10	111	100	-3	7	147	145	-3	10	116	146	-1	5	6B 87	89
4	10	83 92	103 83	-2	8	145 120 208	143 119 223	-4	10	78	62	0	6	239	233
3	11	145	140	-4	8	104	114	-6	1.1	129 85	106	-2	6	798 99	271
2	12	62	175	-8	9	66	43 51	-3	11	134	108	-1	7	8.6	76
-2	12	299 247	272 253	-1 -3	9	306	306 392	-2	12	206	227		4 .	19	
6	12	164	178 76	-5	9	143	149		0.5						
1	15	114	127	0	10	76	73		La	15		1	1	134 83 71	110
1	15	93	80	-2 4	10	184	187	2	0	206	199	-5 -1	3	76	77 96 67
	L =	11		-4	10	149	153	-2	0	177	169	-3	3	251	230
į.	0	466	470	1	11	81 106	53 97	1	1	151	144	9		41	54
2 2 4	0	155	154 124	-1	11	148	116	-1	1	168 86	88				
4	0	59	61	2	12	152 260	155 271	-5 0	2	78 191	138				
6	0	122	119	0	14	84	52 182	2	2	70 111	50 100				
1	1	165	171	- 2	7.44		145	1	3	121	112 153				
3	1	193	184		١.	13		-1 -3	3	180 206	186				
0	2	235	98 232 272	4	0		233	-5 -7	3	98 310	99 324				
2	2	274	272	-4	0	231 231 128	233 277 131	-7 2	3	237	233				
4	2	80	65	-6	0	341	314	-4	4	94	103				
6	2 2 2 2 3	120 63	126	-8 1	0	162 77	163 65 79	-6 -8	4 5	132	122 80				
4 6 6 8	2	115	69	-1	1	57 98	79 104	1	5	115	103				
1	3	173	181	-3 -5 -7	1	103	6.8	-1 3	5	99	101				
1 3 3	3	177	175	-7	1	136	133 71	-3 -5 -7	5 5	70 83	67				
9	3	395 334	395 352	0 2	2	132	138 73	0	6	139	157				
9 9	3	334 453 184	218	-2	2	121	121	-2	6	243 206	228				
7	3	281 146	291 151	-4	2 2 2 2	117	132		6	169	154				
0	4	250	257	-6 -8	2	119 78	115	~6 1	7	72 129	81				
2 2 4	4	159	170	-1	3	461	420 557	-1	7	132	134				
	4	66 81	74	3	3	172	140	-7 -7	7	85	63				
4	4 5	105	97	-3	3	152 79	176	-6	8	74 82	73 107				
4		181	181	-9 -9	3	114	104	-1	9	105	83				
4	5	173							2						
4 8 1 1 1	5	59 70	73	4	4	64	171		9	77	168				
4	5 5 5	59	61	6 -6	4 4 4 5	167 76 132	171 70 127 67	-3 -5 -2	9 9 10 11	77 198 219 76	168 234 31 56				

radiation streaks were deleted from this data set for a total of 1,493 reflections. Because the octahedral coordination of the refined structure was known to be noticeably asymmetric, all of the variable parameters were varied together in one set. No significant correlation effects were indicated, and there were no significant changes in atomic coordinates. The final unweighted and weighted R values were 5.7 and 6.3 percent respectively for the C2 ordered model, and 12.5 and 17.4 percent for the average C2/m structure. Hamilton's (1965) residual-ratio test indicates that the C2 model represents significant improvement over the C2/m model at better than the 1 percent significance level. Tables 2-4 list observed and calculated structure amplitudes based on the anisotropic refinement results in C2 symmetry, final atomic coordinates, and calculated bond lengths and angles.

Tetrahedral ordering

The mean T—O bond lengths for tetrahedra T(1)and T(11) are 1.646 A and 1.639 A, respectively. The standard deviations of the bond lengths (Table 4) were calculated, using the correlation matrix as input to program ORFFE. The standard deviation σ_l of an individual T—O bond length is 0.004 A, the standard deviation of the mean value is $\sigma_n = \sigma_l / \sqrt{n} = 0.002 \text{ A}$, and the standard deviation in their difference is $\sigma =$ $\sqrt{2}\sigma_n = 0.0028$ A. The observed difference of 0.007 A between the two mean T—O values equals 2.50σ for this accuracy, and is just significant at the 1 percent level. From the regression analysis of Hazen and Burnham (1973) for micas, the mean T—O values correspond to tetrahedral contents of 0.233 Al and 0.190 Al in T(1) and T(11), respectively. This gives a total of 0.85 Al^{IV} per formula unit, in close agreement with the composition Si_{3.09}Al_{0.91} determined by microprobe analysis. Thus, a small but statistically significant ordering of tetrahedral cations is indicated by the refinement.

Octahedral ordering

The most important structural feature of zinnwaldite is the unusual octahedral ordering pattern. The mean M—O,F values for M(1), M(2), and M(3) are 2.132, 1.882, and 2.131 A, and the scattering powers are 15.0, 11.5, and 13.5 electrons, respectively. The reduction in symmetry from C2/m to C2, therefore, is due primarily to the difference in sizes and, to a lesser extent, in scattering powers of the M(2) and M(3) cations across the pseudo-mirror plane (Fig. 1). M(1) and M(3) are identical in size and have a difference in scattering power that is probably significant.

TT 1 1					
lable	3. F	ınal	atomic	e para	ameters

Atom	×	g	z	Bequiv.	β11	β22	β33	β12	β13	β23
K(1)	0	0.5028(2)	0	1.93	0.0181(4)	0.0048(1)	0.0054(1)	0.0	0.0013(2)	0.0
M(1)	0	-0.0069(3)	1/2	0.64	0.0067(3)	0.0009(1)	0.00236(8)	0.0	0.0013(1)	0.0
M(2)	0	0.3217(3)	1/2	0.57	0.0050(4)	0.0014(1)	0.0017(1)	0.0	0.0006(2)	0.0
M(3)	1/2	0.1631(3)	1/2	0.62	0.0050(4)	0.0012(1)	0.0023(1)	0.0	0.0000(2)	0.0
r(1)	0.0745(2)	0.1688(3)	0,2276(1)	0.52	0.0053(3)	0.0010(1)	0.0016(1)	0.0003(1)	0.0003(1)	-0.00015(7
r(11)	0.5844(2)	0.3323(3)	0.2275(1)	0.54	0.0044(3)	0.0009(1)	0.0022(1)	0.0002(1)	0.0006(1)	-0.00003(8
7/1)	0.0289(6)	-0.0002(5)	0,1720(3)	1.40	0.0202(9)	0.0021(2)	0.0030(2)	-0.0000(6)	0.0002(3)	-0.0002(3)
0(1)	0.0289(8)	0.2350(5)	0.1725(4)	1.42	0.0125(9)	0.0053(3)	0.0029(3)	-0.0031(5)	0.0019(4)	-0.0005(3)
0(2) 0(22)	0.8177(7)	0.2641(5)	0.1723(4)	1.34	0.0117(9)	0.0046(3)	0.0031(3)	0.0028(4)	0.0010(4)	0.0007(2)
0(3)	0.1155(6)	0.1748(4)	0.3939(3)	0.71	0.0067(7)	0.0019(2)	0.0019(2)	0.0011(3)	0.0008(3)	-0.0001(2)
0(33)	0.6639(5)	0.3271(5)	0.3928(3)	0.75	0.0058(6)	0.0022(2)	0.0022(2)	0.0002(4)	0.0003(3)	0.0000(2)
F(1)	0.1089(5)	0.4715(3)	0.3989(3)	1.10	0.0108(7)	0.0025(2)	0.0032(2)	-0.0007(3)	0.0009(3)	0.0005(2

 B equiv. calculated from anisotropic data according to Hamilton (1959). The anisotropic temperature factor form is exp $(-\Sigma_{i}\Sigma_{j}\beta_{i}_{j},h_{j}h_{j})$.

The size and scattering power of M(2) are consistent with complete ordering of Al into this site. To a first approximation, the remaining octahedral cations and vacancies of the formula unit can be considered randomly distributed over M(1) and M(3). For this random distribution the distance of the average cation from its oxygen neighbors, neglecting the (F,OH) contacts, was calculated from the radii of Shannon and Prewitt (1969, 1970) as 2.133 A. The size of a vacancy was taken as 0.80 A in this calculation. The observed values from Table 4 are 2.133 A for mean M(1)—O and 2.117 A for mean M(3)—O. A better overall fit for calculated sizes and electron counts is obtained by moving all the smaller and lighter Al out of M(1) into M(3) and replacing with a corresponding amount of Fe^{2+} . The calculated M—O values then become 2.140 A for M(1) and 2.127 A for M(3), and the calculated electron counts become 13.9 and 13.1, respectively. Despite the good fit, this should not be considered a unique solution.

The (F,OH) atom has a high temperature factor ($B = 3.0 \text{ A}^2$) when refined in C2/m symmetry. Figure 1 shows that in C2 symmetry the (F,OH) atom has moved off the mirror plane to coordinate more closely with Al in M(2). At the same time the equivalent isotropic B value for (F,OH) has decreased to 1.1 A² (Table 3). The M(3)—(F,OH) distance of 2.159 A (Table 4) is larger than the M(3)—O distances (2.109, 2.125 A). This octahedral irregularity, to be discussed in more detail below, was the reason for using only the oxygen contacts in calculating the octahedral cation distributions mentioned in the preceding paragraph.

The (F,OH) atom is located in the trans-orienta-

tion at opposite apices of the M(1) octahedron but in the cis-orientation along a shared edge between the M(2) and M(3) octahedra (Fig. 1). Because fluorine is smaller than oxygen, two (F,OH) atoms can approach each other more closely along the shared edge between M(2) and M(3) than can two oxygens along any M(1) shared edge. Thus, the smaller Al can be accommodated more readily in either M(2) or M(3)than in M(1). The small amount of tetrahedral ordering present in this crystal also favors location of a high-charge cation in M(2). The shared octahedral edge between M(2) and M(1) that parallels the F—F shared edge between M(2) and M(3) involves two O(3) atoms apical to T(1) cations in the upper and lower tetrahedral sheets (Fig. 1). Because of Al concentration in T(1), the apical O(3) anions are undersaturated and thus favor a trivalent cation in M(2). Reversal of the tetrahedral ordering pattern by placing more Al^{IV} in T(11) would favor octahedral Al in M(3) instead of M(2), and the particular pattern adopted may be a random choice from crystal to crystal. This also may lead to domain structures. It is anticipated that the O-H dipole will be deflected from the sheet normal to point away from M(2) in a direction between M(1) and M(3) and toward the undersaturated O(3)—O(3) shared edge of the next M(2) octahedron.

Ordering model

Rieder (1968) was correct in predicting that zinnwaldite would show octahedral ordering, although the observed ordering pattern is rotated 120° from that predicted. Rieder emphasized the smaller d(001) values of natural zinnwaldites relative to synthetic

Table 4. Calculated bond lengths and angles

В	ond length	Bond angles (°)			
		Tetrahedi	on T(1)		
0(1) 0(2) 0(22) 0(3) Mean	1.646(4) 1.640(4) 1.646(4) 1.653(4) 1.646	0(1)0(2) 0(22) 0(3) 0(2)0(22) 0(3) 0(22)0(3) Mean	2.655(5) 2.653(5) 2.721(4) 2.664(6) 2.727(5) 2.706(5) 2.688	0(1)0(2) 0(22) 0(3) 0(2)0(22) 0(3) 0(22)0(3) Mean	107.8(2) 107.4(2) 111.1(2) 108.3(2) 111.8(2) 110.2(2) 109.4
0(1) 0(2) 0(22) 0(33) Mean	1.637(4) 1.639(4) 1.638(4) 1.643(3) 1.639	Tetrahedr 0(1)0(2) 0(22) 0(33) 0(2)0(22) 0(33) 0(22)0(33) Mean		0(1)0(2) 0(22) 0(33) 0(2)0(22) 0(33) 0(22)0(33) Mean	107.9(2 108.2(2) 111.7(2) 107.2(2) 111.4(2) 110.2(2) 109.4
Inter 0(1)x2 0(2)x2 0(22)x2 Mean	layer cati 2.994(3) 2.999(4) 2.978(4) 2.990 (inner)	3.286(3) 3.291(5) 3.177(4) 3.251 (outer)		T(1) to around O(1) around O(2) around O(22) Mean	T(11) 139.1(2) 139.4(2) 130.5(2) 136.3
		Octahedr	on M(1)		
0(3)x2 0(33)x2 F(1)x2 Mean	2.127(4) 2.138(4) 2.130(3) 2.132		0(3	F(1)x2 F(1)x2 (3)F(1)x2 (unshared)	96.8(1) 100.0(1) 97.9(1) 98.2
0(3)0(33)x2 F(1)x2 0(33)F(1)x2 Mean (unshared)		3.261(4) 3.219(4) 3.223		F(1)x2 (3)0(33) F(1)x2	77.3(2) 88.3(1) 89.6(2) 74.4(1)
0(33)0	1)x2 (33) 1)x2	2.656(6) 2.967(4) 3.013(6) 2.579(4) 2.794	Mea	n (shared)	82.1
0(3)x2 0(33)x2 F(1)x2 Mean	1.889(4) 1.895(3) 1.862(3) 1.882	Octahedr	0(3	F(1)x2 F(1)x2 I3)F(1)x2 In (unshared)	93.5(1) 92.6(1) 91.1(1) 92.4
0(33)F Mean (un	1)x2 (1)x2 shared)	2.755(4) 2.712(3) 2.681(4) 2.716	0(3 F(1	0(33)x2 0(33)x2 (3)F(1)x2 .)F(1) un (shared)	89.4(2) 88.6(1) 86.7(1) 85.4(2)
0(3)0(0(0(33)F F(1)F(Mean (sh	33)x2 (1)x2 1)	2.656(6) 2.643(4) 2.579(4) 2.524(5) 2.604			
0/21-2	2 100/2)	Octahedi		0 (22)2	00 6/1
0(3)x2 0(33)x2 F(1)x2 Mean	2.109(3) 2.125(4) 2.159(3) 2.131		0 (3	F(1)x2 F(1)x2 F(1)x2 In (unshared)	98.6(1) 96.7(1) 99.3(1) 98.2
0(3)0(F(0(33)F Mean (un	1)x2 (1)x2	3.209(4) 3.189(4) 3.265(4) 3.221	0(3 F(1	F(1)x2 F(1)x2 G3)0(33) .)F(1) In (shared)	77.3(1) 88.1(1) 90.3(2) 71.5(1) 82.1
0(3)0(F(0(33)0 F(1)F(Mean (sh	1)×2 (33) 1)	2.643(4) 2.967(4) 3.013(6) 2.524(5) 2.793			

specimens as evidence for ordering, and also noted that the natural specimens were more F-rich than the synthetic specimens, which were grown in equilibrium with a buffer of low fluorine fugacity. The difference of 0.1 A noted in the d(001) values by Rieder

probably is due primarily to the presence of fluorine. Calculation of the octahedral sheet thickness in the present study by considering the apical oxygens separately from the (F,OH) atom does result in a difference of 0.1 A. In addition, Yoder and Eugster (1954) showed a decrease in basal spacing of 0.18 A for fluorophlogopite relative to hydroxyphlogopite, where octahedral ordering cannot be a factor. It should be noted also that the clintonite-1*M* (xanthophyllite), the model for the Rieder ordering pattern, actually has an octahedral sheet not appreciably different in thickness (2.148 A) than for disordered F-poor trioctahedral micas (range 2.09–2.22 A).

We suggest that the ordering pattern observed in this study, namely with the small cation in either M(2) or M(3), should be found in all F-rich zinnwaldites and lepidolites. Some evidence for this view can be found in the anomalously high B values for the (F, OH) atom in micas that show the "normal" ordering pattern, with M(1) larger than the average M(2,3)atom, but that have not been refined in subgroup symmetry, namely $B = 2.78 \text{ A}^2$ in fluor-polylithionite-1M (Takeda and Burnham, 1969), B = 2.92and 1.77 A² in lepidolite- $2M_2$ (Takeda et al., 1971; Sartori et al., 1973), and $B = 2.32 \text{ A}^2$ in lepidolite-1M (Sartori, 1976). Additional evidence comes from the cell dimensions of zinnwaldites. Bailey (1975) has shown that the "normal" ordering pattern necessarily gives rise to an intralayer shift larger than the ideal value of -a/3 and, unless compensated by an offset of adjacent layers, to an observed β angle for 1Mmicas larger than ideal. It is possible to obtain values of $c\sin\beta$, a, and β for 11 zinnwaldites from graphs presented by Rieder (1968, 1970b) and Rieder et al. (1971). In all cases the "normal" ordering pattern is predicted, including specimen # 40 whose structure is presented here ($\beta_{obs} = 100.83^{\circ}$, $\beta_{ideal} = 100.07^{\circ}$). For three synthetic annites the observed and ideal β angles are approximately equal, as should be true where the octahedra are of similar sizes.

Structural distortions

Important structural features of zinnwaldite-1M are summarized in Table 5. Tetrahedra T(1) and T(11) are similar in shape, only slightly elongate, and rotated by 5.8°. The small octahedron M(2) containing Al is nearly regular in shape, but M(1) and M(3) are considerably flattened and distorted (Fig. 1). The individual ψ values of 60.8° (ideal = 54.73°) for M(1) and M(3) are among the largest recorded for micas to date, and are a measure of the amount of flattening required to fit these large Fe,Li-rich octahedra onto

Table 5. Important structural features of zinnwaldite

Parameter	Value		
^a α _{tet} (°)	5.8		
tet	5.8		
b _{ttet} (°)	T(1): 111.0 T(11): 111.1		
^e β _{ideal} (°)	100.07		
oct(°)	M(1), M(3): 60.8 M(2): 56.5 Mean: 59.5		
dSheet thickness (Å)			
tetrahedral Interlayer separation (Å) Basal oxygen Δz_{aye} (Å)	2.252 2.078 3.333 0.124		
Intralayer shift	-0.354a ₁		
Layer offset	-0.004a,		
Resultant shift	-0.358a ₁		

 $^{^{}a}$ Tetrahedral rotation is calculated from α = 1/2|120° - mean $^{O}_{D}$ - $^{O}_{D}$ - $^{O}_{D}$ angle | .

the adjacent smaller and thinner octahedron ($\psi = 56.5^{\circ}$) around the Al in M(2). The sheet thicknesses are similar to those found in other F-rich trioctahedral micas.

Apparent thermal vibrations

Tables 6 and 7 list the orientations of the thermal ellipsoids and the calculated bond lengths after correction for thermal effects. The apparent thermal motions of zinnwaldite resemble those of phlogopite (and annite) in some aspects and are unique in others. Like phlogopite, the apparent thermal motions are large in both magnitude and anisotropy. Hazen and Burnham (1973) suggest that such apparent thermal motions are due, in part, to local variations in atomic positions. Also like phlogopite, equivalent isotropic temperature factors of the apical oxygens [O(3) and O(33)] are considerably smaller in magnitude than those of the basal oxygens. These differences in thermal magnitudes may be related to the differences in bond lengths between tetrahedral Al and Si and to lack of appreciable tetrahedral order. The basal oxygens are coordinated to two tetrahedra and are, therefore, more affected by the different Si-O and Al¹V—O bond lengths than are the apical oxygens.

In phlogopite most of the atoms are elongate along

 Z^* . In zinnwaldite the elongation is along Z^* and Z. However, some atoms also appear to be elongate along the bonds to them. Most significantly, the (F,OH) atom is elongate along the bonds to both M(1) and M(3), and this suggests a small amount of positional disorder for (F,OH). NMR studies of phlogopites with varying amounts of F and OH indicate (1) a tendency for F—F pairing along the same octahedral edge rather than a random distribution. and (2) preferential location of Fe²⁺ in sites close to OH groups (Sanz and Stone, 1977.) This is in accord also with the known preference of Al for F rather than for OH, as summarized by Kampf (1977). It is interesting to note in the Sadisdorf zinnwaldite specimen that OH makes up 39 percent of the F,OH total and that Fe²⁺ also makes up 39 percent of the total occupancy of M(1) + M(3). This may indicate a local clustering or domain structure, in which 39 percent of the volume of the crystal consists of OH-OH pairs located closer to Fe^{2+} in M(1) and M(3) than to Al in M(2), and 61 percent consists of F—F pairs located

Table 6. Orientations of thermal ellipsoids relative to crystal axes

Atom	Axis	rms (Å) displacement	Angle X	(°) with res	pect to Z
K(1)	r ₁	0.143(2)	90	0	90
	r ₂	0.156(2)	150(6)	90	109(6)
	r ₃	0.168(2)	120(6)	90	19(6)
M(1)	r ₁	0.062(4°)	90	0	90
	r ₂	0.091(2)	163(5)	90	62(5)
	r ₃	0.110(2)	73(5)	90	28(5)
M(2)	rl	0.078(4)	90	0	90
	r2	0.082(3)	178(13)	90	77(13)
	r3	0.093(4)	88(13)	90	13(13)
M(3)	r ₁	0.071(5)	90	0	90
	r ₂	0.080(3)	156(5)	90	103(5)
	r ₃	0.110(3)	114(5)	90	13(5)
T(1)	r ₁	0.063(4)	102(5)	14(4)	80(4)
	r ₂	0.084(3)	145(11)	95(6)	113(12)
	r ₃	0.094(3)	122(11)	103(4)	25(11)
T(1))	rl	0.061(4)	100(8)	10(8)	88(3)
	r2	0.078(3)	169(7)	100(8)	84(4)
	r3	0.104(3)	94(4)	91(3)	7(4)
0(1)	r ₁	0.095(5)	89(4)	7(13)	83(12)
	r ₂	0.121(4)	99(3)	83(13)	159(5)
	r ₃	0.171(4)	9(3)	89(4)	110(3)
0(2)	r ₁	0.106(6)	145(7)	118(9)	62(21)
	r ₂	0.117(6)	99(19)	111(11)	151(21)
	r ₃	0.167(5)	123(4)	36(4)	97(5)
0(22)	r ₁	0.102(6)	132(7)	53(4)	106(10)
	r ₂	0.127(6)	61(9)	94(9)	161(10)
	r ₃	0.157(5)	56(5)	37(4)	82(7)
0(3)	r ₁	0.076(6)	130(8)	41(8)	76(12)
	r ₂	0.098(5)	96(22)	86(18)	162(18)
	r ₃	0.106(5)	40(9)	50(8)	100(25)
0(33)	r ₁	0.087(5)	26(18)	106(26)	81(11)
	r ₂	0.097(5)	105(26)	164(27)	93(27)
	r ₃	0.107(5)	110(12)	90(26)	9(12)
F(1)	r ₁	0.096(5)	72(7)	25(5)	110(5)
	r ₂	0.121(4)	145(12)	84(9)	113(13)
	r ₃	0.134(4)	118(13)	66(5)	31(10)

 $[^]b$ The tetrahedral angle is defined as τ = 0 $_{\rm apical}$ ^-T--0 $_{\rm basal}$. The ideal value is 109.47°.

 $^{^{}C}$ The mean octahedral angle, ideally 54.73°, is calculated from cos ψ = oct. thickness/2(M--O,F,OH).

 $^{^{}d}$ Includes the position of F(1) in the calculation.

 $e_{\beta_{ideal}} = 180^{\circ} - \cos^{-1}(\underline{a}/3\underline{c}).$

Table 7. Bond lengths corrected for thermal motion

	Distar	ice (Å)		Distanc	e (Å)
Bond	Correl high	ation none	Bond	Correla high	ntion none
		M(1) octa	hedron		
0(3)x2	2.128(4)	2.136(4)			
0(33)x2 F(1)x2	2.140(4)	2.148(4)			
	2.142(3)	2.149(3)			
Mean	2.137	2.144			
0(3)-0(3)	2.660(6)	2.667(6)	0(3)-0(33)x2	3.189(4)	3.195(4)
F(1)x2 O(33)-O(33)	2.971(4) 3.019(6)	2.977(4) 3.025(6)	F(1)x2 O(33)-F(1)x2	3.275(4) 3.220(4)	3.279(4) 3.226(4)
F(1)x2	2.583(4)	2.590(4)			
Mean	2.798	2.804	Mean (unshared)	3.228	3.233
(shared)			,,		
		M(2) octa	hedron		
)(3)x2	1.890(4)	1.898(4)			
0(33)x2	1.896(3)	1.904(3)			
F(1)x2	1.866(3)	1.874(3)			
Mean	1.884	1.892			
0(3)-0(3)	2.660(6)	2.667(6)	0(3)-0(33)x2	2.755(4)	2.762(4)
0(33)x2	2.645(4)	2.652(4)	F(1)x2	2.714(3)	2.721(3)
0(33)-F(1)x2 F(1)-F(1)	2.583(4) 2.527(5)	2.590(4) 2.537(5)	0(33)-F(1)x2	2.683(4)	2.690(4)
Mean	2.607	2.615	Mean (unshared)	2.717	2.724
(shared)	2.007	2.013	(dishared)		
		M(3) octa	hedron		
O(3)x2	2.109(3)	2.117(3)			
0(33)x2	2.126(4)	2.133(4)			
F(1)x2	2.161(3)	2.169(3)			
Mean	2.132	2.140			
$0(3)-0(33) \times 2$	2.645(4)	2.652(4)	0(3)-0(33)x2	3.210(4)	3.215(4)
F(1)x2 O(33)-O(33)	2.971(4) 3.019(6)	2.977(4) 3.025(6)	F(1)x2 0(33)-F(1)x2	3.189(4) 3.267(4)	3.196(4)
F(1)-F(1)	2.527(5)	2.537(5)	Mean	3.222	3.228
Mean	2.796	2.803	(unshared)	3.222	J. 220
(shared)					
		T(1) tetr	ahedron		
0(1)	1.655(4)	1.664(4)	0(1)-0(2)	2.660(5)	2.674(5)
O(2) O(22)	1.649(4) 1.654(4)	1.656(4) 1.663(4)	0(22) 0(3)	2.655(5) 2.717(4)	2.670(5) 2.732(4)
0(3)	1.654(4)	1.661(4)	0(2)-0(22)	2.664(6)	2.677(6)
Mean	1.653	1.661	0-(3)	2.723(5)	2.738(5)
			0(22)-0(3)	2.702(5)	2.716(5)
			Mean	2.687	2.701
		T(11) tetr	ahedron		
0(1)	1.645(4)	1.655(4)	0(1)-0(2)	2.650(5)	2.666(5)
0(2) 0(22)	1.647(4) 1.645(4)	1.656(4) 1.654(4)	0(22) 0(33)	2.659(5) 2.711(4)	2.673(5)
0(33)	1.645(3)	1.651(3)	0(2)-0(22)	2.637(6)	2.652(6)
Mean	1.646	1.654	0(33)	2.708(5)	2.722(5)
			0(22)-0(33)	2.688(5)	2.701(5)
			Mean	2.676	2.690
		erlayer ca	tion $K(1)$		
0(1)x2 0(2)x2	2.991(3) 2.997(4)	3.007(3) 3.013(4)			
0(22)x2 0(22)x2	2.975(4)	2.991(4)			
Mean (inner)	2.988	3.004			
O(1)x2	3.294(3)	3.307(3)			
0(2)x2	3.292(5)	3.307(5)			
0(22)x2	3.177(4)	3.193(4)			
Mean (outer)	3.254	3.269			

closer to Al in M(2) than to Li and the remaining cations in M(1) and M(3). Under this interpretation the bulk of the thermal ellipsoid of (F,OH) is determined by the positions of the F (61 percent), and the elongation toward M(1) and M(3) is produced by the closer approach of the OH (39 percent) to the Fe in those sites. In addition to the elongation of (F,OH) along the M(1) and M(3) bonds, the M(1) site is observed to be elongate along the M(1)—(F,OH) bond.

Acknowledgments

We thank the donors of the Petroleum Research Fund, administered by the American Chemical Society, for partial support of this research by grant 8425-AC2 and to the National Science Foundation for partial support through grants GA-34918 and EAR76-06620. We also thank Dr. Milan Rieder of Charles University for providing the zinnwaldite crystal used in this study. Mr. Darrell J. Henry of the University of Wisconsin for the electron microprobe analysis, Dr. Wayne A. Dollase of UCLA for use of program OPTDIS, and Dr. Werner H. Baur of the University of Illinois at Chicago Circle for advice on refinement procedures.

References

Bailey, S. W. (1975) Cation ordering and pseudosymmetry in layer silicates, *Am. Mineral.*, 60, 175–187.

Foster, M. D. (1960) Interpretation of the composition of lithium micas. U.S. Geol. Surv. Prof. Pap., 354-E, 115-147.

Guggenheim, S. and S. W. Bailey (1975) Refinement of the margarite structure in subgroup symmetry, *Am. Mineral.*, 60, 1023-1029.

Hamilton, W. C. (1959) On the isotropic temperature factor equivalent to a given anisotropic temperature factor, *Acta Crystallogr.*, 12, 609-610.

—— (1965) Significance tests on the crystallographic *R* factor. *Acta Crystallogr*_{*}, 18, 502–510.

Hazen, R. M. and C. W. Burnham (1973) The crystal structures of one-layer phlogopite and annite. Am. Mineral., 58, 889-900.

Kampf, A. R. (1977) Minyulite: its atomic arrangement. *Am. Mineral.*, 62, 256-262.

Rieder, M. (1968) Zinnwaldite: octahedral ordering in lithium-iron micas. Science, 160, 1338–1340.

— (1970a) Chemical composition and physical properties of lithium-iron micas from the Krušné hory Mts. (Erzgebirge), Part A. Chemical composition. *Contrib. Mineral. Petrol.*, 27, 131-158.

—— (1970b) Lithium-iron micas from the Krušné hory Mountains (Erzgebirge): Twins, epitactic overgrowths and polytypes. Z. Kristallogr., 132, 161-184.

——, A. Píchová, M. Fassová, E. Fediuková, and P. Černý (1971) Chemical composition and physical properties of lithiumiron micas from the Krušné hory (Erzgebirge), Czechoslovakia and Germany, Part B: Cell parameters and optical data. *Mineral. Mag.*, 38, 190–196.

Sanz, J. and W. E. E. Stone (1977) Order-disorder in the octahedral layer of phlogopites: an investigation by NMR (abstr.) *Proc. 3rd European Clay Conf.*, *Oslo*, 167-169.

- Sartori, F. (1976) the crystal structure of a 1M lepidolite. Tscher-maks Mineral. Petrogr. Mitt., 23, 65-75.
- —, M. Franzini and S. Merlino (1973) Crystal structure of a $2M_2$ lepidolite. *Acta Crystallogr.*, B29, 573-578.
- Shannon, R. D. and C. T. Prewitt (1969) Effective ionic radii in oxides and fluorides. *Acta Crystallogr.*, *B25*, 925-946.
- —— (1970) Revised values of effective ionic radii. *Acta Crystallogr.*, B26, 1046-1048.
- Sidorenko, O. V., B. B. Zvyagin and S. V. Soboleva (1975) Crystal structure refinement for 1M dioctahedral mica. Soviet Physics-Crystallogr., 20, 332-335 (English transl.).
- Takeda, H. and C. W. Burnham (1969) Fluor-polylithionite: a lithium mica with nearly hexagonal (Si₂O₅)²⁻ ring. *Mineral. J.* (*Japan*), 6, 102-109.
- ——, N. Haga and R. Sandanaga (1971) Structural investigation of polymorphic transition between $2M_2$ -, 1M-lepidolite and $2M_1$ muscovite. *Mineral. J. (Japan)*, 6, 203–215.
- Yoder, H. S. and H. P. Eugster (1954) Phlogopite synthesis and stability range. *Geochim. Cosmochim. Acta*, 6, 157-185.

Manuscript received, February 7, 1977; accepted for publication, June 20, 1977.