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2 **Effect of faceting on olivine wetting properties**

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16

## 17 **Abstract**

18 Grain-scale pore geometry primarily controls the fluid distribution in rocks, affecting  
19 material transport and geophysical response. The dihedral angle ( $\theta$ ) in the olivine–fluid system is  
20 a key parameter determining pore fluid geometry in mantle wedges. In the system, curved and  
21 faceted olivine–fluid interfaces define  $\theta$ , resulting in faceted–faceted (FF), faceted–curved (FC),  
22 and curved–curved (CC) angles. The effect of faceting on  $\theta$  under various pressure and temperature  
23 (P–T) conditions and fluid compositions, however, have not been constrained, and mineralogical  
24 understanding remains unresolved. This study evaluated facet-bearing  $\theta$  and their proportions in  
25 olivine–multicomponent aqueous fluid systems. Our results show that 1/3 of olivine–fluid  $\theta$  are  
26 facet-bearing angles, regardless of the P–T conditions and fluid composition. Faceting produces  
27 larger dihedral angles than CC angles. The grain boundary plane (GBP) distribution reveals that  
28 the GBPs of faceted interfaces at triple junctions have low Miller index faces ( $\{100\}$ ,  $\{010\}$ , and  
29  $\{101\}$ ). The misorientation angle/axis distributions of adjacent grain pairs are in accord with a  
30 theoretical distribution of random olivine aggregate. Moreover, the calculation of the FF angles  
31 for adjacent grain pairs with low Miller index GBPs reproduces measured angle values based on  
32 the olivine crystal habit. Therefore, our study suggests that the FF angle is strongly affected by  
33 olivine crystallography. The presence of faceting increases  $\theta$  and a critical fluid fraction ( $\phi_c$ ) for  
34 percolation, lowering permeability. In the mantle wedge, where olivine crystallographic preferred  
35 orientation (CPO) is expected owing to corner flow, increasing the FF angle proportion with  
36 associated changes in fluid pore morphology will lead to permeability anisotropy, controlling the  
37 direction of the fluid flow, and resulting in geophysical anomalies such as seismic wave attenuation  
38 and high electrical conductivity.

39 **Keywords:** dihedral angle, faceted plane, Miller index, crystallographic orientation, permeability  
40 anisotropy, mantle wedge.

## 41 1. Introduction

42 Pore geometry significantly controls the distribution of geological fluids (i.e., aqueous  
43 fluids and silicate melt) in deep mantle wedges, thereby affecting element cycling and geophysical  
44 responses in subduction zones (Watson and Brenan 1987; Hermann et al. 2006; Iwamori 1998;  
45 van Keken et al. 2011; Pommier and Evans 2017; Worzewski et al. 2011; Zheng et al. 2016).  
46 Although channelized fluid flow has often been inferred from field studies (Angiboust et al. 2014),  
47 pervasive grain-scale fluid flow may be the most plausible fluid migration regime at high pressure  
48 (P) and high temperature (T) conditions where dissolution–precipitation intensively operates and  
49 interfacial energy minimization (“textural equilibrium”) is quickly attained. Moreover, the  
50 pervasive nature may be suitable for explaining the resistivity anomalies observed at a  
51 magnetotelluric (MT) grid scale (commonly > 10 km) because it would be required for the  
52 channelized flows to be distributed continuously and nearly isotropically over this length scale. In  
53 an olivine-dominant mantle rock, the olivine–fluid dihedral angle ( $\theta$ ) is the primary parameter  
54 controlling grain-scale fluid connectivity (Toramaru and Fujii 1986; Mibe et al. 1999; Huang et al.  
55 2019, 2020). Therefore, a precise constraint on  $\theta$  in the olivine–fluid system is important for a  
56 complete understanding of the fluid distribution and migration in subduction zones.

57 The dihedral angle is a consequence of the fluid–mineral interaction, which changes the  
58 fluid pore geometry through dissolution and precipitation processes to minimize the interfacial  
59 energy in the system. It is defined as the ratio of the grain boundary energy ( $\gamma_{ss}$ ) to the solid–fluid  
60 interfacial energy ( $\gamma_{sf}$ ) (Smith 1964) as follows:

61 
$$2 \cos(\theta/2) = \gamma_{ss}/\gamma_{sf} \quad (1)$$

62 In an isotropic system where solid–fluid interfaces are smoothly curved with a constant mean  
63 curvature; the equilibrium geometry of the intergranular fluid is solely determined by the dihedral  
64 angle. In a fluid-bearing rock with a low fluid fraction ( $\phi$ ), fluids can wet the grain edges well and  
65 migrate along interconnected tubular networks at  $\theta < 60^\circ$  irrespective of the fluid fraction. On the  
66 contrary, at  $\theta > 60^\circ$ , the fluid is distributed as isolated pockets along the grain edges, corners, and  
67 boundaries (Smith 1948; Watson and Brenan 1987; Holness 1992, 1993), although the unstable  
68 interconnected network can be transiently formed above the critical fluid fraction ( $\phi_c$ ) (Park and  
69 Yoon 1985; von Bargen and Waff 1986; Laporte and Provost 2000). However, the fluid  
70 distribution in realistic rocks can deviate from the ideal distribution (Waff and Faul 1992; Laporte  
71 and Watson 1995). Huang et al. (2021) measured the electrical conductivity of a texturally  
72 equilibrated forsterite–saline fluid aggregate at 800°C and 1 GPa and showed that the conductivity  
73 was significantly smaller than that expected from the interconnected tube model, especially at low  
74  $\phi$ . The synchrotron X-ray microtomography (CT) of the post-run products revealed that fluid pores  
75 struggled to become interconnected at  $\phi$  approximately <1.0% even though  $\theta$  defined by the  
76 curved–curved interface was  $< 60^\circ$  under the experimental conditions (Huang et al. 2019).

77 In a realistic mineral–fluid system with interfacial energy anisotropy, facet planes (i.e.,  
78 crystallographically controlled planar solid–liquid interfaces) are often present along with curved  
79 interfaces (Waff and Faul 1992; Watson and Lupulescu 1993; Cmíral et al. 1998; Watson 1999;  
80 Wark and Watson 2000; Price et al. 2006). A curved interface results from a constant mean  
81 curvature that minimizes the surface energy by minimizing the surface area (Bargen and Waff  
82 1986; Waff and Faul 1992), whereas a flat interface is attributed to crystallographically controlled

83 minimum interfacial energy (Yoshino et al. 2006). The coexistence of both curved and faceted  
84 interfaces generates three types of  $\theta$ : curved–curved (CC), faceted–curved (FC), and faceted–  
85 faceted (FF) (Price et al. 2006; Yoshino et al. 2006). Pores surrounded by facet planes are difficult  
86 to connect unless the fluid/melt fraction exceeds a critical value, which depends on  $\theta$  defined by  
87 the faceted interfaces (Waff and Faul 1992; Cmíral et al. 1998; Price et al. 2006). Therefore, the  
88 facet plane may affect the fluid connectivity depending on its  $\theta$  values, which helps to explain the  
89 results of electrical conductivity measurements by Huang et al. (2021). Waff and Faul (1992) and  
90 Cmíral et al. (1998) investigated  $\theta$  in partially molten rock systems and showed that FF angles are  
91 larger than FC and CC angles. Price et al. (2006) reported that FF and FC angles are larger than  
92 CC angles in the quartz–fluid and tremolite–fluid systems at 900–950 °C and approximately 1.5  
93 GPa. These findings imply that facet-bearing pores necessitate a relatively high  $\phi$  for fluid  
94 interconnection.

95         Although many studies have been conducted to investigate the P–T, fluid composition, and  
96 mineral assemblage dependence of  $\theta$  in the olivine–fluid system (Watson and Brenan 1987; Huang  
97 et al. 2019, 2020; Mibe et al. 1998, 1999; Yoshino et al. 2006), the effect of angle type variation  
98 on  $\theta$  is poorly understood. Yoshino et al. (2006) systematically investigated the effect of faceting  
99 on pore geometry in texturally equilibrated rocks and discussed its implications for permeability  
100 in several representative systems, including the San Carlos olivine–MORB melt. However, olivine  
101 (forsterite)–aqueous fluid systems have not yet been studied. Huang et al. (2019, 2020) focused on  
102 the CC  $\theta$  between olivine and multicomponent aqueous fluids over a wide range of pressures and  
103 temperatures and proposed that multicomponent fluids derived from the subducting slab can travel  
104 through the deep fore-arc mantle wedge and cause electrical conductivity anomalies observed in

105 various subduction zones. Laporte and Provost (2000) investigated theoretically  $\theta$  in a system with  
106 simplified surface energy anisotropy and found that the relationship between the mean equilibrium  
107  $\theta$  and the grain boundary to the surface energy ratio was close to the isotropic case. However, few  
108 studies have provided empirical evidence for facet-bearing dihedral angles with respect to  
109 crystallographic orientations for actual mineral–fluid systems with various fluid compositions.

110 Investigating the grain boundary plane distribution (GBPD) can provide a mineralogical  
111 understanding of angle types in terms of FF, FC, and CC angles. Some researchers have  
112 investigated GBPD in fluid-free olivine polycrystals and found that low Miller index planes  
113 preferentially appear in the grain boundaries (Faul and Fitz Gerald 1999; Marquardt et al. 2015).  
114 If such specific grain boundary planes (GBPs) appear preferentially at the facet-bearing triple  
115 junctions in the mineral–fluid system, they may significantly control  $\theta$  of the facet-bearing angles  
116 and their resultant fluid pore geometry. However, previous studies did not distinguish the grain  
117 boundary type in terms of angle type (i.e., FF, FC, or CC), and GBPD has not been examined in  
118 fluid-bearing systems.

119 In this study, we investigated the facet-bearing (i.e., FF and FC)  $\theta$  in the run products of  
120 olivine–fluid systems previously obtained by piston-cylinder experiments at various pressures (P:  
121 1–3 GPa), temperatures (T: 800–1100 °C), and fluid compositions (pure H<sub>2</sub>O, H<sub>2</sub>O–NaCl, and  
122 H<sub>2</sub>O–CO<sub>2</sub> systems) (Huang et al. 2019, 2020). These experiments broadly covered the P–T  
123 conditions and fluid compositions expected in a deep fore-arc mantle wedge. The results were  
124 compared with CC  $\theta$  measured in previous studies to clarify the effect of faceting. Moreover, we  
125 examined the crystallographic orientation of the olivine aggregate and identified the GBPD at  
126 triple junctions with different angle types to evaluate the influence of crystallographic orientation

127 on  $\theta$ . Based on the results, we discussed the origin of faceting and its effect on  $\theta$  and inferred the  
128 consequences of faceting on fluid connectivity, fluid distribution, and permeability anisotropy in  
129 the olivine–fluid system under static and sheared mantle conditions.

## 130 **2. Methods**

### 131 **2.1. Samples**

132 We analyzed the run products of olivine–fluid systems previously obtained by Huang et al.  
133 (2020, 2019). To constrain the P–T and fluid composition dependency of the facet-bearing  $\theta$ , we  
134 selected 19 samples that covered a wide range of experimental P–T conditions (1–3 GPa and 800–  
135 1100 °C) and fluid compositions ( $\text{H}_2\text{O}$ ,  $\text{H}_2\text{O}$ – $\text{CO}_2$  with  $X_{(\text{CO}_2)} = \text{CO}_2/(\text{H}_2\text{O}+\text{CO}_2) = 0.5$  in molar  
136 ratio), and  $\text{H}_2\text{O}$ – $\text{NaCl}$  with 5.0 and 27.5 wt.%  $\text{NaCl}$ ). Run products containing magnesite and  
137 orthopyroxene due to olivine carbonation (Huang et al., 2020) were excluded to avoid the effect  
138 of mineral species other than olivine.

139 Experimental procedures to synthesize the samples followed those of previous studies  
140 (Huang et al., 2020, 2019) and are briefly described here. High P–T experiments were conducted  
141 at Bayerisches Geoinstitut, University of Bayreuth, using end-loaded piston-cylinder apparatus.  
142 Piston diameters of 3/4 inch and 1/2 inch were used for experiments at pressures of 1–2 GPa and  
143 3 GPa, respectively, along with a standard Talc–Pyrex assembly (Keppler et al. 2003). The starting  
144 material was San Carlos olivine powder ( $\text{Fo}_{91}\text{Fa}_9$ ) with a grain size of 38–53  $\mu\text{m}$ . Deionized and  
145 distilled water were used as pure  $\text{H}_2\text{O}$  sources. Pure oxalic acid dihydrate ( $\text{C}_2\text{H}_2\text{O}_4 \cdot 2\text{H}_2\text{O}$ ) was  
146 used as the  $\text{CO}_2$  source. To obtain the  $\text{NaCl}$  solution, we dissolved reagent-grade  $\text{NaCl}$  (99.99%  
147  $\text{NaCl}$ ) in deionized and distilled water at room temperature (approximately 25°C) and atmospheric  
148 pressure. Water and hydrous oxalic acid were mixed to obtain the target  $X_{(\text{CO}_2)}$ . The olivine powder,

149 combined with approximately 10.0 vol.% fluid, was loaded into an end-welded noble metal  
150 capsule that was sealed by arc welding. Au capsules (2.2 mm outer diameter) and Au<sub>80</sub>Pd<sub>20</sub> alloy  
151 capsules (2.0 mm outer diameter) were used for experiments at 800–1000 °C and 1100°C,  
152 respectively. Detailed procedures for the high P–T experiments are shown by Huang et al. (2019,  
153 2020). Briefly, the run duration ranged from 72 to 211 h, depending on the target temperature. The  
154 experimental conditions and results are summarized in Table 1. In our experiments, CO<sub>2</sub> was  
155 assumed to be the predominant C species based on previous studies (Allen 1972; Huang et al.  
156 2020; Médard et al. 2008).

157         The post-run capsules were cut using a diamond wire saw to expose the run products. The  
158 run products were then impregnated with epoxy resin under a vacuum. We first polished these  
159 products using sandpaper and diamond paste with particle sizes down to 1.0 µm and then using a  
160 0.06 µm colloidal silica suspension. For electron backscattered diffraction (EBSD) analysis, the  
161 samples were polished with colloidal silica suspension for more than 10 h using an automatic  
162 vibratory polishing machine (VibroMet, Buehler Ltd.).

## 163 **2.2. Scanning electron microscopy**

164         We observed polished cross-sections of the run charges using a field-emission scanning  
165 electron microscope (FE–SEM; JSM–7100F, JEOL Ltd.) with an accelerating voltage of 15 keV.  
166 The mineral phases and microstructures of samples were observed. We took FF and FC-type triple  
167 junctions for quantitative analysis. More than 300 backscattered electron (BSE) or secondary  
168 electron (SE) images of 1280 × 960 pixels were obtained for each run product with high  
169 magnifications and resolution, depending on the pore size.

170



### 171 **2.3. Dihedral angle measurement**

172 The apparent FF and FC  $\theta$  were measured at triple junctions on the SEM images using the  
173 Image J software (National Institute of Health). We followed the method used for the CC angle  
174 measurements by Huang et al. (2019, 2020). The median of measured angles from a two-  
175 dimensional (2D) section is likely to be close to the median value of the population of the true  
176 three-dimensional (3D) angle (e.g., Harker and Parker 1945; Jurewicz and Jurewicz 1986).  
177 Generally, the median of measured angles is close to the 3D median  $\theta$  value to the greatest extent  
178 when a sufficient number of angles are measured by using high magnification and high-resolution  
179 SEM images. However, a deviation between the estimated and true angle can occur when the  
180 sample suffers from severe plucking during cutting and polishing, which locally reduces the  
181 randomness of the apparent angle distribution. Although Cmíral et al. (1998) demonstrated that  $\theta$   
182 values obtained with transmission electron microscopy (TEM) are smaller than those measured  
183 from low magnification SEM images, our FE–SEM images were taken at high magnification (up  
184 to 150,000) and provided clear images comparable with those from TEM. In this study, more than  
185 100 angles were measured on the acquired SE images for each angle type for each sample. The  
186 statistical error of the median value was estimated to be  $< 1.5^\circ$ , as discussed by Huang et al. (2020).  
187 The details of  $\theta$  selection and measurement are given by Huang et al. (2019).

### 188 **2.4. Electron backscattered diffraction**

189 To identify the crystallographic orientation of the olivine aggregate and its consequences  
190 on  $\theta$ , we mapped 10 representative samples using an FE–SEM equipped with an electron  
191 backscatter diffractometer (EBSD; HKL Channel5, Oxford Instruments plc.) at Tohoku University.  
192 Analyses were conducted at an accelerating voltage of 15.0 kV. We used six Kikuchi reflectors

193 for EBSD analysis, and EBSD mapping was performed using a 250 nm step size. Two types of  
194 EBSD data were collected for each sample. Lower magnification ( $\times 200$ – $500$ ) data were collected  
195 to investigate the crystallographic preferred orientation (CPO) of the entire sample, while higher  
196 magnification ( $\times 5,000$ – $15,000$ ) data were collected to identify crystallographic orientations near  
197 the fluid pool. EBSD data were processed using the MTEX MATLAB toolbox.

198         The degree of mismatch between measured Kikuchi patterns and calculated patterns  
199 expected for a given crystal structure is expressed as the mean angular deviation (MAD) value; a  
200 higher MAD value is likely to have large uncertainties in measurements. In this study, orientation  
201 data with MAD values of  $> 1.0^\circ$  were removed, and grain boundaries were detected using a  
202 threshold misorientation angle of  $10^\circ$ . Indexed grains smaller than  $1 \mu\text{m}$  were removed during the  
203 denoising procedure because they may be caused by mis-indexing. The CPO of analyzed samples  
204 were constructed from one point per grain. The index procedure was performed according to the  
205 manual of the MTEX MATLAB toolbox. For samples selected for EBSD analysis, we measured  
206 the grain size on high resolution SEM images using Image J and normalized grain size using the  
207 mean size value.

## 208 **3. Results**

### 209 **3.1. Product phases and microstructures**

210         In all systems, the recovered samples were composed of olivine aggregates and  
211 intergranular fluid pores that were mostly filled with epoxy resin (Figure 1). The grain size of  
212 olivine in the run products increased with increasing temperature, reaching approximately  $110 \mu\text{m}$   
213 at  $1100^\circ\text{C}$  through grain growth by Ostwald ripening and coalescence of two adjacent grains. The  
214 mean grain size of EBSD-analyzed samples ranged from  $8.2$  to  $32.3 \mu\text{m}$  (Figure 1 and Figure S1

215 of the Supporting Information). Olivine grains were compositionally homogeneous and fluid-filled  
216 pores were generally encompassed by three or more grains. Curved interfaces often coexisted with  
217 flat interfaces, even within a single pore (Figure 1). These two kinds of interfaces produced the  
218 three types of apparent  $\theta$ : CC, FC, and FF (Figure 1b, d, and f). The attainment of local interfacial  
219 energy minimization via balancing of interfacial tensions at triple junctions (i.e., textural  
220 equilibration) was demonstrated by (1) the occurrence of many olivine–olivine–olivine triple  
221 junctions with angles of approximately  $120^\circ$  (Figure 1a, e; e.g., Liu et al., 2018), (2) cumulative  
222 frequency curves for apparent  $\theta$  that showed good agreement with the predicted curve for an  
223 equilibrated texture (Harker and Parker 1945; Elliott et al. 1997), and (3) a normalized grain size  
224 distribution concentrated on the mean grain size (Figure 1g, h; Figure S1 of the Supporting  
225 Information; Faul 1997; Huang et al. 2021). This is supported by the fact that the experimental  
226 durations (72–211 h) were sufficiently long for attaining textural equilibrium compared with those  
227 in previous studies (e.g., 12 h for a grain size of  $10\ \mu\text{m}$  at  $727^\circ\text{C}$ , Holness and Siklos 2000). Once  
228 texture equilibrium is attained in the system, true 3D  $\theta$  likely remains constant with normal grain  
229 growth. We measured the apparent FF and FC angles with clear interfaces (e.g., angles denoted by  
230 green rectangles in Figure 1b) and excluded those affected by cracks (e.g., the angle indicated by  
231 the red rectangle in Figure 1d).

### 232 **3.2. Proportion of faceting-bearing angles**

233 The proportion of facet-bearing angles (i.e., FF and FC angles) was evaluated from the  
234 SEM images of each recovered sample using Image J (Figure 2). A total of 5,025 angles were  
235 counted from the 19 samples (Table 1). To avoid the potential effect of heterogeneity in the angle  
236 type distribution, we processed several images and obtained an average value for each sample.

237 Although minor fluctuations occurred, the facet-bearing angle proportion was almost constant at  
238 approximately 32.5% ( $\pm 0.5$ ) without systematic P–T and fluid composition dependencies. Namely,  
239 the FF and FC angles constituted approximately 1/3 of the dihedral angles. It is well demonstrated  
240 that the proportion of faceting is primarily controlled by  $\phi$  and that the faceting proportion  
241 increases with an increase in the liquid fraction in the solid–liquid system because of increases in  
242 lower surficial energy planes induced by grain rotation at high  $\phi$ . (Watson and Lupulescu 1993;  
243 Wark and Watson 2000; Watson 1999; Yoshino et al. 2005, 2006). Yoshino et al. (2005, 2006)  
244 reported that the faceting interface fraction in the olivine–basaltic melt system mostly fell in the  
245 range of 30.0–35.0% with a melt fraction of 10.1–16.5 vol.%. Given an initial  $\phi$  of approximately  
246 10.0 vol.% in our study, our calculated faceting proportion was roughly consistent with the  
247 previous research (Yoshino et al. 2006). Although the amount of faceting can be affected by  
248 surface adsorption, which is a function of fluid compositions (Kretz 1966), the fluid composition  
249 dependence of the proportion of facet-bearing angles was not obvious in our run products.

### 250 **3.3. Cumulative frequency of apparent dihedral angles**

251 Representative cumulative frequencies of the measured FF and FC angles in the olivine–  
252 fluid systems at 2 GPa are shown in Figure 3; curves of the apparent CC angles from Huang et al.  
253 (2019, 2020) are plotted for comparison. Cumulative frequencies for the other conditions and  
254 histograms for all systems in this study are shown in Figure S2 and Figure S3 of the Supporting  
255 Information. For all systems investigated, the cumulative frequencies mostly showed a sharp  
256 increase around the median  $\theta$ , and the frequency distribution histogram also showed a concentrated  
257 distribution of measured angles around the median value. In most cases, the cumulative frequency  
258 of CC angles sharply increased around the median  $\theta$ , which is in accord with the theoretical

259 prediction for the isotropic system with one true  $\theta$ . In contrast, the cumulative frequency of FF  
260 angles gradually increased around the median angle, which we attributed to an expanded range of  
261 true 3D angles owing to the anisotropy of interfacial energy which acts to rotate interfaces into the  
262 lowest energy orientation (Laporte and Provost 2000). In some cases (Figure 3), the cumulative  
263 frequency of FF and FC  $\theta$  largely deviated from the theoretical curve in regions of high apparent  
264  $\theta$ . This deviation was, first, attributed to the relatively small number of measured apparent FF  
265 angles. However, it can also be attributed to the presence of a very large  $\theta$ , possibly associated  
266 with sub-grain boundaries, in which the misorientation angle between two adjacent grains is very  
267 small (Laporte et al. 1997). These two possible reasons caused the cumulative frequency of FF and  
268 FC angles to deviate far from the theoretical prediction compared with that of the CC angle.  
269 Nevertheless, the median angles in such cases were assumed to represent the true 3D value most  
270 frequently occurring in the system because angles smaller than the median fit the theoretical line  
271 well.

#### 272 **3.4. Faceting effect on the median dihedral angle**

273 The angle type dependence of the median  $\theta$  in the olivine–fluid systems at 1–3 GPa and  
274 800–1100 °C is shown in Figure 4. We classified the experimental systems into two groups (I  
275 group: independent angle type; D group: dependent angle type) based on whether the angle type  
276 effect was prominent or not. The angle values were comparable among the CC, FC, and FF in the  
277 H<sub>2</sub>O system at relatively low P–T conditions and in the H<sub>2</sub>O–CO<sub>2</sub> system (I group). In contrast,  
278 the median angle of the facet-bearing angles was higher than that of the CC angle in the H<sub>2</sub>O  
279 system under higher P–T conditions and in the H<sub>2</sub>O–NaCl system (D group). The corresponding  
280 groups for each run product are summarized in Table 1. The duration of experiments in both the I

281 and D groups was sufficiently long for dihedral angle equilibration, and disequilibrium could not  
282 cause their differences. Under constant P–T conditions, FC and CC angles showed a relatively  
283 wide variation in  $\theta$  between the H<sub>2</sub>O and H<sub>2</sub>O–NaCl systems, except for the results at 3 GPa and  
284 1100°C, for which the difference almost disappeared. In contrast, the variation of FF  $\theta$  was  
285 generally limited among the different fluid systems, with a few exceptions in the H<sub>2</sub>O–CO<sub>2</sub> and  
286 H<sub>2</sub>O–NaCl (5.0 wt.% NaCl) systems. The three types of  $\theta$  in the H<sub>2</sub>O–CO<sub>2</sub> system at 1 GPa and  
287 1000°C were larger than those of the other fluid compositions. In the H<sub>2</sub>O–NaCl (5.0 wt.% NaCl)  
288 system, the  $\theta$  values of the FF angle were similar (Figure 4) or smaller by 5°–10° than those in the  
289 H<sub>2</sub>O and H<sub>2</sub>O–NaCl (27.5 wt.% NaCl) systems at the same P–T conditions (Figure 4d and f).

### 290 **3.5. P–T dependence of median dihedral angles in different fluid systems**

291 The P–T dependencies of the median FF and FC angles in the olivine–fluid systems  
292 measured in this study are shown in Figure 5, along with the CC angles reported by Huang et al.  
293 (2019). CC angles decrease with increasing P and T (Figure 5c, f, and i), possibly corresponding  
294 to the increase in olivine solubility and enhanced adsorption of fluid components on the olivine  
295 surface under high P–T conditions (e.g., Holness 1993; Huang et al. 2019). In contrast, the P–T  
296 dependence of the FC and FF angles was not obvious compared to that of the CC angle. In  
297 particular, the FF angles showed a stepwise change in  $\theta$  but had similar  $\theta$  values (Figure 5a, d, and  
298 g). These findings suggest that factors other than solubility variation control  $\theta$  of FF angles (see  
299 Section 4.1 for a more detailed discussion based on EBSD results). It is worth noting that the  
300 median values of the FF angle were larger than, or close to, 60° in most cases.

301

### 302 **3.6. Crystallographic orientation of olivine**

303 Figure 6 shows the representative EBSD maps and pole figures obtained from the five  
304 recovered samples. Figure S4 of the Supporting Information shows the remaining maps and pole  
305 figures. Our samples exhibited a weak (010) CPO (i.e., *b* axis slightly parallel to the compression  
306 direction of piston cylinder), regardless of the P–T conditions or fluid composition. It was not as  
307 intense as the strong CPO that developed in the deformed olivine aggregate (Pommier et al. 2018).  
308 The weak CPO in this study is reasonable because the cell assembly of the piston cylinder  
309 experiments with materials softened at high P–T conditions have been developed to avoid intense  
310 differential stress. Figure 7 shows olivine grains' high-magnification orientation maps and the  
311 corresponding SE images. Further analyses of the crystallographic orientations of the GBPs are  
312 described in Section 4 (Discussion).

### 313 **3.7. GBP distribution and misorientation angles and axes**

314 To clarify the crystallographic orientation of olivine on GBPs, we evaluated the GBPD for  
315 the grain boundaries of the FF, FC, and CC angles on a high-magnification orientation map (Figure  
316 7). In general, GBPs were not always vertical but tended to incline at various degrees with respect  
317 to the polished cross-section of the sample. In the 2D SEM images, we could not identify the  
318 degree of incline for these planes. In previous studies, GBPD was examined by analyzing the large  
319 number of automatically detected grain boundaries in dry polycrystalline systems, and frequently  
320 appearing planes were detected after statistical treatment (Marquardt and Faul 2018). In our study,  
321 the number of measurements was limited because discrimination of angle types requires a careful  
322 observation of each olivine–fluid–olivine triple junction in a high magnification image. To better  
323 constrain the GBPD with a limited number of measurements, we focused on olivine–fluid–olivine

324 triple junctions with apparent dihedral angles lower than the median value + 5°, and assumed that  
325 their grain boundaries were subvertical to the polished section. For example, in the system with  
326 one true  $\theta$  of 60°, 71% of the apparent dihedral angles fell within the range from 0 to 65°, in which  
327 68% of GBPs formed an angle of  $\geq 67^\circ$  with respect to the sectioning plane. Thus, we inferred  
328 representative errors of approximately 23° in our GBPD analyses, although this is a minimum  
329 estimate in a simplified system. Without this dihedral angle constraint, the estimated error is 35°.  
330 More details about the errors of the GBPD analyses are provided in Supporting Information Note  
331 1 and Figures S5–S6. At such triple junctions, we determined the crystallographic orientations of  
332 olivine sharing the assumed vertical GBPs in the nine samples, including both the I and D groups.  
333 The relationship between the two touching crystal planes was not examined in this study. In the  
334 calculation, we used Euler angles derived from the EBSD analysis and the trend of the GBPs with  
335 respect to the horizontal side of the corresponding SEM image. The symmetrically equivalent  
336 olivine orientations obtained for each angle type were stereologically projected in the crystal  
337 reference frame, as shown in Figure 8. Therefore, the number of geometrically different crystal  
338 planes in our analysis can be 16 ( $= 4^2$ ) with an assumed angle interval of 22.5° ( $= 90^\circ/4$ ), which is  
339 equivalent to our representative GBPD error, for the azimuth and elevation angles in the  
340 stereological projection. More than approximately 75 measurements were required to obtain stable  
341 GBPD results.

342 In the I group, low Miller index planes such as {100} and {010} were dominant at grain  
343 boundaries of the FF triple junctions, whereas the GBPs of the CC triple junctions were often  
344 characterized by a higher Miller index such as {203}. Note that the multiples of uniform density  
345 (MUD) spots indexed as {101} and {100} in Figure 8a1 are indistinguishable because of the tilt  
346 of the GBP. In the D group, the GBPs of the FF triple junctions were focused on {010} and {130},



347 which was interpreted as a broad concentration around {010}, and weakly focused on {110}.  
348 Although the differences in these planes exceeded the possible errors of approximately  $\pm 20^\circ$  in  
349 our analyses, the large variation in true  $\theta$  of the FF angles, especially in the D groups (Figure 3b,  
350 c, e, f; Section 3.3), could cause the apparent broadening of the concentration around {010} beyond  
351 the assumed errors. At the CC triple junctions, high Miller index planes were dominant, and the  
352 {100} was relatively weak. The GBPD at the FC triple junctions tended to exhibit mixed  
353 characteristics of the FF and CC results in both groups. Although the amount of the data was  
354 reduced, the GBPs at the faceted and curved sides were separately analyzed at the FC triple  
355 junctions, as shown in Figure 9. As in the FF and CC junctions, the GBPs at the faceted side of FC  
356 junctions were dominated by low Miller index planes such as {101}, {010}, and {110}, whereas  
357 the GBPs at the curved side of FC junctions preferred high Miller index planes such as {320} and  
358 {151}. This indicates that the GBPDs observed in the FF and CC triple junctions are likely true.  
359 The weak CPO developed in the run products did not significantly affect the results.

360 Marquardt et al. (2015) found that fluid-free olivine aggregates have a preferred {100}  
361 plane of the grain boundary, which is different from the preferential appearance of the {100},  
362 {010}, and {101} planes on the grain boundaries at the FF angle in the present study. This  
363 discrepancy is most likely caused by the high  $\phi$  and various fluid compositions used in our study,  
364 and supports the hypothesis that GBPD might be affected by the low fractions of melt, and/or  
365 contiguity and composition of the melt (Marquardt and Faul 2018). The crystal habit of olivine  
366 grown freely in a fluid-rich system is characterized by the dominant {010} plane (Waff and Faul  
367 1992). Previous studies have shown that the {010} plane of olivine has the lowest energy, followed  
368 by the {100} and {001} planes in fluid/melt systems (Deer et al., 2013; Gurmani et al., 2011; de  
369 Leeuw et al., 2000; Watson et al., 1997).

370 The misorientation angles of two adjacent grains selected for GBPD analysis are shown in  
371 Figure 10. The misorientation angles at the FC and CC triple junctions show an asymmetric  
372 unimodal distribution with a peak at  $90^\circ$  (Figure 10). This is consistent with the theoretical  
373 distribution for randomly distributed orthorhombic crystals (e.g., Mackenzie 1958; Morawiec  
374 2010). The misorientation angle at the FF triple junction shows two maxima at  $70^\circ$  and  $90^\circ$ . In  
375 partially molten olivine aggregates, Faul and Fitz Gerald (1999) also found the two maxima at  $60^\circ$   
376 and  $90^\circ$  in the misorientation angle distribution of two touching olivine grains with the melt-free  
377 boundary. However, owing to a relatively small amount of data in our misorientation analysis, we  
378 considered that the misorientation angle distributions were mostly the same among the three angle  
379 types and generally in accord with the random distribution. The misorientation axes distribution at  
380 the FF, FC, and CC triple junctions were almost indistinguishable from a perfectly random  
381 polycrystal distribution (Figure 11). This was similarly confirmed even when the distributions  
382 were re-analyzed for the fraction of the data with misorientation angles around the peaks at  $90^\circ$   
383 and  $70^\circ (\pm 5^\circ)$ .

#### 384 **4. Discussion**

385 To check the minimum number of measurements required to obtain stable GBPD results,  
386 we performed a series of random tests with dataset b3 in Figure 8, in which 216 measurements  
387 were used to make a contour plot. In this test, 25–150 measurements were randomly selected from  
388 the original dataset without duplication, and this was repeated 10 times to obtain 10 individual  
389 stereograms. We found that the distribution of high MUD derived from the original dataset (b3 in  
390 Figure 8) could be well reproduced within the estimated error margins when the random selection  
391 number was  $\geq 75$  (Figure 12 and Figure S7 of the Supporting Information). In our study, the  
392 number of measurements was  $> 75$  in most cases, suggesting the robustness of our GBPD results.

393 However, for the curved and facet sides of the FC angle in the I-group (a2-1 and a2-2 in Figure 9),  
394 the number of measurements was reduced to 40. Random tests with  $\leq 50$  measurement also often  
395 reproduced parts of the high MUD spots seen in the original figure (b3 in Figure 8). Thus, the high  
396 MUD spots in the above separate analyses may represent part of the true distribution.

397 Under high P–T conditions, enhanced olivine solubility significantly decreased the  
398 interfacial energy with the fluid, resulting in an obvious P–T dependence of  $\theta$  for the CC angle  
399 (Huang et al. 2019, 2020). The effect of fluid composition was also prominent in the case of the  
400 CC angle, reflecting the dependence of interfacial energy on fluid compositions (e.g., Holness  
401 1992, 1993). However,  $\theta$  of the FF angle was less sensitive to P–T conditions and fluid  
402 composition than that of the CC angle, with discrete values. Laporte and Provost (2000)  
403 theoretically investigated an anisotropic system and showed that  $\theta$  of the FF angle was controlled  
404 by the crystallographic orientation of two adjacent minerals. The extent of surface adsorption  
405 among different crystallographic orientations can vary depending on P–T conditions and fluid  
406 composition (Kretz 1966), which may switch the dominant facet planes of olivine and explain the  
407 stepwise change of the FF  $\theta$  value in our study.

408 In this study, GBPD analyses revealed that low Miller index planes, such as  $\{100\}$ ,  $\{010\}$ ,  
409 and  $\{101\}$ , preferentially appeared at the GBPs of the FF angle (Figure 8 and Figure 9). As the  
410 faceted mineral–fluid interfaces appeared to have low Miller indices, the  $\theta$  value of the FF angle  
411 could be estimated from the angles between the GBPs of  $\{100\}$ ,  $\{010\}$ ,  $\{101\}$ , and the other low  
412 Miller index olivine surfaces. To test this inference, we calculated the angles between the GBPs  
413 of  $\{100\}$  and  $\{010\}$ , and the interfacial boundary crystal planes (IBCPs; i.e., faceted planes with  
414 fluids) of  $\{001\}$ ,  $\{011\}$ ,  $\{110\}$ ,  $\{101\}$ , and  $\{120\}$  appearing in the ideal habit of olivine crystals

415 (Figure 7e). The calculations were made for asymmetrical configurations in which two touching  
416 grains have different crystallographic orientations, allowing the grain boundary to be defined. A  
417 weak CPO supports the dominance of these configurations in our run products. Triple junctions at  
418 which one extended grain boundary plane acts as one of the mineral–fluid interfaces were  
419 considered as an extreme case (Flat face in Table 2). All the calculated configurations are presented  
420 in Table 1 of the Supporting Information and angles consistent with the experimentally obtained  
421 FF  $\theta$  (i.e., 50°–55°, 55°–65°, 65°–70°, and 75°–80°) are summarized in Table 2. The calculated  
422 candidates cover these experimental values in Table 2. For instance, the measured I type FF  $\theta$  in  
423 the H<sub>2</sub>O–CO<sub>2</sub> system at 1 GPa and 1000°C was 79.2°. This may correspond to the calculated  
424 dihedral angles of 79.2°, 80.1°, and 81.5°, from an asymmetrical triple junction composed of the  
425  $\{100\}_{\text{GB}}\text{--}\{101\}_{\text{IB}}$  and  $\{010\}_{\text{GB}}\text{--}\{021\}_{\text{IB}}$ ,  $\{100\}_{\text{GB}}\text{--}\{110\}_{\text{IB}}$  and  $\{101\}_{\text{GB}}\text{--}\{120\}_{\text{IB}}$ , and  $\{100\}_{\text{GB}}\text{--}$   
426  $\{101\}_{\text{IB}}$  and  $\{100\}_{\text{GB}}\text{--}\{120\}_{\text{IB}}$ , respectively. The misorientation angle and axis distributions at the  
427 FF triple junctions are indicative of nearly random alignment of two touching grains, showing that  
428 the weak CPO developed in the sample had little effect on the FF angles. The  $\theta$  value of the FC  
429 angle may be controlled by both interfacial energy and crystallographic orientation, resulting in  
430 FC angles showing intermediate characteristics between FF and CC angles.

## 431 **5. Implications**

### 432 **5.1. Consequences of faceting on fluid connectivity in an undeformed olivine–fluid system**

433 Our study demonstrates that approximately 1/3 of dihedral angles in the olivine–fluid  
434 system are facet-bearing, irrespective of P–T conditions or fluid composition (Table 1). Fluid pores  
435 surrounded by faceted interfaces struggle to connect with each other, even at  $\theta < 60^\circ$ , which  
436 requires a threshold  $\phi$  for the establishment of a fluid network, as in the case for CC angles of  $>$

437 60° (Price et al. 2006). Thus, a system that includes both curved and faceted interfaces with low  $\phi$ ,  
438 bulk permeability may be reduced. Huang et al. (2021) measured the electrical conductivity of  
439 fluid-bearing forsterite aggregate with various  $\phi$  under textural equilibrium states at 1 GPa and  
440 800°C in the H<sub>2</sub>O–NaCl system with 5.0 wt.% NaCl. The electrical conductivity measurements  
441 and synchrotron X-ray computed CT imaging of the post-run products showed that fluid pores  
442 were not interconnected at  $\phi$  of 0.51 vol.%. In contrast, they started to form the fluid network at  $\phi$   
443 of > 2.14 vol.%. Although the CC angle can be lower than 60° under this experimental P–T  
444 condition (Huang et al. 2019), fluid interconnection was not established at  $\phi$  below approximately  
445 1.0–2.0 vol.%. This is most likely be attributable to the presence of a substantial number of faceted  
446 interfaces that increase  $\phi_c$  and decreases permeability, as pointed out by Price et al. (2006).

447 Toramaru and Fujii (1986) examined the melt connectivity in peridotites composed of  
448 olivine, clinopyroxene, and orthopyroxene based on a bond percolation model with the melt  
449 stability at the grain edges and corners, namely, dihedral angles. They found that the melt was not  
450 stable at pyroxene-dominated grain edges and that the interconnection was established when  
451 pyroxene modal composition was approximately < 25–20 vol.% when the grain size of olivine and  
452 pyroxenes were similar. If we assume that faceting triple junctions hamper fluid interconnection,  
453 an analogous discussion will be possible for fluid connectivity in facet-bearing olivine aggregates.  
454 Given the slightly higher proportion of faceting triple junctions (28%–36%; Table 1) than the  
455 above pyroxene modal composition of Toramaru and Fujii (1986), the electrical conductivity  
456 results of Huang et al. (2021), in which the fluid percolation was prohibited at  $\phi = 0.51$  vol.% but  
457 established at a small critical fraction (2.14 vol.%), seems consistent with the model prediction.

## 458 **5.2. Preferential appearance of faceted fluid pores in sheared mantle**

459 Our study demonstrates that faceted olivine–fluid interfaces are preferentially  
460 accompanied by low Miller index GBPs such as {100}, {010}, and {101}. The extensive  
461 occurrence of faceted crystallographic faces in deep-seated rocks can change bulk permeability,  
462 elastic, anelastic, and electrical properties (Waff and Faul 1992). Waff and Faul (1992)  
463 investigated melt distribution in the texturally equilibrated olivine–melt system and found that the  
464 presence of melt film along the pervasive faceted crystal interface significantly reduced  $\theta$  and  
465 increased the permeability of the bulk rock. However, the effect of the faceted interface on pore  
466 morphology and permeability obtained from olivine–fluid systems in the present study is different  
467 from that of the olivine–melt system. Our results show that grain boundaries associated with the  
468 flat interface are dry, and that FF and FC angles are comparable to or larger than CC angles,  
469 working against the establishment of fluid connectivity. This effect could be magnified under shear  
470 stress, such as within the corner flow of subduction zones. In deformed olivine aggregate with  
471 strong CPO, crystal axes (i.e.,  $a$ ,  $b$ , and  $c$  axes) of olivine grains are aligned in specific directions  
472 depending on temperature, olivine water content, and stress state (Jung and Karato 2001; Kneller  
473 et al. 2005, 2008; Jung et al. 2006; Karato et al. 2008). In addition, Ferreira et al. (2021) found that  
474 deformation not only causes strong CPO, but also increases the proportion of particular grain  
475 boundary planes (e.g., {010}). This may lead to low Miller index planes dominating grain  
476 boundaries, thereby increasing the proportion of facet-bearing angles in sheared mantle.

477 Jung and Karato (2001) examined a water-saturated olivine fabric under shear strain and  
478 found that the  $c$  axis was subparallel to the shear direction, and the  $b$  axis was perpendicular to the  
479 shear direction (B-type fabric). Therefore, in a water-rich subduction zone, a B-type fabric is  
480 expected, in which the  $c$  axis is subparallel to the subduction direction and the  $b$  axis is  
481 perpendicular to the plate interface. Liu and Zhao (2017) detected  $V_p$  anisotropy in the mantle

482 wedge beneath Japan, supporting the hypothesis that the B-type fabric is dominant in the fore-arc  
483 mantle wedge. This type of grain alignment could lead to grain boundaries composed of the {100}  
484 and {010} planes parallel to the subduction direction to form abundant FF angles that are  
485 comparable to or larger than CC angles, decreasing permeability and electrical conductivity along  
486 the subduction direction. That is to say, the presence of faceting may change pore geometry and  
487 direction of fluid flow, contributing to anisotropy of permeability. The preferred fluid flux induced  
488 by faceting in the subduction zone controls flux melting, anomalies of electrical conductivity, and  
489 seismic wave velocity attenuation in subduction systems.

## 490 **6. Conclusions**

491 In this study, we quantitatively determined the effect of faceting on the olivine–fluid  $\theta$  in  
492 different fluid systems ( $\text{H}_2\text{O}$ ,  $\text{H}_2\text{O}\text{--}\text{CO}_2$  ( $X_{\text{CO}_2} = 0.5$ ),  $\text{H}_2\text{O}\text{--}\text{NaCl}$  (5.0 and 27.5 wt.% NaCl)) at 1–  
493 3 GPa and 800–1100 °C. The results show that 1/3 of olivine–fluid  $\theta$  are faceted plane-bearing  
494 angles, regardless of the P–T conditions or fluid composition. Our  $\theta$  measurements show that in  
495 the  $\text{H}_2\text{O}$  systems at relatively low P–T conditions and in the  $\text{H}_2\text{O}\text{--}\text{CO}_2$  system, facet-bearing angle  
496 values (i.e., FF and FC) are comparable to those of the CC angle. However, in the  $\text{H}_2\text{O}$  system at  
497 high P–T conditions and in the  $\text{H}_2\text{O}\text{--}\text{NaCl}$  system, facet-bearing angle values (i.e., FF and FC) are  
498 larger than those of the CC angle. EBSD analyses show that the run products did not have an  
499 intense CPO corresponding to static compression conditions. Strikingly, the GBPD revealed that  
500 faceted and curved interfaces at facet-bearing triple junctions have GBPs with low (e.g., {100},  
501 {010}, {101}) and high (e.g., {130}, {203}, {320}) Miller index faces, respectively. The  
502 misorientation angle/axis distributions of adjacent grain pairs were in accord with a theoretical  
503 distribution of random olivine aggregate. The calculation of  $\theta$  values between two adjacent crystal

504 planes strongly reproduced the measured values of the FF angles, which further supports the results  
505 of our GBPD analyses. Therefore, our results suggest the importance of crystallographic  
506 orientation in determining the origin of the FF angle. The presence of the FF angle and associated  
507 changes in fluid pore morphology require a high fluid fraction for establishing fluid networks. This  
508 further leads to permeability anisotropy and changes in geophysical characteristics, particularly in  
509 mantle wedge settings where olivine CPO is expected.

510

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525



526 **References**

- 527 Allen, C.J. (1972) The role of water in the mantle of the earth: the stability of amphiboles  
528 and micas. 24th International Geological Congress, Montreal, 2, 231–240.
- 529 Angiboust, S., Pettke, T., de Hoog, J.C.M., Caron, B., and Oncken, O. (2014) Channelized  
530 Fluid Flow and Eclogite-facies Metasomatism along the Subduction Shear Zone.  
531 Journal of Petrology, 55, 883–916.
- 532 Cmíral, M., Fitz Gerald, J.D., Faul, U.H., and Green, D.H. (1998) A close look at dihedral  
533 angles and melt geometry in olivine-basalt aggregates: A TEM study. Contributions to  
534 Mineralogy and Petrology, 130, 336–345.
- 535 de Leeuw, N.H., Parker, S.C., Catlow, C.R.A., and Price, G.D. (2000) Modelling the effect  
536 of water on the surface structure and stability of forsterite. Physics and Chemistry of  
537 Minerals 27:5, 27, 332–341.
- 538 Deer et al. (2013) An Introduction to the Rock-Forming Minerals (third edition), 663–664.  
539 The Canadian Mineralogist Vol. 51.
- 540 Elliott, M.T., Cheadle, M.J., and Jerram, D.A. (1997) On the identification of textural  
541 equilibrium in rocks using dihedral angle measurements. Geology, 25, 355–358.
- 542 Faul, U.H. (1997) Permeability of partially molten upper mantle rocks from experiments and  
543 percolation theory. Journal of Geophysical Research: Solid Earth, 102, 10299–10311.
- 544 Faul, U.H., and Fitz Gerald, J.D. (1999) Grain misorientations in partially molten olivine  
545 aggregates: an electron backscatter diffraction study. Physics and Chemistry of Minerals  
546 1999 26:3, 26, 187–197.

- 547 Ferreira, F., Hansen, L.N., and Marquardt, K. (2021) The effect of grain boundaries on plastic  
548 deformation of olivine. *Journal of Geophysical Research: Solid Earth*, 126(7),  
549 e2020JB020273.
- 550 Gurmani, S.F., Jahn, S., Brasse, H., and Schilling, F.R. (2011) Atomic scale view on partially  
551 molten rocks: Molecular dynamics simulations of melt-wetted olivine grain boundaries.  
552 *Journal of Geophysical Research: Solid Earth*, 116, 12209.
- 553 Harker, D., and Parker, E.R. (1945) Grain shape and grain growth. *Transactions of the*  
554 *American Society for Metals*, 34, 156–201.
- 555 Hermann, J., Spandler, C., Hack, A., and Korsakov, A. v. (2006) Aqueous fluids and hydrous  
556 melts in high-pressure and ultra-high pressure rocks: Implications for element transfer  
557 in subduction zones. *Lithos*, 92, 399–417.
- 558 Holness, M.B. (1992) Equilibrium dihedral angles in the system quartz–CO<sub>2</sub>–H<sub>2</sub>O–NaCl at  
559 800°C and 1–15 kbar: the effects of pressure and fluid composition on the permeability  
560 of quartzites. *Earth and Planetary Science Letters*, 114, 171–184.
- 561 Holness, M.B. (1993) Temperature and pressure dependence of quartz-aqueous fluid dihedral  
562 angles: the control of adsorbed H<sub>2</sub>O on the permeability of quartzites. *Earth and*  
563 *Planetary Science Letters*, 117, 363–377.
- 564 Holness, M.B., and Siklos, S.T.C. (2000) The rates and extent of textural equilibration in  
565 high-temperature fluid-bearing systems. *Chemical Geology*, 162, 137–153.
- 566 Huang, Y., Nakatani, T., Nakamura, M., and McCammon, C. (2019) Saline aqueous fluid  
567 circulation in mantle wedge inferred from olivine wetting properties. *Nature*  
568 *communications*, 10, 5557.

- 569 Huang, Y., Nakatani, T., Nakamura, M., and McCammon, C. (2020) Experimental constraint  
570 on grain-scale fluid connectivity in subduction zones. *Earth and Planetary Science*  
571 *Letters*, 552, 116610.
- 572 Huang, Y., Guo, H., Nakatani, T., Uesugi, K., Nakamura, M., and Keppler, H. (2021)  
573 Electrical Conductivity in Texturally Equilibrated Fluid-Bearing Forsterite Aggregates  
574 at 800°C and 1 GPa: Implications for the High Electrical Conductivity Anomalies in  
575 Mantle Wedges. *Journal of Geophysical Research: Solid Earth*, 126, e2020JB021343.
- 576 Iwamori, H. (1998) Transportation of H<sub>2</sub>O and melting in subduction zones. *Earth and*  
577 *Planetary Science Letters*, 160, 65–80.
- 578 Jung, H., and Karato, S.I. (2001) Water-Induced Fabric Transitions in Olivine. *Science*, 293,  
579 1460–1463.
- 580 Jung, H., Katayama, I., Jiang, Z., Hiraga, T. and Karato, S.I. (2006) Effect of water and stress  
581 on the lattice-preferred orientation of olivine. *Tectonophysics*, 421(1–2), 1–22.
- 582 Jurewicz, S.R., and Jurewicz, A.J.G. (1986) Distribution of apparent angles on random  
583 sections with emphasis on dihedral angle measurements. *Journal of Geophysical*  
584 *Research*, 91, 9277.
- 585 Karato, S.I., Jung, H., Katayama, I. and Skemer, P. (2008) Geodynamic significance of  
586 seismic anisotropy of the upper mantle: New insights from laboratory studies. *Annual*  
587 *Review of Earth and Planetary Sciences*, 36(1), 59–95.
- 588 Keppler, H., Wiedenbeck, M. and Shcheka, S.S. (2003) Carbon solubility in olivine and the  
589 mode of carbon storage in the Earth's mantle. *Nature*, 424(6947), 414–416.

- 590 Kneller, E.A., van Keken, P.E., Karato, S.I., and Park, J. (2005) B-type olivine fabric in the  
591 mantle wedge: Insights from high-resolution non-Newtonian subduction zone models.  
592 Earth and Planetary Science Letters, 237, 781–797.
- 593 Kneller, E.A., Long, M.D. and van Keken, P.E. (2008). Olivine fabric transitions and shear  
594 wave anisotropy in the Ryukyu subduction system. Earth and Planetary Science Letters,  
595 268(3–4), 268–282.
- 596 Kretz, R. (1966) Interpretation of the Shape of Mineral Grains in Metamorphic Rocks.  
597 Journal of Petrology, 7, 68–94.
- 598 Laporte, D., and Provost, A. (2000) Equilibrium geometry of a fluid phase in a polycrystalline  
599 aggregate with anisotropic surface energies: Dry grain boundaries. Journal of  
600 Geophysical Research: Solid Earth, 105, 25937–25953.
- 601 Laporte, D., and Watson, E.B. (1995) Experimental and theoretical constraints on melt  
602 distribution in crustal sources: the effect of crystalline anisotropy on melt  
603 interconnectivity. Chemical Geology, 124.
- 604 Laporte, D., Rapaille, C., and Provost, A. (1997) Wetting Angles, Equilibrium Melt  
605 Geometry, and the Permeability Threshold of Partially Molten Crustal Protoliths pp. 31–  
606 54.
- 607 Liu, X., and Zhao, D. (2017) P-wave anisotropy, mantle wedge flow and olivine fabrics  
608 beneath Japan. Geophysical Journal International, 210, 1410–1431.
- 609 Liu, X., Matsukage, K.N., Li, Y., Takahashi, E., Suzuki, T. and Xiong, X. (2018) Aqueous  
610 fluid connectivity in subducting oceanic crust at the mantle transition zone conditions.  
611 Journal of Geophysical Research: Solid Earth, 123(8), 6562–6573.

- 612 Mackenzie, J.K., (1958) Second paper on statistics associated with the random disorientation  
613 of cubes. *Biometrika*, 45(1–2), 229–240.
- 614 Marquardt, K., Rohrer, G.S., Morales, L., Rybacki, E., Marquardt, H., and Lin, B. (2015) The  
615 most frequent interfaces in olivine aggregates: the GBCD and its importance for grain  
616 boundary related processes. *Contributions to Mineralogy and Petrology*, 170, 1–17.
- 617 Marquardt, K., and Faul, U.H. (2018) The structure and composition of olivine grain  
618 boundaries: 40 years of studies, status and current developments. *Physics and Chemistry  
619 of Minerals*, 45(2), 139–172.
- 620 Médard, E., McCammon, C.A., Barr, J.A., and Grove, T.L. (2008) Oxygen fugacity,  
621 temperature reproducibility, and H<sub>2</sub>O contents of nominally anhydrous piston-cylinder  
622 experiments using graphite capsules. *American Mineralogist*, 93, 1838–1844.
- 623 Mibe, K., Fujii, T., and Yasuda, A. (1998) Connectivity of aqueous fluid in the Earth's upper  
624 mantle. *Geophysical Research Letters*, 25, 1233–1236.
- 625 Mibe, K., Fujii, T., and Yasuda, A. (1999) Control of the location of the volcanic front in  
626 island arcs by aqueous fluid connectivity in the mantle wedge. *Nature*, 401, 259–262.
- 627 Morawiec, A. (2010) Volume of intersection of two balls in orientation space. *Acta  
628 Crystallographica Section A: Foundations of Crystallography*, 66(6), 717–719.
- 629 Park, H.H., and Yoon, D.N. (1985) Effect of dihedral angle on the morphology of grains in  
630 a matrix phase. *Metallurgical Transactions A*, 16, 923–928.
- 631 Pommier, A., and Evans, R.L. (2017) Constraints on fluids in subduction zones from  
632 electromagnetic data. *Geosphere*, 13, 1026–1049.

- 633 Pommier, A., Kohlstedt, D.L., Hansen, L.N., Mackwell, S., Tasaka, M., Heidelbach, F., and  
634 Leinenweber, K. (2018) Transport properties of olivine grain boundaries from electrical  
635 conductivity experiments. *Contributions to Mineralogy and Petrology*, 173, 41.
- 636 Price, J.D., Wark, D.A., Watson, E.B., and Smith, A.M. (2006) Grain-scale permeabilities of  
637 faceted polycrystalline aggregates. *Geofluids*, 6, 302–318.
- 638 Smith, C. S. (1948) Grains, phases and interfaces. *Transactions of the American Institute of*  
639 *Mining and Metallurgical Engineers*, 175, 15–51.
- 640 Toramaru, A., and Fujii, N. (1986) Connectivity of melt phase in a partially molten peridotite.  
641 *Journal of Geophysical Research: Solid Earth*, 91, 9239–9252.
- 642 van Keken, P.E., Hacker, B.R., Syracuse, E.M., and Abers, G.A. (2011) Subduction factory:  
643 4. Depth-dependent flux of H<sub>2</sub>O from subducting slabs worldwide. *Journal of*  
644 *Geophysical Research*, 116, B01401.
- 645 von Bargen, N., and Waff, H.S. (1986) Permeabilities, interfacial areas and curvatures of  
646 partially molten systems: Results of numerical computations of equilibrium  
647 microstructures. *Journal of Geophysical Research*, 91, 9261.
- 648 Waff, H.S., and Faul, U.H. (1992) Effects of crystalline anisotropy on fluid distribution in  
649 ultramafic partial melts. *Journal of Geophysical Research*, 97, 9003.
- 650 Wark, D.A., and Watson, E.B. (2000) Effect of grain size on the distribution and transport of  
651 deep-seated fluids and melts. *Geophysical Research Letters*, 27, 2029–2032.
- 652 Watson, E.B., and Brenan, J.M. (1987) Fluids in the lithosphere, 1. Experimentally-  
653 determined wetting characteristics of CO<sub>2</sub>–H<sub>2</sub>O fluids and their implications for fluid  
654 transport, host-rock physical properties, and fluid inclusion formation. *Earth and*  
655 *Planetary Science Letters*, 85, 497–515.

- 656 Watson, E.B., and Lupulescu, A. (1993) Aqueous fluid connectivity and chemical transport  
657 in clinopyroxene-rich rocks. *Earth and Planetary Science Letters*, 117, 279–294.
- 658 Watson, E.B. (1999) Lithologic partitioning of fluids and melts. *American Mineralogist*, 84,  
659 1693–1710.
- 660 Watson, G.W., Oliver, P.M., and Parker, S.C. (1997) Computer simulation of the structure  
661 and stability of forsterite surfaces. *Physics and Chemistry of Minerals*, 25, 70–78.
- 662 Worzewski, T., Jegen, M., Kopp, H., Brasse, H., and Taylor Castillo, W. (2011)  
663 Magnetotelluric image of the fluid cycle in the Costa Rican subduction zone. *Nature*  
664 *Geoscience*, 4, 108–111.
- 665 Yoshino, T., Takei, Y., Wark, D.A., and Watson, E.B. (2005) Grain boundary wetness of  
666 texturally equilibrated rocks, with implications for seismic properties of the upper  
667 mantle. *Journal of Geophysical Research: Solid Earth*, 110, 1–16.
- 668 Yoshino, T., Price, J.D., Wark, D.A., and Watson, E.B. (2006) Effect of faceting on pore  
669 geometry in texturally equilibrated rocks: implications for low permeability at low  
670 porosity. *Contributions to Mineralogy and Petrology*, 152, 169–186.
- 671 Zheng, Y.F., Chen, R.X., Xu, Z., and Zhang, S.B. (2016) The transport of water in subduction  
672 zones. *Science China Earth Sciences*, 59(4), 651–682.
- 673
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678 **Figure captions**

679 **Figure 1. Representative scanning electron microscope (SEM) images and normalized**  
680 **grain size distribution of run products. a** Backscattered electron (BSE) image of the run  
681 product in the olivine–H<sub>2</sub>O system at 1000°C and 1 GPa for 120 h. **b** High-magnification  
682 secondary electron (SE) image of a typical triple junction in the olivine–H<sub>2</sub>O system at  
683 1000°C and 1 GPa 120 h, which shows apparent FF angles. **c** Backscattered electron (BSE)  
684 image of the run product in the H<sub>2</sub>O–CO<sub>2</sub> system ( $X_{\text{CO}_2} = 0.5$ ) at 1000°C and 1 GPa for 211  
685 h. **d** High-magnification secondary electron (SE) image of the typical apparent  $\theta$  in the H<sub>2</sub>O–  
686 CO<sub>2</sub> system ( $X_{\text{CO}_2} = 0.5$ ) at 1000°C and 1 GPa for 211 h, which shows the coexistence of  
687 three types of apparent  $\theta$ . **e** Backscattered electron (BSE) image of the run product in the  
688 H<sub>2</sub>O–NaCl system (salinity = 27.5 wt.%) at 2 GPa and 1000°C for 120 h. **f** High-  
689 magnification secondary electron (SE) image of typical apparent  $\theta$  in the H<sub>2</sub>O–NaCl system  
690 (salinity = 27.5 wt.%) at 2 GPa and 1000°C for 120 h. The run products are composed of  
691 mineral grains and epoxy resin-filled pores previously filled with aqueous fluids during the  
692 experiment. Orange marks denote olivine–olivine–olivine triple junctions with intersection  
693 angles of approximately 120°, indicating the attainment of textural equilibrium. White single  
694 and double arrows represent curved and faceted interfaces, respectively. For the CC, FC, and  
695 FF angles, the interfaces in **b**, **d**, and **f** are highlighted by cyan curves. Green and red  
696 rectangles illustrate suitable and unsuitable angles (i.e., with a crack) for measurement,  
697 respectively. **g** Histogram of normalized grain size distribution in the H<sub>2</sub>O system at 1 GPa  
698 and 1000°C. **h** Histogram of normalized grain size distribution in the H<sub>2</sub>O–NaCl system  
699 (salinity = 27.5 wt.%) at 2 GPa and 1000°C. The grain size distribution was normalized by  
700 the mean grain size of the recovered sample. The grain size peak is concentrated around the



701 mean grain size. The mean grain size and measured grain number are shown in the panel.  
702 Abbreviations: ol = olivine, FF = faceted–faceted angle, FC = faceted–curved angle, CC=  
703 curved–curved angle.

704 **Figure 2. Faceting-bearing angle proportion in the olivine–fluid system.** Results are  
705 calculated from 19 run products with a 10.0 vol.% fluid fraction. The average of 19 values is  
706 shown in the panel. The analytical error of each value is 0.5%. The P–T condition and fluid  
707 composition for each sample are shown in the panel.

708 **Figure 3. Representative cumulative frequency curves of measured apparent dihedral**  
709 **angles ( $\theta$ ) in the olivine–fluid system at 2 GPa and 800–1000 °C.** P–T conditions and fluid  
710 composition are shown at the top of each panel. Facet-bearing angles (FC and FF) were  
711 measured in this study, while data for the CC angles are from Huang et al. (2019, 2020). The  
712 median angle value and number (N) of the measured angles are shown for each case. Thin,  
713 solid curves represent theoretical cumulative frequency curves obtained from the isotropic  
714 system with one true  $\theta$ . Abbreviations: FF = faceted–faceted angle, FC = faceted–curved  
715 angle, CC = curved–curved angle.

716 **Figure 4. Faceting dependence of median dihedral angles ( $\theta$ ) in the olivine–fluid system**  
717 **at 1–3 GPa and 800–1100 °C.** Colors in each panel denote fluid compositions. CC data are  
718 from Huang et al. (2019, 2020). An error bar of  $\pm 1.5^\circ$  is shown in the panels. Abbreviations:  
719 FF = faceted–faceted angle, FC = faceted–curved angle, CC = curved–curved angle, I =  
720 independent angle type, D = dependent angle type.

721 **Figure 5. Pressure and temperature dependence of median dihedral angles ( $\theta$ ) in the**  
722 **olivine–fluid system. a–f** Pressure dependence at 800–1000 °C. **g–i.** Temperature  
723 dependence at 2 GPa. P–T conditions and fluid composition are shown in each panel. CC data  
724 are from Huang et al. (2019). An error bar of  $\pm 1.5^\circ$  is shown along with the median angle.  
725 The blue dash line represents  $\theta$  of  $60^\circ$ . Abbreviations: FF = faceted–faceted angle, FC =  
726 faceted–curved angle, CC = curved–curved angle.

727 **Figure 6. Representative EBSD maps and corresponding pole figures. a1–e1** Raw EBSD  
728 maps of recovered olivine aggregate in olivine–fluid systems. Small points within grains are  
729 attributed to noise, crystal defects, and fluid inclusions. **a2–e2** Denoised EBSD maps of  
730 recovered olivine aggregate in olivine–fluid systems. Points smaller than 1  $\mu\text{m}$  have been  
731 removed. **a3–e3** Pole figures showing the crystallographic orientations of (100), (010), and  
732 (001) corresponding to **a2–e2**. Intensities in the color bar are multiples of the uniform  
733 distribution (MUD). All grains defined by different colors are olivine with different  
734 orientations. P–T conditions and fluid composition are shown at the top of **a–e**. The arrow on  
735 the left represents the compaction direction (parallel to the direction of piston movement).

736 **Figure 7. Olivine grains' representative 3D crystal orientation and corresponding**  
737 **secondary electron image in the olivine–fluid system. Raw** EBSD maps in the  $\text{H}_2\text{O}$  system  
738 at 1 GPa and 1000°C (**a1**), in the  $\text{H}_2\text{O}-\text{CO}_2$  ( $X_{(\text{CO}_2)}=0.5$ ) system at 1 GPa and 1000°C (**b1**), in  
739 the  $\text{H}_2\text{O}$  system at 2 GPa and 1000°C (**c1**), and in the  $\text{H}_2\text{O}-\text{NaCl}$  (27.5 wt.%) system at 2 GPa  
740 and 1000°C (**d1**). **a2–d2**, Denoised EBSD maps with 3D crystal orientations of olivine  
741 corresponding to **a1–d1**. Colored areas denote olivine grains with different orientations, and  
742 white areas represent noise, defects, fluid inclusions, and fluid pools. The 3D crystal

743 orientation was visualized for each grain in the high magnification images by showing the  
744 oriented olivine crystal with an idealized morphology. Even though the grain configurations  
745 in **a2–d2** were slightly altered during EBSD data processes (MTEX MATLAB toolbox) such  
746 as denoising and binarization, this essentially has no effect on orientation identification. **a3–**  
747 **d3** Secondary electron (SE) images of recovered olivine aggregates corresponding to **a1–d1**.  
748 Olivine grains have a grey color; fluid pools are black areas that are sometimes filled by resin.  
749 An orange outline visually emphasizes the interface between the olivine and fluids, and  
750 apparent angles with a sequenced number are examples of analyzed angles and corresponding  
751 grain boundaries. **a4–d4** Dihedral angle types, and their values in terms of FF, FC, and CC  
752 corresponding to measured angles in **a3–d3**. These angles have values close to the  
753 corresponding median  $\theta$ . **e** Crystal habit of a single olivine crystal derived using the MTEX  
754 MATLAB toolbox. The blue squares in **a3**, **c3**, and **d3** represent  $\theta$  that has one shared flat  
755 plane for both the grain boundary and interfacial boundary. The white circles in **b3** and **d3**  
756 represent the angle defined by interfaces of one grain/sub-grain, which was not included in  
757 our discussion. Abbreviations: ol = olivine, FF = faceted–faceted angle, FC = faceted–curved  
758 angle, CC = curved–curved angle.

759 **Figure 8. Pole figure of grain boundary plane distribution. a1–a3** Grain boundary plane  
760 distribution (GBPD) in the I group (i.e., systems where  $\theta$  values are independent of angle  
761 type). **b1–b3** Grain boundary plane distribution (GBPD) in the D group (i.e., systems where  
762 faceting increases  $\theta$ ). The equivalent olivine orientations obtained in the analysis were  
763 rearranged into the first quadrant to better highlight the results. Subsequently, by assuming  
764 that the grain boundary planes were equally distributed in each quadrant, the orientations in  
765 the first quadrant were copied in the other quadrants to obtain the pole figure. The analyzed

766 number (N) in the first quadrant is shown at the lower left in each pole figure. MUD is the  
767 multiples of uniform density, shown by the color bar's intensities. The Miller indices were  
768 marked around the high MUD. The assumed error is  $\pm 20^\circ$ . Abbreviations: FF = faceted–  
769 faceted angle, FC = faceted–curved angle, CC = curved–curved angle, I = independent angle  
770 type, D = dependent angle type.

771 **Figure 9. Pole figure of FC angles.** Pole figure of grain boundary plane distributions  
772 (GBPDs) of faceted (**a2–1** and **b2–1**) and curved (**a2–2** and **b2–2**) sides at the FC triple  
773 junction for the I group (i.e., the systems where the  $\theta$  values are independent of the angle  
774 type; **a2–1** and **a2–2**) and D group (i.e., the systems where the faceting increases  $\theta$ ; **b2–1** and  
775 **b2–2**). The data were plotted in the same way as in Figure 8. The analyzed number (N) in the  
776 first quadrant is shown at the lower left in each pole figure. MUD represents the multiples of  
777 uniform density, which shows the intensities in the color bar. The Miller indices were marked  
778 around the high MUD. The assumed error is  $\pm 20^\circ$ . Abbreviations: FC = faceted–curved angle,  
779 F = grain boundary at the faceting side, C = grain boundary at the curved side, I = angle type  
780 independent group, D = angle type dependent group.

781 **Figure 10. Distribution of misorientation angles.** The misorientation angle between each  
782 pair of two measured olivine grains was calculated based on the Euler angle derived from  
783 EBSD. **a** Misorientation angle of the FF type angle. **b** Misorientation angle of the FC type  
784 angle. **c** Misorientation angle of the CC type angle. The red curve represents the theoretical  
785 random distribution of olivine grains. The number of angles (N) is shown in the panel.

786 **Figure 11. Misorientation axis distribution.** All axes of measured grain pairs were mapped  
787 into a unit quadrant owing to symmetry equivalence. **a** All misorientation axis of the FF type

788 angle. **b** All misorientation axis of the FC type angle. **c** All misorientation axis of the CC type  
789 angle. **d** Theoretical misorientation axis distribution for a perfectly random polycrystal based  
790 on the analytical solution (Mackenzie 1958). The number of angles (N) is shown in the panel.  
791 The scale is in units of multiples of uniform distribution (MUD). The highest MUD is the  
792 most frequent axis (121).

793 **Figure 12. Random test of the DCC dataset (b3 in Figure 8).** **a** Test repeated 10 times with  
794 50 randomly selected measurements. **b** Test repeated 10 times with 75 randomly selected  
795 measurements. **c** Test repeated 10 times with 125 randomly selected measurements. The  
796 number of random selections (RN) is shown for each series of tests. The scale is in units of  
797 multiples of uniform distribution (MUD). The assumed error is  $\pm 20^\circ$ .

#### 798 **Supplementary figure captions**

799 **Figure S1. Histogram of normalized grain size distribution in the olivine–fluid system.**  
800 The grain size distribution was normalized by the mean grain size of the recovered sample.  
801 The grain size peak is concentrated around the mean grain size. The mean grain size and  
802 measured grain number are shown in the panel. Abbreviation: ol = olivine.

803 **Figure S2. Cumulative frequency curves of measured apparent dihedral angles ( $\theta$ ) in**  
804 **olivine–fluid systems at 1–3 GPa and 800–1100 °C.** The median value and number (N) of  
805 the measured angles are shown for each experimental condition. The solid lines represent the  
806 theoretical cumulative frequency curves of the isotropic system with one true  $\theta$  (Jurewicz and  
807 Jurewicz 1986). This angle is assumed to coincide with the obtained median value. P–T  
808 conditions and fluid composition are shown for each system. Abbreviations: ol=olivine, FFT

809 = faceted–faceted angle, FCT = faceted–curved angle, CCT= curved–curved angle, AllT= all  
810 types of measured angle.

811 **Figure S3. Frequency distribution histograms of measured apparent dihedral angles ( $\theta$ )**  
812 **in olivine–fluid systems at 1–3 GPa and 800–1100 °C.** Theoretical distributions (orange  
813 curves) for mono–mineral and isotropic systems are also shown in the histograms along with  
814 the median values (Jurewicz and Jurewicz 1986). The P–T and fluid composition are shown  
815 for each system. Abbreviations: ol=olivine, FFT = faceted–faceted angle, FCT = faceted–  
816 curved angle, CCT= curved–curved angle, AllT= all types of measured angle.

817 **Figure S4. EBSD maps and corresponding pole figures under static compression**  
818 **conditions. a1–e1** Raw EBSD maps of recovered olivine aggregate in olivine–fluid systems.  
819 Small points within grains are attributed to noise, crystal defects, and fluid inclusions. **a2–e2**  
820 Denoised EBSD maps corresponding to **a1–e1**. Points smaller than 1  $\mu\text{m}$  have been removed.  
821 All grains defined by different colors are olivine with different orientations. **a3–e3** Pole  
822 figures showing the crystallographic orientation of (100), (010), and (001) corresponding to  
823 **a2–e2**. Color intensities are multiples of the random distribution (MUD). P–T conditions and  
824 fluid composition are shown along with the corresponding system. Abbreviation: ol=olivine.

825 **Figure S5. Schematic olivine–olivine–fluid triple junction with a sectioning plane after**  
826 **Harker and Parker (1945) and Jurewicz and Jurewicz (1986).** True dihedral angle ( $\theta$ )  
827 formed by two olivine–fluid interfaces (pale blue planes).  $Y$  is the apparent dihedral angle  
828 observed on the sectioning plane (pale orange plane). The bold red line represents the unit  
829 normal of the sectioning plane defined in the angular coordinates  $Q$  and  $\phi$ .  $F$  is the angle  
830 formed by the sectioning plane and grain boundary plane (deep sky blue plane).

831 **Figure S6. Sectioning calculation at the olivine–olivine–fluid triple junction. a** Contours  
832 of the apparent dihedral angle,  $Y$  in the  $\sin^2Q$  versus  $\phi$  diagram calculated according to Harker  
833 and Parker (1945) assuming a true dihedral angle,  $\theta$  of  $60^\circ$ . **b** Contours of the angle formed  
834 by the grain boundary plane and sectioning plane,  $F$ , in the  $\sin^2Q$  versus  $\phi$  diagram. **c** Area of  
835  $F \geq 67^\circ$  within the  $Y$  window of  $0^\circ$ – $65^\circ$  in the  $\sin^2Q$  versus  $\phi$  diagram (orange). The ratio of  
836 this area to the area of  $Y = 0^\circ$ – $60^\circ$  (orange + pale yellow) yields the probability of  $F \geq 67^\circ$  in  
837 the selected  $Y$  window.

838 **Figure S7. Random test of the DCC dataset (b3 in Figure 8). a** Test repeated 10 times with  
839 25 randomly selected measurements. **b** Test repeated 10 times with 100 randomly selected  
840 measurements. **c** Test repeated 10 times with 150 randomly selected measurements. The  
841 number of random selections (RN) is shown for each series of tests. The scale is in units of  
842 multiples of uniform distribution (MUD). The assumed error is  $\pm 20^\circ$ .

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## 844 **Supplementary Information**

### 845 **Supplementary Note 1. Errors in GBPD analyses**

846 Our GBPD analyses focused on the olivine–olivine–fluid triple junction with apparent  
847 dihedral angles lower than the median value +  $5^\circ$  and assumed a vertical grain boundary plane.  
848 Based on a simple theoretical calculation, we show that grain boundary planes at such triple  
849 junctions are dominantly subvertical with respect to the polished section. Although we cannot  
850 exactly determine the extent of grain boundary plane tilting in cross-sectional images, the  
851 apparent dihedral angles can be used to constrain the extent of tilting statistically.

852           Following the method of Harker and Parker (1945), we can calculate the apparent  
853           dihedral angle,  $Y$  on an arbitrary sectioning plane at the mineral–mineral–fluid triple junction  
854           in an isotropic system, with one true dihedral angle  $\theta$ . This method is identical to that used to  
855           compute the theoretical cumulative frequency curve of the apparent dihedral angle, as shown  
856           in Figure 3. A schematic of the triple junction with a sectioning plane is shown in Figure S5.  
857           The unit normal of the sectioning plane is defined in angular coordinates  $Q$  and  $\phi$  ( $Q, \phi = 0^\circ$ –  
858            $90^\circ$ ), and  $Y$  is a function of  $\theta$ ,  $Q$ , and  $\phi$  (Harker and Parker 1945). In Figure S6a, the contours  
859           of  $Y$  for a representative  $\theta$  of  $60^\circ$  are shown in the  $\sin^2Q$  versus  $\phi$  diagram. In this diagram,  
860           the area fraction of angles  $\leq Y$  corresponds to the probability that the observed apparent  
861           dihedral angles become  $\leq Y$  (Harker and Parker 1945). The apparent dihedral angles around  
862            $\theta$  were more likely to be observed on the polished section than the other angles. The median  
863           of the  $Y$  values closely corresponds to  $\theta$  (Jurewicz and Jurewicz 1986). We noted that a  $Y$   
864           value smaller than  $\theta$  required a smaller  $\phi$ , and vice versa. Increasing  $\sin^2Q$  (i.e.,  $Q$ ) tended to  
865           cause  $Y$  values to deviate from the median (i.e.,  $\theta$ ).

866           The angle between the grain boundary plane and the arbitrary sectioning plane,  $F$  ( $F$   
867            $= 0^\circ$ – $90^\circ$ ) can be calculated from their normals.  $F$  is dependent on  $Q$  and  $\phi$ , but independent  
868           of  $\theta$  (Figure S1). In Figure S6b, the contours of  $F$  are shown in the  $\sin^2Q$  versus  $\phi$  diagram.  
869           As in the case of  $Y$ , the area fraction of angles of  $\geq F$  should correspond to the probability  
870           that the angles become  $\geq F$ . At  $\sin^2Q$  (i.e.,  $Q$ ) = 0 or  $\phi = 0$ , the grain boundary plane is vertical  
871           ( $F = 90^\circ$ ). With increasing  $Q$  and  $\phi$ ,  $F$  tends to deviate from  $90^\circ$ ; that is, the grain boundary  
872           plane becomes tilted. Therefore, subvertical (i.e.,  $F$  close to  $90^\circ$ ) grain boundary planes can  
873           be expected at triple junctions with  $Y$  smaller than the median, because such  $Y$  values can  
874           only be observed at low  $\phi$ .



875           Combining the Y and F contours in the  $\sin^2Q$  versus  $\phi$  diagram allows us to compute  
876 the probability of observing sub-vertical grain boundary planes at triple junctions in an  
877 arbitrary Y window on the polished section. We regarded the minimum deviation of F from  
878  $90^\circ$ , which satisfies probability of more than approximately 68%, as the representative error  
879 (1s) of our GBPD analyses. In Figure S6c, the area of  $F \geq 67^\circ$  in our preferred Y window  
880 from  $0^\circ$  to  $65^\circ$  (i.e., median +  $5^\circ$ ) is shown in the  $\sin^2Q$  versus  $\phi$  diagram for  $\theta = 60^\circ$ . We  
881 found that 71% of the apparent dihedral angles fell within the range of  $0^\circ \leq Y \leq 65^\circ$ , in which  
882 68% of grain boundary planes formed an angle  $\geq 67^\circ$  with respect to the sectioning plane.  
883 Thus, we inferred a representative error of approximately  $23^\circ$  in our GBPD analyses.  
884 Although this value slightly increased and decreased at lower and higher  $\theta$ , respectively, it  
885 was not significantly dependent on  $\theta$  in the range of interest ( $23^\circ$ – $24^\circ$  at  $\theta = 50^\circ$ – $80^\circ$ ). If we  
886 do not use dihedral angle constraints (i.e., a Y window of  $0^\circ$ – $180^\circ$ ), the probability of  $F \geq 67^\circ$   
887 decreases to 49% and the estimated error becomes  $35^\circ$ .  
888

889        **References**

890        Harker, D., and Parker, E.R. (1945) Grain shape and grain growth. Transactions of the  
891                American Society for Metals, 34, 156–201.

892        Jurewicz, S.R., and Jurewicz, A.J.G. (1986) Distribution of apparent angles on random  
893                sections with emphasis on dihedral angle measurements. Journal of Geophysical  
894                Research, 91, 9277.

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Table 1. Experimental conditions and results

Run no.	P (GPa)	T (°C)	Duration (h)	Fluid phase	Salinity (wt%)	X(CO <sub>2</sub> )	Statistic angle number for faceting ratio	Faceting proportion ( % )	EBSD	GBPD	Angle Type	Number of angle measured	Median (Degree)	I&D group
1	1	800	192	H <sub>2</sub> O	0	-	187	32.6	-	-	FF	112	70.8	I
											FC	174	70.4	
											CC <sup>a</sup>	200	71.2	
											All	486	71.0	
2	1	800	192	H <sub>2</sub> O-NaCl	5	-	218	31.9	-	-	FF	137	68.4	D
											FC	176	63.1	
											CC <sup>a</sup>	222	58.5	
											All	535	62.0	
3	1	800	192	H <sub>2</sub> O-NaCl	27.5	-	309	32.9	-	-	FF	120	66.3	D
											FC	126	62.3	
											CC <sup>a</sup>	207	58.3	
											All	453	60.6	
4	1	1000	120	H <sub>2</sub> O	0	-	509	30.4	√	√	FF	134	63.6	I
											FC	143	64.2	
											CC <sup>a</sup>	210	64.1	
											All	487	63.8	
5	1	1000	120	H <sub>2</sub> O-CO <sub>2</sub>	0	0.5	383	32.1	√	√	FF	148	79.2	I
											FC	171	78.4	
											CC <sup>b</sup>	213	78.2	
											All	532	78.5	
6	1	1000	120	H <sub>2</sub> O-NaCl	27.5	-	302	32.0	√	√	FF	115	62.0	D
											FC	174	59.7	
											CC <sup>a</sup>	208	56.1	
											All	497	57.8	
7	2	800	212	H <sub>2</sub> O	0	-	205	31.8	√	-	FF	136	68.7	I
											FC	178	68.8	
											CC <sup>a</sup>	212	69.9	
											All	526	69.1	
8	2	800	192	H <sub>2</sub> O-NaCl	5	-	201	33.3	√	√	FF	118	64.3	D
											FC	172	57.9	
											CC <sup>a</sup>	247	53.6	
											All	537	56.4	
9	2	800	210	H <sub>2</sub> O-NaCl	27.5	-	269	31.3	√	√	FF	107	66.5	D
											FC	178	59.0	
											CC <sup>a</sup>	210	56.4	
											All	495	59.0	
10	2	1000	120	H <sub>2</sub> O	0	-	582	33.5	√	√	FF	102	58.4	I
											FC	249	57.5	
											CC <sup>a</sup>	213	58.0	
											All	564	57.8	
11	2	1000	133	H <sub>2</sub> O-NaCl	5	-	314	32.8	√	√	FF	101	50.7	D
											FC	188	46.5	
											CC <sup>a</sup>	201	43.4	
											All	490	45.0	
12	2	1000	120	H <sub>2</sub> O-NaCl	27.5	-	136	33.8	√	√	FF	120	57.0	D
											FC	159	55.6	
											CC <sup>a</sup>	304	51.7	
											All	583	53.7	
13	2	1100	72	H <sub>2</sub> O	0	-	284	34.2	-	-	FF	134	59.1	D
											FC	153	56.9	
											CC <sup>b</sup>	210	53.8	
											All	497	55.9	
14	2	1100	72	H <sub>2</sub> O-NaCl	27.5	-	201	30.0	-	-	FF	104	58.3	D
											FC	224	52.6	
											CC <sup>a</sup>	202	47.1	
											All	530	50.2	
15	3	800	211	H <sub>2</sub> O	0	-	100	35.8	-	-	FF	123	64.9	I
											FC	135	65.0	
											CC <sup>a</sup>	200	66.2	
											All	458	66.0	
16	3	800	192	H <sub>2</sub> O-NaCl	5	-	166	33.6	-	-	FF	107	55.0	D
											FC	132	53.2	
											CC <sup>a</sup>	242	50.2	
											All	481	52.4	
17	3	800	211	H <sub>2</sub> O-NaCl	27.5	-	157	36.9	-	-	FF	103	64.0	D
											FC	164	60.2	
											CC <sup>a</sup>	200	56.4	
											All	467	60.0	
18	3	1000	120	H <sub>2</sub> O	0	-	344	28.3	-	-	FF	110	60.6	D
											FC	199	54.7	
											CC <sup>a</sup>	205	49.2	
											All	514	53.7	
19	3	1000	120	H <sub>2</sub> O-NaCl	27.5	-	158	30.0	√	√	FF	102	61.4	D
											FC	147	55.5	
											CC <sup>a</sup>	220	47.4	
											All	469	51.6	

913 Note: The fluid fraction in each experiment was ~10.0 vol.%. The true  $\theta$  value for each system  
 914 was a median value  $\pm 1.5^\circ$ . The analytical error for the faceting proportion was 0.5%. **a** the  
 915 CC angle data is cited from Huang et al. (2019); **b** the CC angle data is cited from Huang et  
 916 al. (2020). The run products that were employed for the electron backscattered diffraction  
 917 (EBSD) and grain boundary plane distribution (GBPD) are marked in the table.  $X_{(\text{CO}_2)}$  =  
 918  $\text{CO}_2/(\text{H}_2\text{O}+\text{CO}_2)$  in mole. Abbreviations: ol = olivine, FF = faceted–faceted angle, FC =  
 919 faceted–curved angle, CC = curved–curved angle, All = all measured angles, I = angle type  
 920 independent, D = angle type dependent.

921 **Table 2. Theoretical FF angle between two crystal planes**

Crystal	GBCP	IBCP	ACP( $^\circ$ )	Calculated FF Angle( $^\circ$ )	Configuration	Measured FF angle( $^\circ$ )
C1	101	001	51.5	51.5	Flat face	50–55
C1	101	120	55.1	55.1	Flat face	
C1	100	110	25.0	63.5	Asymmetrical	55–65
C2	100	101	38.5			
C1	100	110	25.0	65.7	Asymmetrical	
C2	010	021	40.7			
C1	101	021	66.2	66.2	Flat face	65–70
C1	100	110	25.0	68.0	Asymmetrical	
C2	100	120	43.0			
C1	100	110	25.0	69.8	Asymmetrical	
C2	101	110	44.8			
C1	100	110	25.0	76.5	Asymmetrical	
C2	101	001	51.5			
C1	100	101	38.5	79.2	Asymmetrical	75–80
C2	010	021	40.7			
C1	100	110	25.0	80.1	Asymmetrical	
C2	101	120	55.1			
C1	100	101	38.5	81.5	Asymmetrical	
C2	100	120	43.0			

922

923 Note: The calculated FF angle was obtained by summing the two angles between two crystal  
 924 planes (ACP). This configuration shows the geometry of the calculated dihedral angle. We  
 925 assume that the two crystals (crystal 1, C1; crystal 2, C2) are touching with different (i.e.,  
 926 asymmetrical configuration) low–Miller index planes to form the FF angle. Additionally, we  
 927 show extreme cases where one flat plane is shared for both grain boundary and interfacial  
 928 boundary (i.e., flat face). GBCP, grain boundary crystal plane; IBCP, interfacial boundary  
 929 crystal plane. The crystal cell parameters ( $a=4.7540 \text{ \AA}$ ,  $b=10.1971 \text{ \AA}$ , and  $c=5.9806 \text{ \AA}$ )  
 930 employed for the angle calculation were cited from Deer et al. (2013).

### 931 Supplementary Table

932 **Table S1. List of calculated FF–type dihedral angles formed by low–Miller Index grain**  
 933 **boundary planes and interfaces with fluid at a triple junction**

Grain Boundary Plane	Faceted Interface												
		C1	C2	25.0	43.0	38.5	47.0	40.7	66.2	55.1	51.5	44.8	Flat face
100	110	25.0	<b>50.0</b>	68.0	63.5	72.0	65.7	91.2	80.1	76.5	69.8	25.0	
100	120	43.0	68.0*	<b>86.0</b>	81.5	90.0	83.7	109.2	98.1	94.5	87.8	43.0	
100	101	38.5	63.5*	81.5*	<b>77.0</b>	85.5	79.2	104.7	93.6	90.0	83.3	38.5	
010	120	47.0	72.0	90.0	85.5	<b>94.0</b>	87.7	113.2	102.1	98.5	91.8	47.0	
010	021	40.7	65.7*	83.7	79.2*	87.7	<b>81.4</b>	106.9	95.8	92.2	85.5	40.7	
101	021	66.2	91.2	109.2	104.7	113.2	106.9	<b>132.4</b>	121.3	117.7	111	66.2*	
101	120	55.1	80.1*	98.1	93.6	102.1	95.8	121.3	<b>110.2</b>	106.6	99.9	55.1*	
101	001	51.5	76.5*	94.5	90.0	98.5	92.2	117.7	106.6	<b>103</b>	96.3	51.5*	
101	110	44.8	69.8*	87.8	83.3	91.8	85.5	111.0	99.9	96.3	<b>89.6</b>	44.8	

934  
 935 Note: Bold represents symmetric configuration, which are not dihedral angles because a grain  
 936 boundary cannot be defined between the two adjacent grains with the same crystallographic  
 937 orientation (i.e., symmetrical configurations). The star superscript represents calculated  
 938 angles that are consistent with experimental data.

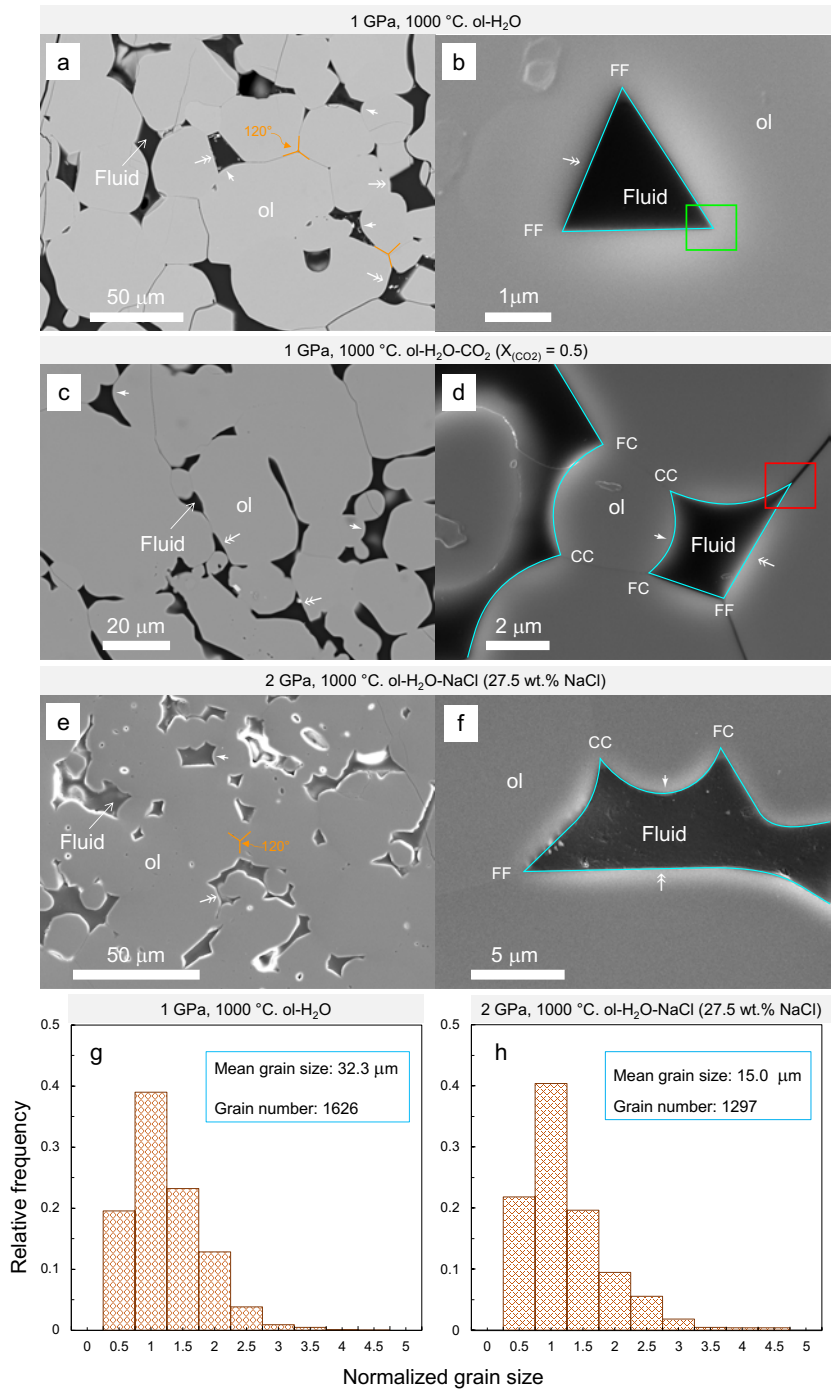
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942 **Figures**

943 **Figure 1**

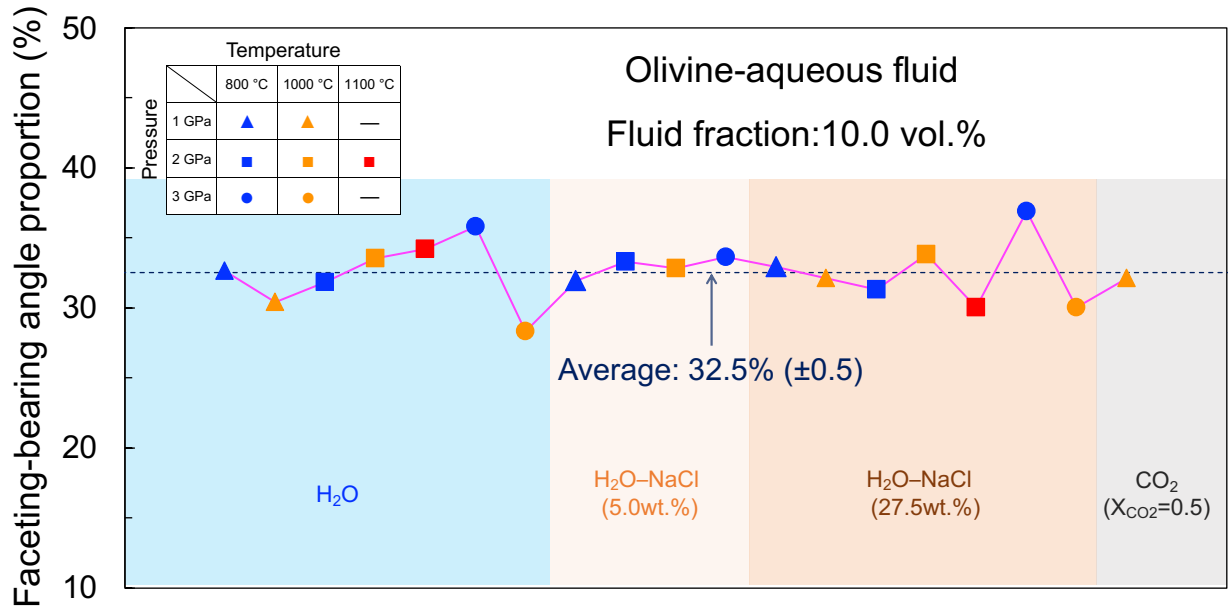


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Figure 2



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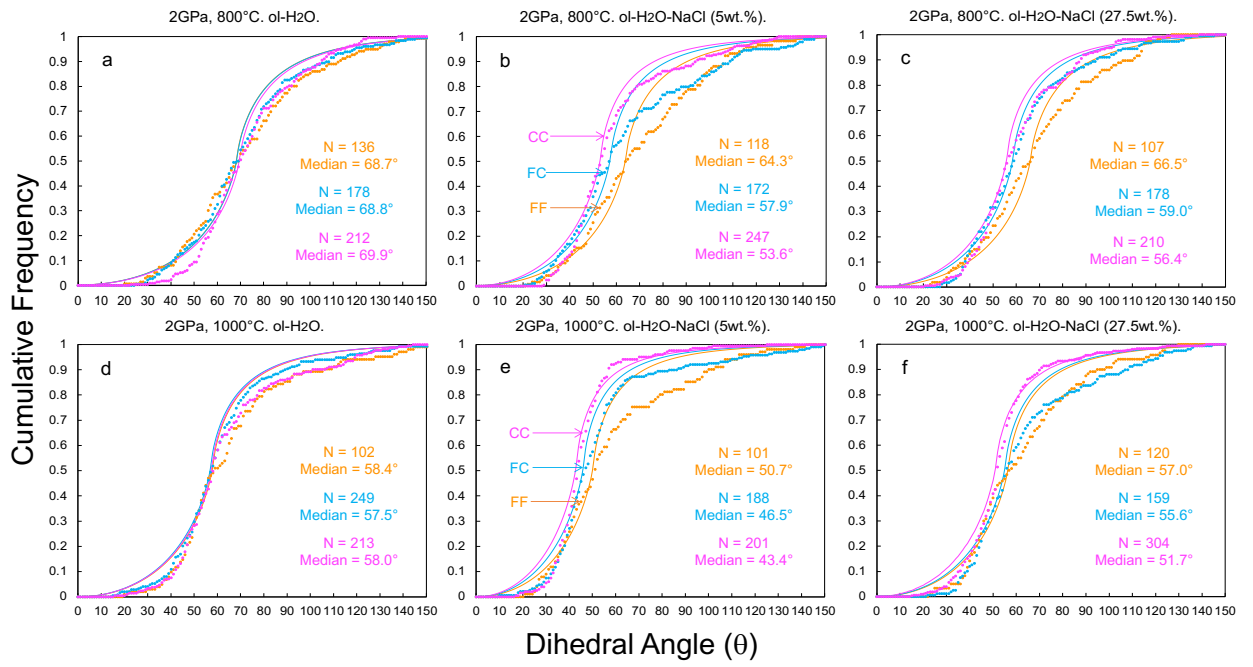
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Figure 3



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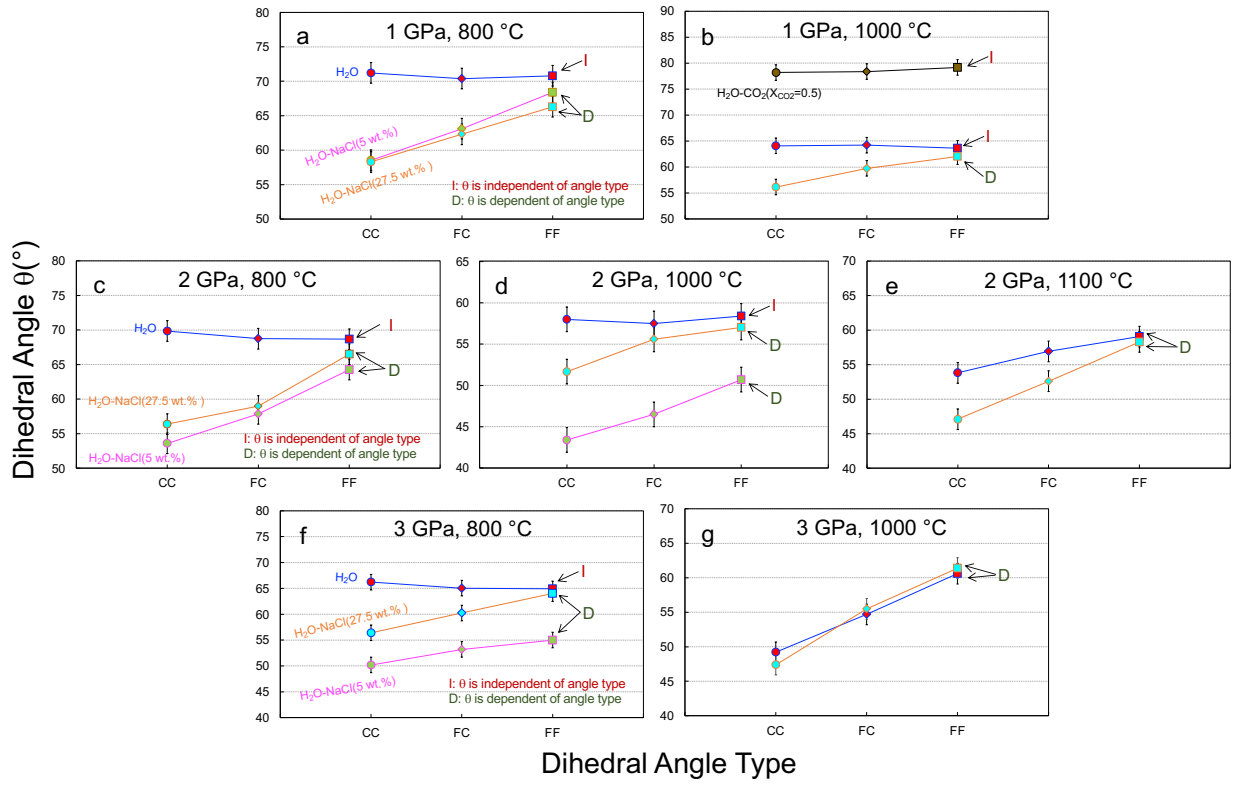
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Figure 4



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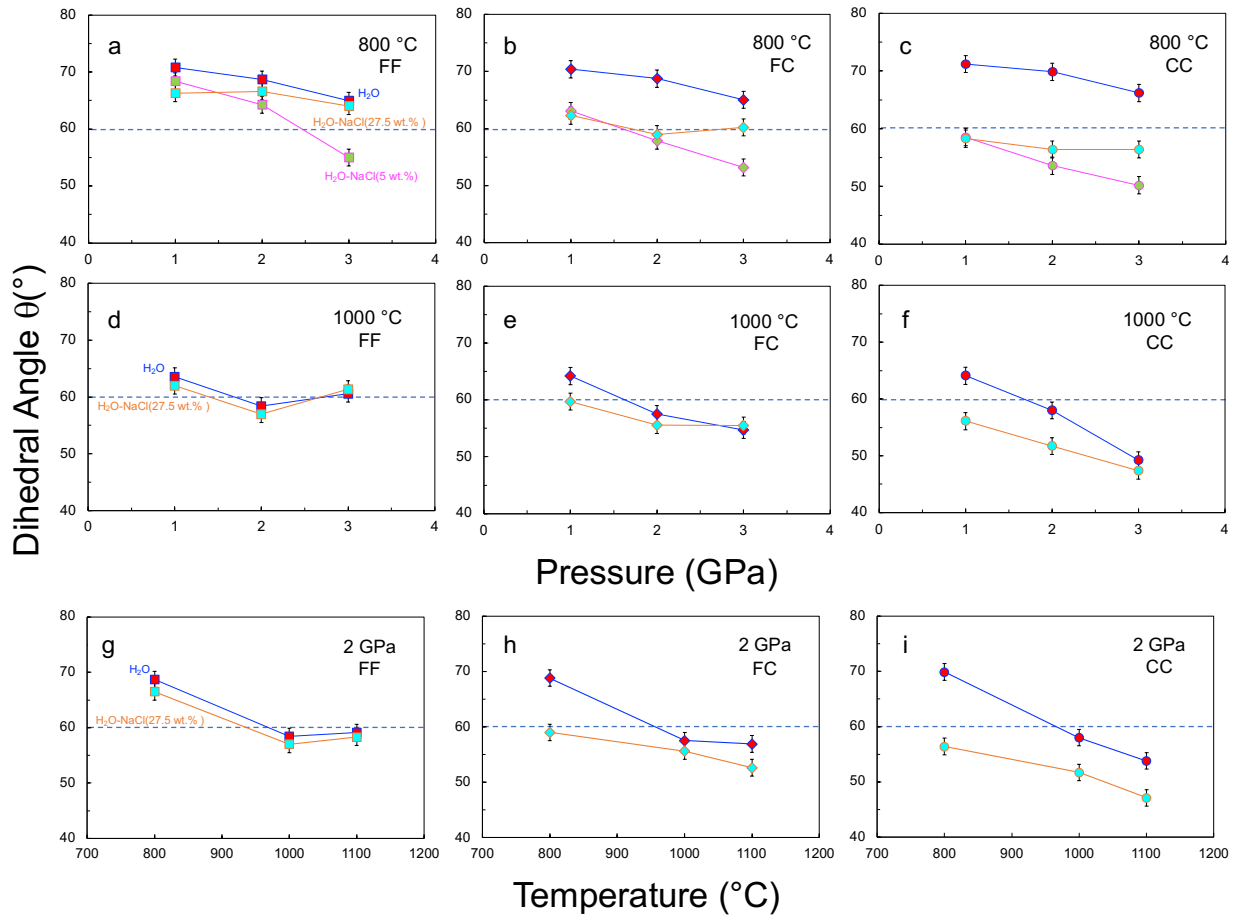
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Figure 5



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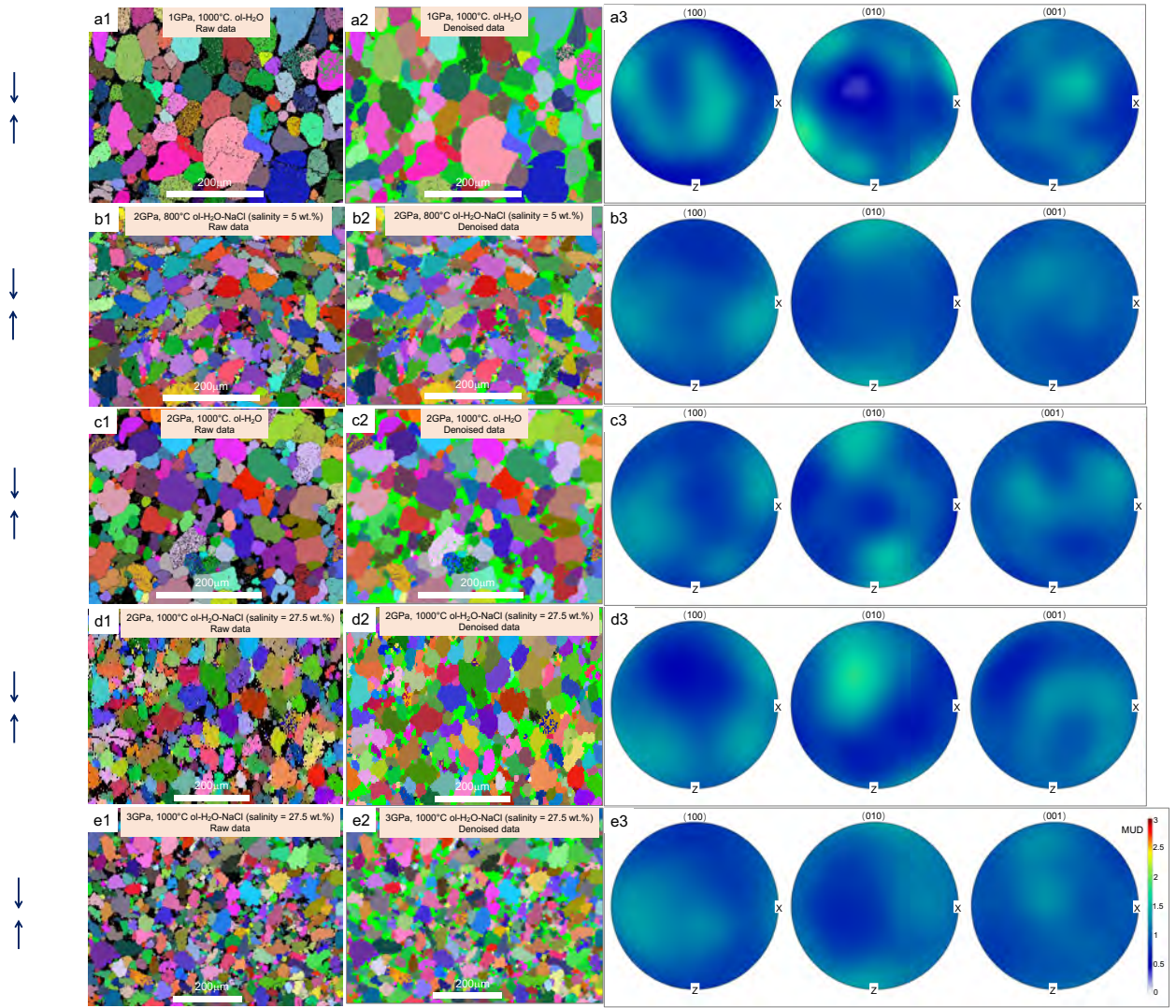
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Figure 6



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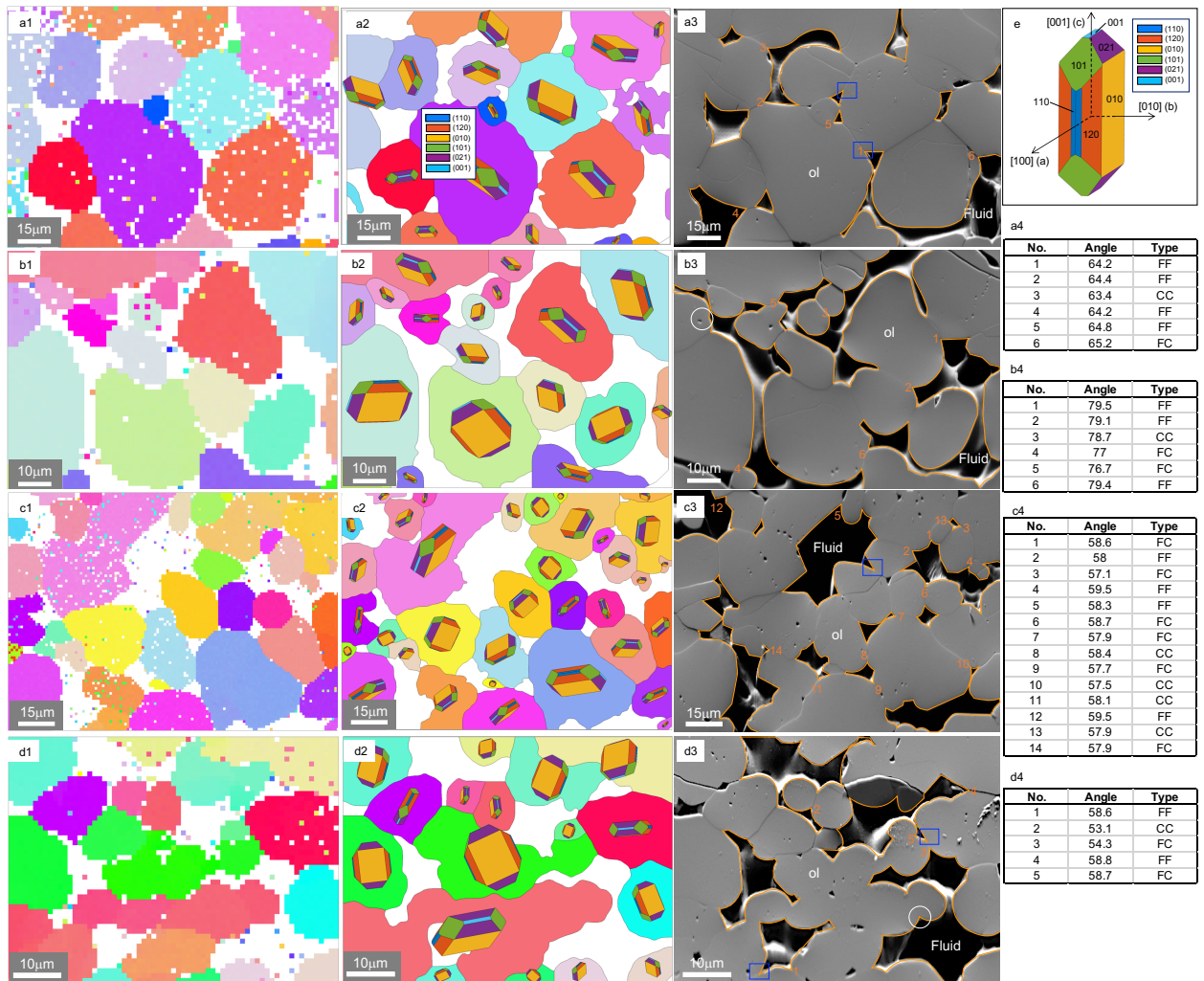
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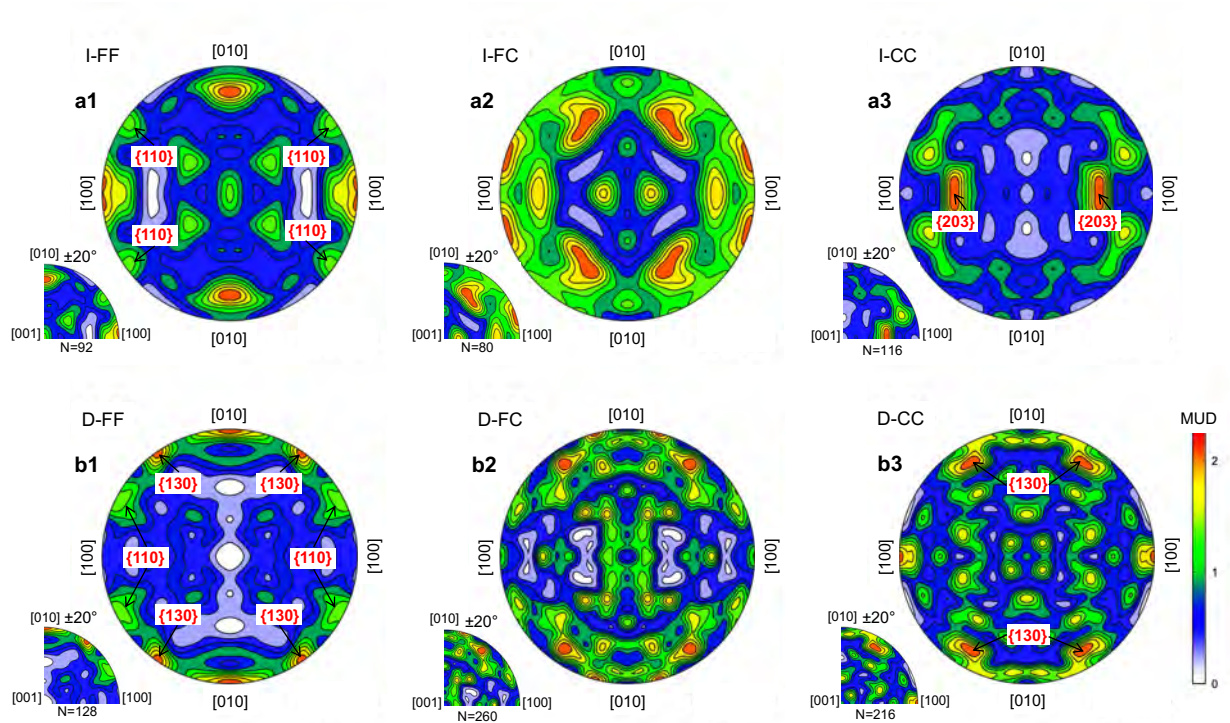
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Figure 7



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Figure 8



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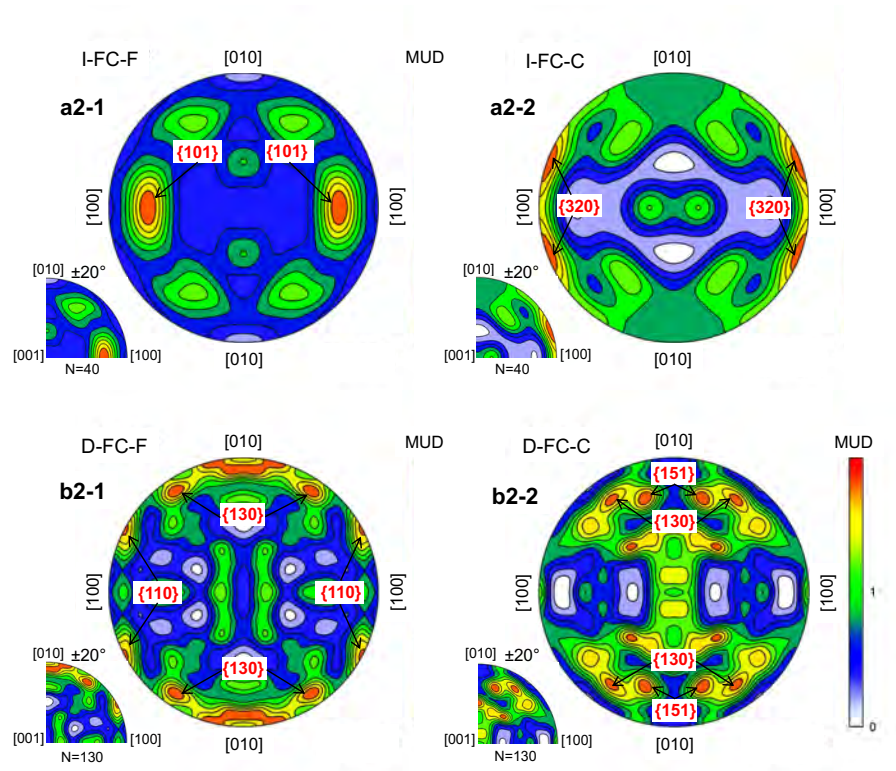
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Figure 9



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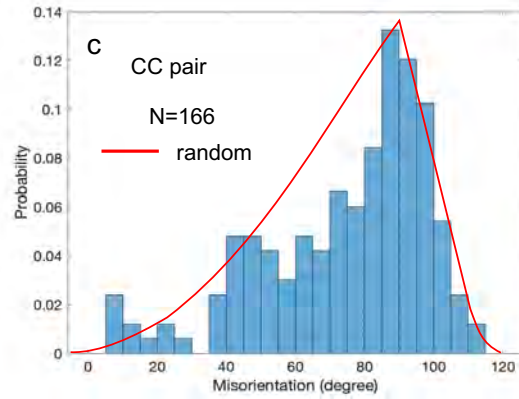
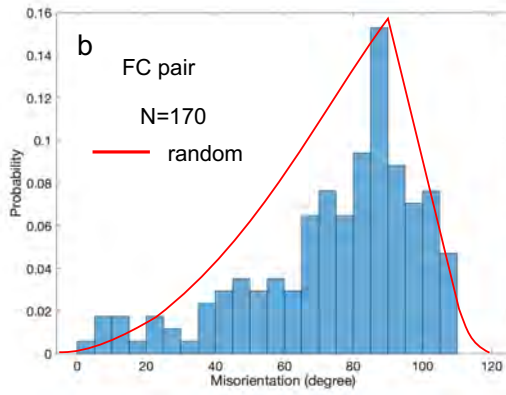
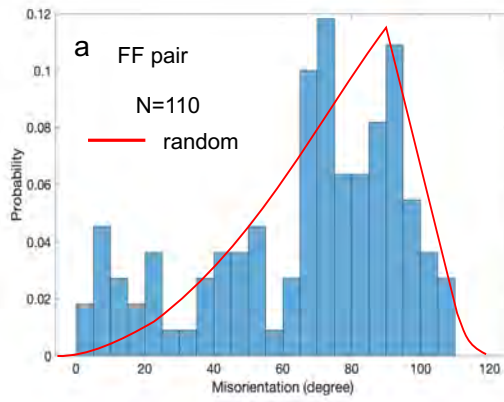
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Figure 10



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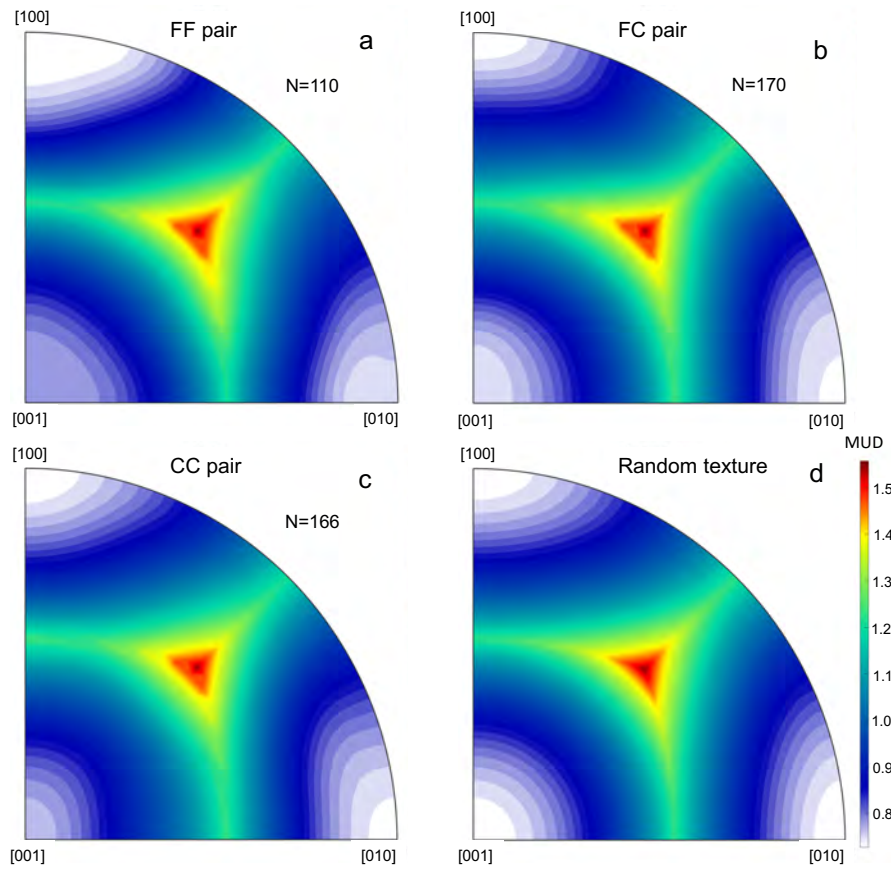
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Figure 11



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Figure 12

