Klebelsbergite, Sb₄O₄(OH)₂SO₄: redefinition and synthesis

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Abstract

Klebelsbergite is verified as a valid mineral species, in museum specimens from the type locality, Baia Sprie (Felsöbánya), Rumania.

Klebelsbergite occurs as pale yellow to orange-yellow crystals radially grown on stibnite. Wet-chemical analysis of synthetic klebelsbergite, supported by electron microprobe analyses of natural material, gives Sb 71.78, S 4.65, H_2O 2.69, O 20.88% (by difference). The empirical cell contents are $Sb_{4.07}O_{5.00}(SO_4) \cdot 1.03H_2O$ with an ideal formula $Sb_4O_4(OH)_2SO_4$. The crystals are tabular or acicular elongated [001], flattened {010}. The mineral is orthorhombic, Pcmb or $Pc2_1b$, a = 11.279(2), b = 14.909(3), c = 5.7648(6)A; Z = 4. D(meas) = 4.62(6), D(calc) = 4.69g/cm³. The strongest lines of the X-ray powder pattern are 6.22(s)(120), 3.892(s)(211), 3.545(s)(221), 3.150(s)(301), 3.131(vs)(231), 2.830(s)(012), 2.435(s)(132), 1.805(s)(213). Klebelsbergite has vitreous luster, yellow streak, and no distinct cleavage. The mineral is optically biaxial with large $2V_x$ about 70° , $\beta \cong 1.95$, and is not pleochroic. VHN = 200 (10g load).

Single crystals of klebelsbergite were synthesized by boiling 2 g of Sb_2O_3 with 600 ml of 1 N H_2SO_4 for four days. Synthetic crystals are tabular, elongated [001], flattened {010}. The space group, cell parameters, and powder-diffraction data are essentially identical to those of natural klebelsbergite. The crystals decompose by dehydration at 273°C.

Introduction

Klebelsbergite was first described by Zsivny in 1929 from Felsöbánya, Hungary (now Baia Sprie, Rumania). He reported that the mineral was essentially a basic antimony sulfate, containing a little H₂O, Fe, Mg, Na, K and traces of Bi and P₂O₅ by microchemical tests. He also showed that it was monoclinic with a β angle of 91.8°, biaxial negative, X = b, $Y: c = 1.8^{\circ}$, n > 1.74. In the following 50 years, several reports concerning this mineral have appeared (Manilici et al., 1965; Radulescu and Dimitrescu, 1966; Antonovici and Vanghelie, 1972); however, none of them gave a complete description of this rare mineral. Antonovici and Vanghelie (1972), in the only detailed study of this mineral since Zsivny, examined museum specimens of klebelsbergite. They reported that the mineral contained 61.86 percent antimony by X-ray fluorescence analysis and more than one percent sulfur by spectrographic analysis.

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They also presented an X-ray powder diffraction chart with the three strongest lines given as 6.19(100), 3.43(80), and 2.47A(90).

Our work gives for the first time the chemical formula, crystal data, and physical and chemical properties of klebelsbergite. This study has been done using museum specimens from the Smithsonian Institution labelled as klebelsbergite from the type locality, Baia Sprie, Rumania; NMNH #137949, R6535, B12631, B12640, B12647, and B16076.

A short version of this manuscript was submitted to the IMA Commission on New Minerals and Mineral Names. It was determined that a vote of the commission was not necessary (Max H. Hey, written communication).

Physical properties

Klebelsbergite occurs as aggregates of very small platy or acicular crystals (Fig. 1), maximum size about $0.8 \times 0.08 \times 0.02$ mm, with a pale yellow to orange-yellow color. The streak is white to yellow and the luster is vitreous. The mineral is brittle and has no distinct cleavage. The density determined by Ber-

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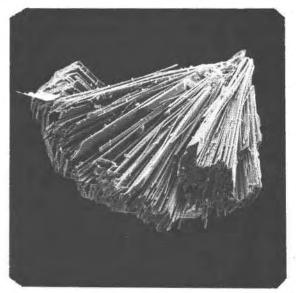


Fig. 1. SEM photograph of natural klebelsbergite from Baia Sprie, Rumania (225×).

man balance is 4.62(6) g/cm³ (average of 6 measurements using more than 10mg of sample each time). The Vickers hardness number (VHN) is 200 with a 10g load.

Klebelsbergite is optically biaxial, negative, with 2V about 70° . Because of its high reactivity with index-of-refraction liquids of high index, the refractive indices could not be measured; however, quick observation before the crystals dissolved showed that β is near 1.95. Birefringence $\beta - \gamma$ is about 0.035 (by retardation). The mean index of refraction, 1.95, was calculated from the Gladstone-Dale relationship using the ideal chemical formula, $\mathrm{Sb_4O_6H_2SO_4}$, calculated density of 4.64 g/cm³, and the constants given by Mandarino (1976), and this supports the above observation that β is about 1.95. The optical orientation is X = b, Y = c and Z = a. The mineral is not pleochroic.

Klebelsbergite is weakly magnetic when it contains iron as an impurity. This feature enabled us to separate klebelsbergite from other minerals with the Franz isodynamic separator. The mineral is not luminescent in either long- or short-wave ultraviolet radiation.

Crystallography

In the original description by Zsivny (1929), klebelsbergite was described as monoclinic; however, the true symmetry has been revealed to be orthorhombic by our X-ray diffraction studies of single crystals. Weissenberg and precession photographs show that the space group is either Pcmb or $Pc2_1b$. The lattice constants are a=11.279(2), b=14.909(3), and c=5.7648(6)A, calculated from the powder-diffraction data using the least-squares refinement program of Appleman and Evans (1973). The powder-diffraction pattern of klebelsbergite was obtained with a Guinier-Hägg camera using $CuK\alpha_1$ radiation and BaF_2 internal standard; the data are given in Table 1. Indexing of ambiguous reflections was aided by intensity measurements from single-crystal photographs.

Crystals of klebelsbergite are generally flattened to tabular on $\{010\}$ and elongated in the [001] direction. Well-developed crystals were occasionally observed; however, the small crystal size and especially the thickness of less than 20 μ m prevented us from carrying out detailed measurements of the interfacial angles. Under the optical microscope, the apical angle, indicated by an arrow in Figure 2, was measured to be about 124° . This value is found to be close to the supplemental angle of $(101) \land (\overline{101}) = 54.14^{\circ}$ calculated from the lattice constants. A section normal to the c axis shows a flattened hexagon; therefore the platy crystal can be described as having a major $\{010\}$ pinacoid and a minor probably $\{120\}$ prism.

As is seen in Figure 2, klebelsbergite crystals tend to grow as parallel blades elongated along the c axis, stacked in the b-axis direction in lamellar fashion. Precession photographs parallel to the (010) plane revealed very slight distortion among the stacked crystals; *i.e.* the (010) plane rotates slightly around the b axis. As a result, the crystals show a monoclinic appearance, and this is considered to be the reason klebelsbergite was described as monoclinic in the original description (Zsivny, 1929). Zsivny reported a beta angle of 91.8° with optical orientation $Y:c = 1.8^{\circ}$, and this is well explained as the distortion around the b axis.

Synthesis

Single crystals of klebelsbergite (maximum size about $1 \times 0.1 \times 0.03$ mm) have been synthesized by boiling 2 g of Sb₂O₃ with 600 ml of 1 N H₂SO₄ solution for four days.

The crystals obtained are colorless and transparent, having the same luster as natural crystals. The crystal habit is also similar to that of natural crystals: tabular on {010}, elongated along [001], with the dominant form {010} (Fig. 3). The crystals often developed radially with a parallel growth along [001] as in natural crystals. The VHN is 130 with 10 g

Table 1. X-ray powder-diffraction data for klebelsbergite

Natural klebelsbergite Synthetic klebelsbergite hk1 d calc d calc d obs 7.45 6.21 5.136 020 7.47 6.22 7.44 6.21 7.45 020 120 6.22 120 101 111 5.13 5,14 ms 5.137 ms m mw 4.856 4.228 4.854 4.85 4.85 4.228 4.231 121 m 3.893 3.722 3.569 211 3.892 3.892 3.731 211 3,892 3,720 3,568 040 131 3,727 221 3.546 3.545 3.350 3,546 3.535 3.545 3.534 320 3.357 301 3.149 3.150 301 3.150 3.148 vs vw ms 231 3.131 231 3,130 3.129 3.105 3.110 311 311 3.081 3.081 321 2.901 2.902 mw 321 2.901 2.901 2.882 2.882 002 012 400 2.830 2.820 2.830 012 2.833 2.833 2.819 2.735 2.691 2.821 2.736 m m 400 2.819 241 022 241 022 2.688 2.688 2,690 ms 420 2.637 2.615 2.639 2.618 w vw 420 2.636 2.637 mw 2_617 2.617 vw 2.577 2.531 2.496 151 2.576 2.529 2.497 2.531 2.497 212 2.531 2.497 212 m ms 060 2-485 2,485 060 2.482 2.435 2.435 222 2.428 2-428 vw 2.398 2.396 421 2.396 251 042 2,397 2,280 251 2,281 2.280 2.280 m mw 2,249 2,237 W VW 440 22.248 440 2.247 2.247 2.239 2.235 322 2.188 2.188 2.164 2,164 2,158 520 2,159 2.159 520 VW 2.114 2.094 2.078 2.115 2.094 2.079 242 ms mw w 332 2.078 2.078 2 072 2 072 152 2.038 2.038 2.038 vvw VW 2.016 1.997 2.015 1.998 402 412 402 2,016 2.017 ms 1.998 1.965 1.950 1.998 ms w 1.967 1.965 1.949 1.967 171 1.950 mic ms 1.946 1.946 252 1.945 1.945 ms 1.945 451 1.930 1.931 1,929 1,929 m 1.928 9 540 1,896 1,881 1,879 103 mw mw 103 1.896 1.88111 1.894 1.883 062 1.882 600 1.880 v 600 1.880 1.879 460 1.864 1,861 460 1.863 1.863 m 162 1,857 vvw 123 1.838 1.837 vw 1.829 1.822 1.815 1,830 1,822 1,815 541 vw 620 1.823 1.823 ms vw 213 1.806 1.805 8 213 1.807 1.807 1.775 1.774 1.7743 1.7735 ms m 1.7730 442 1.7688 1.767 1.768 223 VW 1 7684 280 512 1.7674 512 1.764 1.764 w 1.7627 1.7170 1 7627 1.7165 vw 1.7121 1.7124 072 1.713 1.711 mw 072 1.7012 1.6895 1.6813 313 143 1.700 mw vvw 313 143 1.7011 631 m vw 631 1,682 1.682 1.6813 640 1.6777 1.6774 532 1.673 1.673 532 1.6729 1.6727 m 452 1,6697 1.6696 νw 1.639 1.635 1.640 1.6383 1.6384 243 1.635 vw 471 1.6289 1.6291 333 542 153 1.6187 1.6193 1.6035 1.6032 vw 153 1.599 1.598 vvw 1.5994 413 1.575 1.575 1.5746 720 1.575 602 1.5745

Table 1. (continued)

Nat	ural klebe	lsbergite	Synthetic klebelsbergite				
hk1	d calc	d obs	I*	hk1	d calc	d obs	1
462	1,565	1,5655	m	462	1.5649	1.5644	ms
082 423	1.565	1.5529	m	082 423	1.5640	1,5534	ms
253	1.553			253	1,5533		
182 711	1.550	1.549	VVW	711 622	1.5432	1,5431	VW
522	1.5435	1.5401	VVW	651	1.5323	1.5321	mw.
291	1.5322	1.5318	IB	291	1.5306	1.5306	m
433	1.5126	1.5132	w	433	1.5133	1,5132	m
163	1.5065	1.5067	vw	163 481	1.5067	1.5067 1.4997	w vw
				353	1.4844	1.4841	W
1.10.0	1.4781	1.4777	vvw	1.10.0	1.4762	1.4764	w
472	1.4640	1.4630	vw	472	1.4634	1.4633	w
				642	1.4503	1.4502	w
523	1.4354	1.4355	W	523 014	1.4360	1.4358	m
				741	1.4321	1.4318	vw
192	1,4248	1.4238	nw	114	1,4244	1.4242	ms
114	1.4238			192	1.4237		
024	1.4150	1.4149	w vw	024 800	1.4163	1.4163	m mw
800 363	1.4093	1,4093	V W	363	1.4094	11.071	
124	1.4040	1.4040	VW	124	1,4053	1.4053	mw
712	1.4002	1.4005	νW	712	1.4003	1.4003	m
3.10.0	1.3859	1.3859	mw	2.10.1	1.3968	1.3969	w mw
034	1.3842	1.3839	w	652 034	1.3921	1.3850	III
034	1.5042	11,5057		820	1.3849		
				273	1,3831	1.3833	mw
				751	1.3760		VW
671	1.3691	1,3686	VVW	671 482	1.3682	1.3681	W
482 811	1.3684	1.3636	inw	811	1.3635	1.3631	ms
				234	1.3453	1.3456	mw
314	1 2122	1.3401	vw.	044 392	1.3452	1.3409	w
314	1.3403	1.3401	vw.	613	1,3388	1.3387	vv
0.10.2	1.3243	1.3240	vw	0.10.2	1.3232	1.3225	w
680	1.3235	1.5240	300	680	1.3225		
				840	1.3182	1.3180	vw
761 334	1.3162	1.2988	VVW	761 334	1.2999	1,2998	mw
334	1.2,00	112,00	* "	591	1.2997		
841	1.2855	1.2856	VVW	841	1.2851	1.2851	vw
473	1.2731	1.2733	VVW	752	1.2718	1.2713	W
424 254	1.2647	1.2646	m	802 672	1.2665	1.2656	TP:
234	1.2045			900	1.2529	1.2531	W
822	1 . 2486	1.2483	vw	822	1.2486	1.2481	w
851	1.2445	1,2446	W	851	1.2441	1,2438	mw
0.12.0	1.2424	1.2421	vw	5.10.0 0.12.0	1.2426	1.2425	W
0.12.0	1.6727	1,2-121		713	1.2308	1.2306	vw
				354	1.2272	1.2274	vw
				832	1.2272		
				0.11.2 653	1.2255	1.2251	mw
				901	1.2244		
				911	1.2203	1.2205	vw
				264 444	1.2178	1.2179	w mw
				691	1.2140	1.1137	III W
				592	1.2108	1.2110	W
				921	1.2082	1.2080	vw
4.10-2	1,1987	1.1985	Total Control	573 842	1.2057	1.2057	VV m 5.7
4.10,2	1,170/	1.1703	w	772	1.1732	1.1991	mw vw
				3.12.1	1.1545	1.1547	W
				2.10.3	1.1526	1.1526	vv
				583	1.1505	1.1508	vw
				493 604	1.1460	1.1460	w
				464	1.1406	1.1406	W
				374	1.1379	1,1381	vw
				5.11.1	1.1378		
				673 184	1.1364	1.1363	vw w
				951	1.1324	1.1325	vw
				10.0.0	1.1276	1.1277	W
				554	1.1252	1.1253	vw
				4.11.2	1.1239	1.1239	w
				634	1.11/9	1.1151	w mw
				833	1.1083	1.1086	mw
				235	1.1024	1.1027	vv
				145	1.0971	1.0974	W
				564	1.0914	1.0918	W

^{*} Relative intensities estimated visually (s, strong; m, medium; w, weak; v, very; b, broad.)

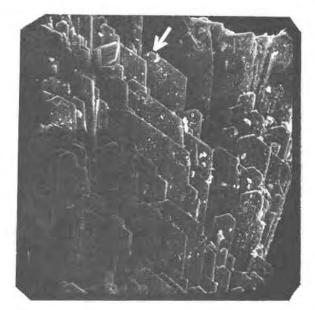


Fig. 2. SEM photograph of natural klebelsbergite from Baia Sprie, Rumania (400×).

load, as compared to 200 measured on natural crystals. The powder-diffraction data obtained with a Guinier-Hägg camera using $CuK\alpha_1$ radiation and CaF₂ standard are given in Table 1. The data agree very well with those for natural klebelsbergite. The crystal symmetry and the space group are the same as the natural mineral, i.e. orthorhombic Pcmb or $Pc2_1b$. Cell dimensions refined from the powder data are a = 11.2763(7), b = 14.890(1) and c =5.7707(4)A, in excellent agreement with natural klebelsbergite; the difference is almost within one standard deviation. The measured density of the synthetic material is 4.61(4) g/cm³ which agrees well with the calculated density 4.69 g/cm³ and that measured on natural crystals, 4.62(6) g/cm³. Thus the synthetic crystals have been identified as klebelsbergite.

Chemistry

Wave-length scans using an ARL-SEMQ electron microprobe disclosed that klebelsbergite contains major Sb and S and minor Fe. No other elements ($Z \ge 9$) were detected. The presence of sulfate ion was confirmed by a spot test with BaCl₂ solution.

As very little material was available, wet-chemical analyses were carried out only on synthetic klebels-bergite. Sb was determined by KBrO₃ titration, and SO₄² gravimetrically with BaCl₂ solution (Hillebrand and Lundell, 1953). Water was determined by the Penfield method and the sample was heated to

300°C. The result is given in Table 2, which yields the empirical formula $Sb_{407}O_{500}SO_4 \cdot 1.03H_2O$ (on the basis of $SO_4 = 1$), and ideally $Sb_4O_6H_2SO_4$, which can be expressed as either $Sb_4O_4(OH)_2SO_4$ or $Sb_4O_5SO_4H_2O$. On the basis of IR spectroscopy, $Sb_4O_4(OH)_2SO_4$ is favored (see below); however, $Sb_4O_5SO_4 \cdot H_2O$ cannot be ruled out completely.²

Quantitative analyses of natural klebelsbergite were carried out with the electron microprobe using an operating voltage of 15kV and a sample current of 0.15 microamps. Valentinite, barite, and olivine were used as standards for Sb, S, and Fe, respectively. The results are given in Table 2. Concentration of iron varies from 0.0 to 0.5 percent and is thought to be present substituting for antimony as explained below. The rather large discrepancy of sulfur content between the two analyses, wet-chemical and microprobe, is considered to be due to inadequate data correction of the microprobe data. The ideal chemical formula $Sb_4O_6H_2SO_4$ with Z = 4 is suitable for the space groups Pcmb or Pc21b. The theoretical density for the ideal formula is 4.69 g/cm³, which agrees well with the measured density 4.62(6) g/cm³. Klebelsbergite is readily soluble in HCl(1:1), and NaOH(10 percent), slowly soluble in tartaric acid (20

² After this paper was accepted for publication, we received a preprint from Drs. S. Menchetti and C. Sabelli of the Università di Firenze describing their recent determination of the klebelsbergite structure (Menchetti and Sabelli, in preparation). They found that the correct formula is indeed $Sb_4O_4(OH)_2SO_4$, and that the space group is $Pc2_1b$.

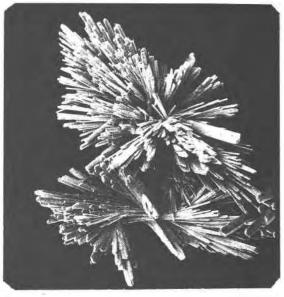


Fig. 3. SEM photograph of synthetic klebelsbergite (60×).

percent), and very slowly soluble in HNO₃(1:1) and $H_2SO_4(1:1)$.

The color of klebelsbergite is more or less yellow; however, the depth of the yellow color varies from specimen to specimen. To elucidate the origin of the color, microprobe analyses were made on various specimens with different depths of color. The result indicates that there is a clear positive correlation between the iron concentration and the depth of the color. Iron-free synthetic klebelsbergite is colorless. Therefore, the origin of the yellow color is ascribed to the iron impurity. To establish the valence of iron, spot tests were made using 0.1 N potassium ferri- and ferro-cyanide solutions. HCl(1:3), HNO₃(1:1) and tartaric acid (20 percent) were used to dissolve the crystals. The tests clearly indicate the presence of both trivalent and divalent iron; however, trivalent iron is much more abundant.

Thermal treatments

DTA/TGA

DTA/TGA spectra were taken under He atmosphere on a 100 mg sample of synthetic klebelsbergite. Heating rate was 4°C/min, and the temperature ranged from room temperature to 940°C. The spectra are characterized by an endothermic peak at 273°C with a weight loss of 2.4 percent, which is attributed to a dehydration reaction. The value corresponds well with the result of quantitative analyses of water (see Table 2). After the thermal analysis, the specimen was X-rayed and was found to be Sb₂O₄.

Decomposition

Decomposition experiments were carried out in an evacuated Pyrex glass tube (15 cm in length; 5 mm ID). A 100 mg sample of synthetic klebelsbergite was placed at the end of a sealed glass tube, then the other end was closed. The tube was placed in a horizontal furnace; the end of the tube containing the sample was located inside the furnace along with a thermocouple, while the other end (about 4 cm) was kept outside of the furnace. On heating to 300°C, a condensation of vapor was observed at the cool end. The specimen was heated for 1 hour, after which the dehydration product and the condensed liquid were tested.

The product kept the original shape; however, it became translucent with a slightly brownish color. The precession photograph of a "single crystal" which had been heated in the same way as above gave diffuse spots which approach a powder pattern,

Table 2. Chemical analysis of klebelsbergite

	1	2	3	4	5	6
Sb	71.78	0.5896	4.066	4	71.51	70.8
S	4.65	0.1450	1	1	4.71	3.2
Fe						0.5
H ₂ 0 0*	2,69	0.1493	1.030	1	2.64	
0*	20.88	1.305	9.000	9	21.14	
Total	100.00				100.00	

- Synthetic klebelsbergite by wet-chemical analysis (wt.%)
- 2. Molecular quotient
- Mole number basis S=1 (or SO₄=1) Ideal mole number
- Theoretical $\mathrm{Sb_40_6H_2S0_4}$ Natural klebelsbergite by microprobe analysis

and this indicates that a single crystal changes into the polycrystalline state by dehydration. The X-ray powder diffraction pattern of the dehydration product has the following 8 strongest lines: 4.55(4), 3.35(1), 3.12(10), 2.80(1), 2.665(6), 1.926(3), 1.886(3), 1.615(3)A; the pattern as a whole shows a similarity to that of senarmontite, Sb₂O₃. The product dissolved in HCl solution (1:1) and gave a brown colloidal residue. Successive addition of BaCl₂ solution (10 percent) into the HCl solution caused an abundant white precipitate to form. This observation indicates that the sulfate ion remains in the structure after heating to 300°C. However, there is evidence of a partial decomposition of sulfate ion; the condensed liquid was strongly acidic, indicating that a small part of the sulfate ion decomposed on heating and the resultant SO₂ gas dissolved in the liquid.

Infrared spectroscopy

IR spectra were measured by the conventional KBr method for the region from 4000 to 645 cm⁻¹. Figure 4 shows the spectra of synthetic klebelsbergite (a), of synthetic klebelsbergite heated at 300°C for 1 hour (b), and of natural klebelsbergite (c). Since spectra of both natural and synthetic klebelsbergite have essentially the same features, we treat them together in the following discussion; the values in the parentheses are those for synthetic klebelsbergite. A sharp absorption of 3435 (3435) cm⁻¹ is attributed to the O-H stretching vibration, and it disappeared on heating to 300°C for 1 hour: spectrum (b). Judging from the absence of any sizable peak in the 1650 cm⁻¹ region, this absorption is more likely to be due to the stretching vibration of hydroxyl ion than to that of water of hydration; the O-H bending frequency usually occurs near 1650 cm⁻¹ in the cases of water of hydration and absorbed water, but not in the case of hydroxyl ion (Gadsden, 1975). Although there is a very weak broad absorption in the 1650

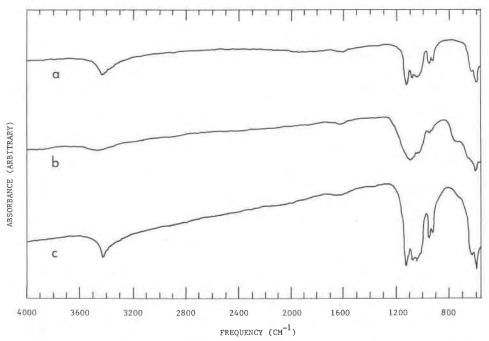


Fig. 4. Infrared spectra: (a) synthetic klebelsbergite; (b) synthetic klebelsbergite heated at 300°C for 1 hr.; (c) natural klebelsbergite from Baia Sprie, Rumania.

cm⁻¹ region, it can be attributed to the absorbed water in the KBr, since spectrum (b) has a similar peak. The broad absorption in the 3450 cm⁻¹ region in spectrum (b) can be assigned to the same origin.

Absorptions at 1130 (1130), 1080 (1085), 1050 (1045) cm⁻¹ are here attributed to the ν_3 vibration of SO_4^{2-} ion, 955 (955) and 935 (930) cm⁻¹ are to ν_1 , 635 (635) and 600 (600) cm⁻¹ are to ν_4 (Gadsden, 1975; Miller and Wilkins, 1952; Miller et al., 1960). In the free tetrahedral SO_4^{2-} ion, ν_1 is the non-degenerate stretching vibration and is infrared inactive, and ν_3 and v_4 are triply degenerate vibrations and are infrared active. Distortion of the tetrahedron causes ν_1 to be active, and removes the degeneracy of ν_3 and ν_4 . therefore the clear multiplet character of the above peaks in the ν_3 and ν_4 regions is explained by the presence of significant distortion of the SO₄²⁻ tetrahedron in the klebelsbergite structure. The doublet character of ν_1 may possibly be explained by the presence of several kinds of S-O bonds. On heating, as seen in spectrum (b), the above splittings in the ν_3 and ν_4 regions become degenerate, and ν_1 absorption almost extinguished. This implies that the symmetry of SO₄² increases and becomes close to T_d symmetry. There is also evidence of a partial decomposition of SO₄²⁻ ion on heating (see "Thermal treatment"). This may be related to the apparent increase in SO₄² symmetry shown in the IR spectra.

In addition, spectrum (b) shows the appearance of a rather strong peak at 745 cm⁻¹. This value reminds us that senarmontite (the low-temperature form of Sb₂O₃) has very strong absorption at 740 cm⁻¹ (Miller and Wilkins, 1952). X-ray powder-diffraction patterns of the heated sample show a possible structural similarity to senarmontite (see "Thermal treatment"). Combined with the above change in the absorption peaks of SO₄²⁻ ion, this observation may indicate that the structure of klebelsbergite changes to an antimony oxide-type structure on heating.

Association and paragenesis

Klebelsbergite is found as fan-shaped aggregates composed of many small crystals (Fig. 1) from Baia Sprie, Rumania. Associated minerals are stibnite, barite, and valentinite. We observed that klebelsbergite crystals almost always grew on stibnite crystals. Large euhedral crystals of barite (more than 1 cm) with stibnite inclusions are commonly observed. Surfaces of the stibnite crystals often show evidence of corrosion. It seems clear that the source of the antimony in klebelsbergite is stibnite. The mineral relationships suggest that stibnite crystallized first, followed by barite, and finally klebelsbergite.

A characteristic feature of the mineral assemblage is the intimate association of valentinite with klebelsbergite. Valentinite is found as small massive spherical crystals on the surface of stibnite crystals. The color is pale yellow to dark yellow, and the appearance resembles that of klebelsbergite except for the habit. Microprobe analyses of the valentinite crystals showed the presence of iron in unusually high concentrations. The concentration of iron was generally higher at the rim than at the core of the spherical crystals; in some cases iron concentrations of up to 5 weight percent were observed. The yellow color is likely due to the iron. Valentinite is very probably an alteration product of stibnite.

The above investigation clearly shows that klebelsbergite is a valid mineral species. It is the only naturally-occurring sulfate of antimony. Hereby the NMNH specimens are all designated as neotypes.

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