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Revision 1

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Machine-learning oxybarometer developed using zircon trace-element

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chemistry and its applications

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Abstract

12

Magmatic oxygen fugacity (fO_2) is a fundamental property to understanding the

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long-term evolution of the Earth's atmosphere and the formation of magmatic–

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hydrothermal mineral deposits. Classically, the magmatic fO_2 is estimated using the

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mineral chemistry, such as Fe–Ti oxides, zircon, and hornblende. These methods,

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however, are only valid within certain limits and/or requires a significant amount of a

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priori knowledge. In this contribution, a new oxybarometer, constructed by

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data-driven machine learning algorithm using the trace elements of zircons and their

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corresponding independent fO_2 constraints, is provided. Seven different algorithms

20 are initially trained and then validated on a data set that was never utilized in the
21 training processes. Results suggest that the oxybarometer constructed by the
22 extremely randomized trees model has the best performance, with the largest R^2 value
23 (0.91 ± 0.01), smallest RMSE (0.45 ± 0.03), and low propagated analytical error
24 (~ 0.10 log units). Feature importance analysis demonstrates that U, Ti, Th, Ce, and Eu
25 in zircon are the key trace elements that preserved fO_2 information. This newly
26 developed oxybarometer has been applied in diverse systems, including the arc
27 magmas and mid ocean ridge basalts, the fertile and barren porphyry systems, and the
28 global S-type detrital zircon, which provide fO_2 constraints that are consistent with
29 other independent methods, suggesting that it has wide applicability. To improve
30 accessibility, the oxybarometer was developed into a software application aimed at
31 enabling more consistent and reliable fO_2 determinations in magmatic systems,
32 promoting further research.

33 **Keywords:** Machine learning; Zircon; Trace elements; Magmatic oxygen fugacity;
34 Oxybarometer.

35 **1. Introduction**

36 Oxygen fugacity (fO_2) is a fundamental thermodynamic property governing the
37 speciation and behavior of multivalent elements (e.g., S, Ce, Eu, Fe, and V) during
38 magma evolution, which in turn controls their solubility, mobility, and compatibility
39 in silicate magmas (Brounce et al. 2014; Ni et al. 2020). Estimates of fO_2 have been

40 used to help address a range of important questions, including the compositional
41 evolution of the atmosphere (Trail et al. 2011b; Lee and Bachmann 2014) and the
42 mineralization potential of igneous rocks (Jugo et al. 2010; Sillitoe 2010; Richards
43 2015). Traditionally, the oxidation state of ancient rocks is constrained using the
44 bulk-rock $\text{Fe}^{2+}/\text{Fe}^{3+}$ ratio (Kress and Carmichael 1991; Brounce et al. 2014; Zhang et
45 al. 2018), the Fe–Ti oxide oxybarometer (Ghiorso and Evans 2008), and the
46 hornblende oxybarometer (Ridolfi and Renzulli 2012; Ridolfi 2021). Previous studies,
47 however, have pointed out that these oxybarometers are only applicable in a limited
48 range of conditions. For example, the bulk $\text{Fe}^{2+}/\text{Fe}^{3+}$ ratio of a glass is easily reset by
49 the subsequent metamorphism or alteration (Trail et al. 2011b), the Fe–Ti oxide
50 oxybarometer is only applicable rapidly quenched volcanic rocks (Loucks et al. 2018),
51 and the hornblende oxybarometer is only suitable for rocks emplaced deep in the crust
52 because hornblende is unstable in the shallow crust (Rutherford and Hill 1993).
53 Therefore, a more broadly applicable and robust oxybarometer is needed to unravel
54 the complex interplay between $f\text{O}_2$, magmatic evolution, and metallogenesis through
55 geological time.

56 Zircon is an ubiquitous accessory mineral in crustal rocks and is geochemically
57 robustness, even in rocks that were hydrothermally altered, metamorphosed and
58 weathered (Cherniak and Watson 2003; Trail et al. 2011a). Zircon contains a variety
59 of trace elements, most of which include REEs, Ti, Y, U, Th, and Hf. These elements
60 mainly enter into the zircon crystal lattice via isomorphic and coupled substitution

61 mechanisms (Hoskin and Schaltegger 2003). The former includes $U^{4+} \leftrightarrow Zr^{4+}$, $Ti^{4+} \leftrightarrow$
62 Zr^{4+} , $Hf^{4+} \leftrightarrow Zr^{4+}$, and $Ce^{4+} \leftrightarrow Zr^{4+}$, and the latter includes $(REEs + Y)^{3+} + Na^+ / K^+ /$
63 $H^+ \leftrightarrow Zr^{4+}$ and $P^{5+}/Nb^{5+} + (REEs + Y)^{3+} \leftrightarrow 2Zr^{4+}$. Substitution of these elements into
64 zircon is related to the similarity of their ionic radius and charge to Zr^{4+} , which are
65 controlled by the physicochemical conditions (e.g., temperature, oxidation state, and
66 magma composition) of magma during mineral growth (Ballard et al. 2002; Ferry and
67 Watson 2007). Oxidation state mainly affects substitution efficiency by changing the
68 valence state of redox-sensitive elements. Consequently, the concentrations of the
69 multivalent REEs (i.e., Ce and Eu) relative to the other monovalent REEs in zircon
70 and the parent magma has become a popular proxy to characterize the fO_2 of magmas
71 (Ballard et al. 2002; Trail et al. 2011b; Smythe and Brenan 2016; Loader et al. 2017;
72 Zhong et al. 2019; Loucks et al. 2020). However, recent studies (Zhong et al. 2019;
73 Zou et al. 2019) have demonstrated that such zircon REE oxybarometers may be
74 unreliable given the difficulty of accurately determining the composition of the parent
75 magma and the existence of REE-mineral inclusions within zircon. Accordingly,
76 Loucks et al. (2020) proposed a new oxybarometer using ratios of Ce, U, and Ti in
77 zircon (the U–Ti–Ce equation), which does not require knowledge of the parent
78 magma composition. We have evaluated the reliability of the U–Ti–Ce equation of
79 Loucks et al. (2020) by applying it to the global detrital zircon database from Tang et
80 al. (2021) and reference therein. The results of this assessment demonstrate that some
81 detrital zircons crystallized from a magma with an extremely (almost impossibly)

82 reduced fO_2 (less than FMQ-5) (Fig. S1), indicating that the linearly regressed U–Ti–
83 Ce equation does not capture all of the redox information of magmas and is not
84 always valid. In addition, the U–Ti–Ce equation may not work well in the strongly
85 peraluminous or peralkaline felsic magmas because their differentiation index (i.e.,
86 U/Ti) is not consistent with the variation of Ce/U ratio in melts (Loucks et al. 2020).

87 Recently, several studies have demonstrated that the data-driven machine learning
88 methods can be powerful tools for solving complex problems in mineralogy, petrology,
89 and geochemistry (Petrelli and Perugini 2016; Chen et al. 2021; Huang et al. 2022;
90 Lin et al. 2022; Nathwani et al. 2022; Qin et al. 2022; Wang et al. 2022; Zou et al.
91 2022), and for the construction of thermobarometers (Petrelli et al. 2020; Higgins et al.
92 2022; Jorgenson et al. 2022; Li and Zhang 2022), without having any a priori
93 knowledge. These research advances suggest that the machine learning method has
94 the potential to be used for calibrating a mineral chemical-based oxybarometer. In this
95 contribution, we present a novel machine learning-based approach to develop an
96 oxybarometer using zircon trace-element chemistry and their corresponding fO_2
97 values obtained by other independent methods. Initially, seven different
98 oxybarometers were first constructed and trained using prevalent machine learning
99 algorithms, including linear regression, decision trees (Cramer et al. 1976), random
100 forest (Breiman 2001), extremely randomized trees (ERT, Geurts et al. 2006), support
101 vector machines (Smola and Schölkopf 2004), K nearest neighbors (Fukunaga and
102 Narendra 1975) and XGboost (Chen and Guestrin 2016), and their performance was

103 then benchmarked on a testing set to select an optimal algorithm. The calibrated
104 oxybarometer was then applied in three different geological situations to test its
105 performance and to explore the breadth of its applicability. To enhance accessibility,
106 we developed a user-friendly web app and intuitive software with a graphical
107 interface, which allow free online or offline fO_2 estimation from zircon chemistry
108 using the machine learning oxybarometer, respectively. This study contributes a new
109 tool to determine oxidation states in silicate melts without a priori knowledge. The
110 approach circumvents limitations of existing methods dependent on specific
111 equilibrium mineral pairs or intensive variables. With appropriate training data, it
112 potentially provides a widely applicable oxybarometer to address outstanding
113 questions on magma fO_2 .

114 **2. Data and methods**

115 ***2.1. Data compilation and filtration***

116 The data used to construct the machine learning models in this study were
117 collated from published articles, and comprise zircon trace-element chemistry and
118 independently constrained fO_2 values of their host rocks. The unfiltered calibration
119 data set is composed of 1,450 data points of zircons from volcanic and plutonic rocks
120 from 37 locations around the world; this data is summarized in Table S1 and the full
121 data set can be found in Table S2. For the calibration of a robust oxybarometer,
122 unreliable zircon analyses have been removed based on the following criteria. First,

123 analyses with Th/U ratios less than 0.1 have been excluded to avoid metamorphic
124 zircon (Kirkland et al. 2015). Second, analyses with La and P concentrations greater
125 than 1 and 2000 ppm, respectively, have been removed as these analyses were likely
126 contaminated by mineral inclusions (Zou et al. 2019; Zhu et al. 2020). Third,
127 elemental data with >30% of the values as missing or 0 were not included. These
128 criteria led to the selection of P, Ti, Y, Nb, Hf, Th, U, and other 14 REEs as the
129 elements of interest and a dataset consisting of 1,369 zircon compositions with fO_2
130 values ranging from FMQ-4.9 to FMQ+2.75.

131 ***2.2. Data treatment of null values***

132 Following the filtration, a log-transformation has been employed to the dataset,
133 with the pre- and post-transformation features of the collected data illustrated in
134 Figure S2. The purpose of this transformation is to ensure the distribution of elements
135 approximate normality (Fig. S2b), thereby enhancing the performance of machine
136 learning models, as disputed in previous studies (e.g., Frenzel et al. 2016; Petrelli et al.
137 2021). Regrettably, a significant amount (<30%) of the included elements contained
138 null values or zeros, i.e., invalid data, which is an inevitability and may derived from
139 different analytical procedures and instruments utilized to obtain the data.
140 Consequently, the natural logarithm computation of null or 0 values becomes
141 meaningless, producing misleading outcomes. To address this issue, various types of
142 invalid data can be imputed using appropriate multiple strategies before applying the
143 logarithmic transformation. In this study, the REEs are all analyzed in the collected

144 data set, and the invalid values in the REEs data means they are below the detection
145 limit; while invalid values in other elements indicate that they are not detected.
146 Accordingly, we addressed this matter by employing the method proposed by van den
147 Boogaart and Tolosana-Delgado (2013) to fill invalid data using normal distributions.
148 For the REEs, the normal distribution was constructed with the mean and standard
149 deviation parameters set to their respective detection limits, while for the other
150 elements, the normal distribution was established based on the mean and standard
151 deviation of not null data.

152 ***2.3. Dimensionality reduction***

153 The correlation between the concentration of trace elements in zircon, especially
154 the heavy REEs, is high (Fig. S3). If these highly correlated trace elements were to be
155 utilized in the machine learning models, their interaction between elements would be
156 masked and the noise information would be magnified, leading to overfitting and
157 spurious interpretations (Hall 1999; Lösing and Ebbing 2021). To solve the problem
158 that trace elements in zircon are highly correlated, the strategy of reducing data
159 dimension and filtering noise by Principal Component Analysis (PCA) first, and then
160 training the machine learning models was adopted in this study. The PCA method is a
161 statistical procedure that identifies patterns among variables according to their
162 correlations (Jackson 2005), and can map high-dimensional data to low-dimensional
163 data while retaining as much information as possible from the original dataset (Abdi
164 and Williams 2010). In addition, principal component loadings for each variable in the

165 compositional biplots can provide the contribution of the variable to the principal
166 component (Aitchison and Greenacre 2002), which can be used to discover the
167 representative element associations.

168 ***2.4 Machine learning model***

169 *2.4.1. Data labeling*

170 Supervised machine learning algorithms were adopted to calibrate the
171 oxybarometer. Supervised machine learning is an algorithms that is trained by using
172 labeled datasets to learn an appropriately function that maps the inputs (i.e., the
173 trace-element composition of zircon in this study) to the predicted outputs (i.e., the
174 fO_2 values in this study) using an the optimization algorithm (Chen et al. 2022). Once
175 the training processes are complete, the trained model can be used as an oxybarmeter.
176 Accordingly, the preparation of a dataset with calibrated labels is the key to
177 constructing supervised machine learning models. In our case, the trace-element
178 composition of zircon can be labeled using fO_2 values (expressed as ΔFMQ) obtained
179 from the independent fO_2 constraints. Considering that the independent fO_2
180 constraints have some errors (Table S1), in this study, the trace-element composition
181 of zircon from the same locality was labeled with random values from a normal
182 distribution, and its mean and standard deviation are equal to the average fO_2 value
183 and uncertainty in this area, respectively.

184 *2.4.2. Machine learning model construction workflow*

185 In this study, seven widely used supervised machine learning methods, including
186 linear regression, decision trees (Cramer et al. 1976), random forest (Breiman 2001),
187 extremely randomized trees (ERT, Geurts et al. 2006), support vector machines
188 (Smola and Schölkopf 2004), K nearest neighbors (Fukunaga and Narendra 1975),
189 and XGboost (Chen and Guestrin 2016), were used to calibrate the oxybarometer.
190 Data processing and machine learning model construction procedures were completed
191 in Python 3.10 utilizing the python packages, including Scikit-Learn
192 (<https://scikit-learn.org/stable/index.html>) and XGboost
193 (<https://xgboost.readthedocs.io/en/latest/>). A detailed description of the model
194 building workflow used in this study (modified after (Li and Zhang 2022) and (Zou et
195 al. 2021)) is illustrated in Fig. 1 and can be summarized in the following nine steps: (1)
196 Data was collected from the pre-reviewed literatures and filtered based on the
197 aforementioned criteria. (2) The collected data was preprocessed to meet the
198 requirements of machine learning, including imputation of null data,
199 log-transformation, data dimensionality reduction, and data labeling. (3) The filtered
200 dataset was divided into a training (80% of the data) and testing (20% of the data) by
201 stratified random splitting. The training set was used to train the model and optimize
202 the hyperparameters, while the testing set was only used to evaluate the model. (4) A
203 grid searching algorithm with 5-fold cross validation (CV) was used to select the
204 optimal hyperparameters combination based on the performance metrics (i.e.,

205 coefficient of determination [R^2] score and root-mean-square error [RMSE]). Detailed
206 regarding the tuning of hyperparameters can be found in Zou et al. (2022). (5) With
207 the optimal hyperparameters combination, the training set was used to retrain the
208 machine learning model, the performance of which was then evaluated using the
209 testing set. (6) To avoid overfitting, the data preprocessing, random splitting, training,
210 hyperparameters tuning, and model evaluation (i.e., steps 2, 3, 4 and 5) were repeated
211 1000 times, and 1000 paired R^2 scores and RMSE values were obtained. (7) By
212 comparing the mean R^2 scores and RMSE values generated by different machine
213 learning models, the best model (i.e., the one with the largest R^2 scores and the
214 smallest RMSE values) was chosen as the target model. (8) The target model was then
215 tuned using the entire filtered dataset using a grid searching algorithm with 5-fold CV
216 to figure out the optimal hyperparameters combination. (9) The final, calibrated
217 model was then trained using the entire dataset and the best hyperparameters
218 combination obtained in step (8).

219 **3. Results**

220 ***3.1. Principal component analysis***

221 PCA was performed on the entire filtered dataset to determine its dominant
222 geochemical features. As illustrated in the compositional biplots, which combines the
223 datapoints and PCA loadings (Figs. 2a and 2b), apart from a few abnormal data from
224 the Lunar Highlands and kimberlites in southern Africa, most of the data are clustered

225 together, and the 21 elements can be roughly be divided into three groups by the PC
226 loadings — group 1 comprises U–Th–Ce, group 2 comprises of Ti, and group 3
227 comprises the REEs, with two subgroups for the light REEs and heavy REEs (Figs. 2a
228 and 2b). The PCA loadings also confirm the strong correlation between the REEs
229 (Figs. 2a and 2b). All of the elements, except for Ti, have negative loadings for the
230 PC1. PC2 is mainly expressed by negative loadings for U, Th, and Ce and a positive
231 loading for Ti. PC3 is dominated by a positive loading of La. The loadings for the
232 other PCs can be found in the Figure S4. In the scree plot, the first seven PCs (from
233 PC1 to PC7) cumulatively explain more than 90% of the total variance of the dataset
234 (Fig. 2c), and the last nine PCs (from PC13 to PC21) cumulatively explain less than 1%
235 of the variance (Fig. 2c). This means that PC1–PC7 preserve almost all the
236 information of the original dataset. Accordingly, the noise in the entire dataset can be
237 filtered out by reducing the dataset from its original 21 dimensions to 7 dimensions
238 via the PCA method; we use these 7 dimensions to train our machine learning-based
239 oxybarometer.

240 ***3.2. Comparison the performance of the different machine learning models***

241 Previous studies have revealed that different machine learning algorithms exhibit
242 unique strengths and weaknesses, with the most suitable model for a given task
243 contingent on the data characteristics and intended outcomes (e.g., [Petrelli et al. 2020](#);
244 [Li and Zhang 2022](#)). As a result, it is customary to leverage a range of diverse
245 machine learning algorithms for specific tasks, followed by an assessment of each

246 model's performance (quantifying the discrepancy between predicted and observed
247 values) to identify the optimal algorithm. Typically, R^2 and RMSE values between
248 predicted and observed values serve as evaluation metrics. A higher R^2 value implies a
249 better model fit to true values, whereas a smaller RMSE value typically indicates
250 reduced deviation between predicted and observed values. Consequently, both metrics
251 are frequently employed to assess the performance of machine learning models. In
252 this study, we apply this approach, where we computed the mean R^2 and RMSE
253 values for 1000 repeated calculations using the identical dataset to assess and compare
254 the performance of several machine learning models used to estimate fO_2 . Figure 3
255 illustrates the probability density distributions of the R^2 and RMSE values for fO_2
256 estimations performed on the testing dataset.

257 Our results demonstrate that the performance of different machine learning
258 models varies significantly, with R^2 values ranging from 0.3 to 0.9 and RMSE values
259 ranging from 0.3 to 1.25. Among all the models, the ERT algorithm exhibits the best
260 performance, with the highest R^2 value of 0.91 ± 0.01 and the lowest RMSE value of
261 0.45 ± 0.03 . In addition, the ERT model performs better than the previously proposed
262 equation by Loucks et al. (2020), with higher R^2 and lower RMSE value during the
263 1000 times repeated calculations.

264 The ERT algorithm is a powerful machine learning technique that constructs an
265 ensemble of decision trees and aggregates their predictions to make a final prediction.
266 Previous studies have demonstrated that the ERT algorithm performs well on

267 high-dimensional datasets with noisy features, due to its ability to reduce the impact
268 of noisy features and avoid overfitting (Petrelli et al. 2020). Therefore, an ERT-based
269 model was chosen to construct the final, calibrated machine learning-based
270 oxybarometer in this study.

271 **3.3. Estimation of error**

272 The accuracy, stability, and uncertainty of the calibrated machine learning model
273 must be assessed before it can be used as a reliable oxybarometer. During training and
274 evaluation of the models, the probability density distributions of the R^2 and RMSE
275 values of the extremely randomized trees model had the highest peaks and most
276 constrained ranges, indicating that the model has limited variance and high stability
277 (Fig. 3). The cross-validation error of the model can be expressed as RMSE of $0.45 \pm$
278 0.03 .

279 The propagated error of the analysis uncertainties should also be evaluated; this
280 can be estimated using the residual values (i.e., the difference between the measured
281 fO_2 values and the predicted values from our model, (Jorgenson et al. 2022)). To avoid
282 self-validation and overfitting, the uncertainties should be evaluated using data that
283 was never used in the training process. Considering that the final constructed model
284 was trained using the entire dataset, we performed a bootstrap resampling ($n = 1000$)
285 Monte Carlo calculation (Keller and Schoene 2012) using zircon data with analytical
286 errors equivalent to that of sample HB-18 ($\Delta FMQ = + 1.24$) from (Meng et al. 2021)

287 to generate new data that can maximally represent HB-18. As shown in Table S3,
288 none of the newly generated zircon compositions appear in the training dataset. Thus,
289 we used these new data to predict their fO_2 values using the calibrated machine
290 learning-based oxybarometer; the predicted results and residuals are illustrated in
291 Figure 4. The results demonstrate that the predicted mean fO_2 value ($\Delta FMQ = +1.39 \pm$
292 0.27) is slightly higher than the measured value ($\Delta FMQ = +1.24$), but consistent with
293 the measured value within the 1 sigma standard error estimate (SEE). As shown in
294 Figure 4a, about 70% of predictions fall within the ± 1 SEE range and the median
295 residual of the 1000 values is 0.10 (Fig. 4b), suggesting that the analytical propagated
296 error is about 0.10 log units, which is negligible relative to the model error ($0.45 \pm$
297 0.03). For comparison, the equation from Loucks et al. (2020) was also used to predict
298 the fO_2 values of the new data. Figure 4c and 4d demonstrates that this oxybarometer
299 gives a predicted mean fO_2 value of 1.86 ± 0.72 , with an analytical propagated error
300 of 0.46 log units. This comparison reveals that the machine learning-based
301 oxybarometer performs better under these conditions.

302 **4. Discussion**

303 ***4.1. Mapping relationships between trace elements in zircon and oxygen fugacity***

304 In order to understand the relationship between zircon trace elements and their
305 oxygen fugacity in machine learning model, it is imperative to discern the pivotal
306 input data that governs the model's predictions. The Shapley Additive exPlanations

307 (SHAP), a game theory-based approach that indicates the output of machine learning
308 models, is a widely recognized technique. It facilitates the interpretation of results by
309 recognizing the contribution of each feature to the model's output. By utilizing the
310 SHAP method, we can gain a better comprehension of how the model arrived at its
311 predictions and identify the features (i.e., elements in this study) that significantly
312 influence the model's prediction-making.

313 In this study, the SHAP method was used to help explain the relative importance
314 of each PC to the output of the model; this was completed using the python package
315 of SHAP (<https://shap.readthedocs.io/en/latest/index.html>). This method involves
316 calculating the SHAP value for each feature (i.e., the PC in this study) of the sample,
317 which represents the influence of that feature on the prediction (Lundberg and Lee
318 2017). A higher average absolute-size SHAP values for a feature indicates a higher
319 influence on the prediction and vice versa. In addition, positive or negative SHAP
320 values of each datapoint represents whether its output is positive or negative. Figure 5
321 shows the relative importance of features (i.e., the 7-dimentional data) are ordered
322 based on their mean absolute-size SHAP value. As shown in Figure 5a, relative
323 importance scores, from high to low, are PC2, PC4, PC3, PC5, PC7, PC6, and PC1.
324 Figure 5b shows the SHAP value of features for all individual analysis and also shows
325 that how the high and low SHAP values of the features (i.e., PC1 to PC7) impact the
326 model output. As shown in Figure 5b, the high values of PC2 and PC4 have the
327 greatest contribution for the model to obtain negative values, whereas the low values

328 of PC2 and PC4, and high values of PC5 have the greatest contribution to obtain
329 positive values. These results reveal that PC2, PC4, and high values of PC5
330 significantly impact the model output. It is worth noting that PC2 always contributes
331 more than the other PCs on the model output.

332 The features (i.e., PC1 to PC7) fed into the machine learning model are derived
333 from the first seven principal components that cumulatively explain 90% of the
334 variance of the original dataset. The PCA-derived loadings of elements can help
335 characterize the relationships between the PC values and the elemental concentrations.
336 Figure 2a and 2b illustrate that the low PC2 values are related to U–Th–Ce and its
337 high values are related to Ti. Similarly, low PC4 values are mainly related to Eu and
338 its high values are related to La (Fig. S4). High values of PC3 and PC5 are mainly
339 derived from La and Eu, respectively (Fig. S4). Given that the La content of zircon is
340 typically very low to the point that it is typically hard to measure accurately (Zhong et
341 al. 2019; Zou et al. 2019) and the fact that La is missing in nearly 20% of the original
342 dataset, the importance of La is likely being magnified, but the impact is still limited
343 relative to other elements (i.e., U, Ti, Th, Ce, and Eu) used in the model because PC2
344 always contributes more than the other PCs. Accordingly, we can conclude that the U,
345 Ti, Th, Ce, and Eu are the most important elements for the machine learning model
346 and preserved fO_2 information.

347 Loucks et al. (2020) discussed, in detail, the relationships between the U, Ti, and
348 Ce contents of zircons and the fO_2 conditions of their formation; this relationship was

349 described using the equation of $\log fO_2(\Delta FMQ) = 3.998(\pm 0.124) \times \log$
350 $[Ce/\sqrt{(U_{\text{initial}} \times Ti)}] + 2.284 (\pm 0.101)$ (i.e., the U–Ti–Ce equation). Along with the
351 elements used in this equation, our machine learning model demonstrates that Th and
352 Eu also preserve fO_2 information and are important to predicting fO_2 of a magma.
353 This may be one of the reasons for the comparatively larger error of the fO_2 estimates
354 using the above equation. Previous studies have demonstrated that the compatibility
355 Th and Eu entering into zircon is controlled by oxygen fugacity (Burnham and Berry
356 2012). Because uranium is a redox-sensitive element with two major valence states
357 (U^{4+} and U^{6+}), with U^{4+} being more compatible into zircon substitution of Zr^{4+} .
358 Thorium in the tetravalent state behaves geochemically similar to U^{4+} because of their
359 similar ionic radius, which may cause Th^{4+} and U^{4+} to compete for the Zr^{4+} position in
360 zircon. Therefore, the Th/U ratio of zircon may be controlled by the environmental
361 oxidation state of the magma. Similarly, Eu is also a redox-sensitive element and can
362 exist as Eu^{3+} and Eu^{2+} . Eu^{3+} is more compatible into zircon than Eu^{2+} , and so a Eu
363 anomaly in zircon can be used to estimate prevailing oxygen fugacity during its
364 formation (Burnham and Berry 2012; Loader et al. 2017). Taken together, we can
365 conclude that our constructed machine learning model correctly captures the
366 relationship between trace elements in zircon and oxygen fugacity.

367 ***4.2. Applications of the machine learning-based oxybarometer to natural systems***

368 To further explore the range of applicability of this machine learning-based
369 oxybarometer, we have applied it to three different geological scenarios in which the

370 fO_2 has been well constrained by previous studies — i) arc magmas and mid ocean
371 ridge basalts (MORB), ii) fertile and barren porphyry systems, and iii) a global
372 database of S-type detrital zircon.

373 *4.2.1. Estimation of fO_2 of arc magmas and MORB*

374 Arcs and the mid ocean ridge are the two most important tectonic settings for the
375 production of magmas globally, the former of which producing more oxidized
376 magmas, on average, than the latter (Christie et al. 1986; Carmichael 1991; Brounce
377 et al. 2014; Wang et al. 2019). Evidence from whole-rock Fe^{3+}/Fe^{2+} , V/Sc, and Zn/ Fe_T
378 ratios (Carmichael 1991; Lee et al. 2010; Brounce et al. 2014), Eu anomalies in zircon
379 (Burnham and Berry 2012), and the composition of spinel and olivine (Evans et al.
380 2012; Wang et al. 2019) indicate that the measured fO_2 value at different times in most
381 arc magmas vary from $\Delta FMQ = +0.5$ to $+2$ (locally up to $+3$), while MORBs have fO_2
382 values in the range of $\Delta FMQ = -1$ to $+0.5$ (Jugo et al. 2010; Evans et al. 2012;
383 Richards 2015; Wang et al. 2019).

384 In this study, the composition of zircons from arcs and MORBs were used to
385 validate the machine learning-based oxybarometer. Data for arcs (the Alenutian Arc,
386 Andean Arc, and Central American Arc) were collated from the GEOROC database
387 (<https://georoc.eu/georoc/new-start.asp>, accessed on October, 2022), and data for the
388 Vema MORB are collected from (De Hoog et al. 2014); these data were filtered using
389 criteria described above and are presented in Table S4. As illustrated in Fig. 6, the
390 predicted mean fO_2 values of arcs are consistent with each other and are greater than

391 that of the Vema MORB by approximately 1 log unit. In addition, the average fO_2
392 values for arcs range from 1.05 to 1.17 log units above the FMQ buffer (FMQ+1.05 to
393 FMQ+1.17), whereas the average fO_2 value of the Vema MORB is similar to the FMQ
394 buffer (FMQ-0.01). These results are in excellent agreement with the fO_2 constraints
395 from previous works (Ballhaus 1993; Zhang et al. 2018). It is crucial to recognize that
396 the most fO_2 proxies from previous studies on arc magmas and MORB most likely
397 represent various stages of magma system evolution. Zircons are particularly effective
398 at capturing fO_2 information from the more evolved intermediate to felsic magma
399 products, whereas other proxies may be more sensitive to earlier, more mafic
400 conditions. From this point, our machine learning-based oxybarometer, when
401 combined with other proxies, has the potential to offer valuable insights into the
402 changes in redox conditions experienced by arc magmas and MORB during their
403 evolution.

404 *4.2.2. Fertile and barren porphyry systems*

405 Porphyry deposits are thought to form as a result of fluid exsolution derived from
406 hydrous, volatile-rich and oxidized magmas (Sillitoe 2010; Richards 2015). It has
407 been well established that, in oxidized magmas, S occurs as sulfate; in such
408 environments, the sulfur content at sulfur saturation is higher, leading to the presence of
409 S undersaturated magmas that can interact with wall-rocks to form porphyry deposits
410 (Streck and Dilles 1998; Jugo et al. 2010; Sillitoe 2010; Richards 2015; Shen et al.
411 2015; Sun et al. 2015). Therefore, the redox state of an intrusion is considered to have

412 a significant impact on the formation of porphyry deposits, and has been used as one
413 of the most important indicators to evaluate mineralization potential (Shen et al. 2015;
414 Lu et al. 2016). Previous studies have demonstrated that the fertile rocks generally
415 crystallized from highly oxidized ($\Delta\text{FMQ} > + 1$) and hydrous magmas, whereas
416 barren rocks generally crystallize from the relatively reduced ($\Delta\text{FMQ} < + 1$) and dry
417 melts (Richards 2015; Rezeau and Jagoutz 2020). Given the apparent link between
418 $f\text{O}_2$ and mineralization potential, we also applied our machine learning-based
419 oxybarometer to a large database of fertile and barren rocks from (Zou et al. 2022).
420 Our calculation revealed the average $f\text{O}_2$ value of fertile rocks in this database is
421 approximately 0.4 log units higher than that of barren rocks (FMQ + 1.40 vs. FMQ
422 +0.62), with about 30% of data overlapping each other (Fig. 7). The overlap between
423 the two rock groups may be attributed to the fact that the formation of porphyry
424 deposits is not independently dictated by oxygen fugacity, but is also related to the
425 water and volatile contents of the magma. It is, therefore, difficult to distinguish
426 between fertile and barren rocks based solely on their $f\text{O}_2$ (Zou et al. 2022). Although
427 overlap does exist, these results are consistent with others that fertile rocks generally
428 form from higher $f\text{O}_2$ magmas than barren rocks.

429 *4.2.3. Global S-type detrital zircons*

430 The rise of atmospheric oxygen has had a profound impact on the chemical
431 environment of the Earth's surface, ultimately leading to the evolution of a diverse
432 biosphere (Kaufman et al. 2007; Lyons et al. 2014). It is widely agreed upon that two

433 events caused the oxygen content of the atmosphere to rise to current levels: the Great
434 Oxidation Event (GOE) that occurred approximately 2.5–2.2 billion years ago, and
435 the later Neoproterozoic Oxidation Event (NOE) that occurred around 800–540 Myr
436 ago (Bekker et al. 2004; Frei et al. 2009; Crowe et al. 2013; Planavsky et al. 2014; Liu
437 et al. 2019; Chen et al. 2022; Hodgskiss and Sperling 2022). This established
438 knowledge provides a basis for testing our oxybarometer. In this study, the detrital
439 zircon derived from S-type granite (referred to as “S-type zircon”) is employed to
440 reconstruct the history of atmospheric oxygenation. S-type granites crystallize from
441 strongly peraluminous magmas derived (dominantly) from metasedimentary rocks
442 (Chappell and White 1992), which are derived from sedimentary rocks through
443 several steps, including metamorphism, partial melting, and subsequent crystallization.
444 As sediments provide a record of Earth surface processes, including variations in
445 atmospheric oxygen levels, and are the precursors to S-type granites, it is reasonable
446 to assume that the latter would also retain a record of atmospheric oxygen changes.

447 Trace element concentrations in zircon have been employed to distinguish S type
448 zircon from others. One method for distinguishing S-type zircon involves examining
449 its phosphorus concentration (e.g., [Burnham and Berry 2017](#); [Zhu et al. 2020](#)). This
450 method assumes that the higher solubility of apatite in peraluminous S-type granite
451 magma results a relatively high phosphorus content in S-type zircon, assuming
452 consistent zircon-melt P partition coefficients across diverse melts. However, recent
453 study has shown that the reliability of this approach is limited to zircon samples

454 younger than 720 Ma (Bucholz 2022; Bucholz et al. 2022), as the phosphorus content
455 in S-type granite before this age does not consistently exceed that of other granites
456 (Bucholz et al. 2022). As an alternative method, a machine learning approach utilizing
457 trace element data from known zircon type has been proposed by Zhong et al. (2023).
458 This method considers the influence of trace elements and has demonstrated high
459 identification accuracy (about 0.85). It is important to note, however, that this
460 approach did not incorporate ancient zircon samples predating 800 Ma in the model
461 training process and, as a result, may not be suitable for analyzing zircon from these
462 ages.

463 Therefore, this study aims to investigate the NOE through the fO_2 value of the
464 global S-type detrital zircon with an age younger than 720 Ma. To achieve this goal,
465 the trace-element composition of detrital zircons globally were collected from Tang et
466 al. (2021) and references therein, and S-type zircons (< 720 Ma) were identified based
467 on the criteria of Zhu et al. (2020): S-type zircons have a the molar concentration of
468 REE +Y greater than $0.77 * P$ and less than $1.23 * P$, and the molar concentration of P
469 is greater than $15 \mu\text{mol/g}$. After data filtering, a total of 116 data points of S-type
470 detrital zircon were recognized, as shown in Table S5, and their fO_2 of formation was
471 estimated using our machine learning-based oxybarometer. The predicted fO_2 values
472 of S-type zircon is displayed in Figure 8, with a curve fit in every 100 Myr interval.
473 Our findings are consistent with prior studies on sediments (e.g., [Partin et al. 2013](#);
474 [Reinhard and Planavsky 2022](#)), verifying a gradual increase in the fO_2 value of S-type

475 detrital zircon from 800 Ma to 200 Ma (Fig. 8). Notably, our analysis reveals a
476 distinctive shift in fO_2 levels of S-type zircons during the late Neoproterozoic period
477 (650 Ma) (Canfield et al. 2007), coinciding with the timing of the NOE; and the sharp
478 increase from 550–450 Ma, which is consistent with independent evidence indicating
479 comprehensive ocean oxidation during that time (Fig.8b-c). Accordingly, although
480 S-type zircon collected in this study do not provide a complete and continuous record
481 of fO_2 fluctuations (Fig. 8), our study identifies key oxidation events. This highlights
482 the effectiveness of the machine learning-based oxybarometer we developed, even in
483 in strongly peraluminous environments.

484 ***4.3. Implications and possible uncertainty***

485 Compared to the traditional oxybarometer, the data-driven machine learning
486 oxybarometer in this study is constructed based on the contents of P, Ti, Y, Nb, Hf, Th,
487 U and other 14 REEs in zircon. The greater number of elements used in the
488 calibration process means that less meaningful information is lost, leading to higher
489 precision and accuracy. More importantly, the successful application of the
490 oxybarometer to the three different geologic systems (i.e., MORB–arc, porphyry
491 systems, detrital zircons) suggests that it has the potential to be more widely
492 applicable than traditional equations. The calibration approach does not require any a
493 priori knowledge of physicochemical information and magma composition, which
494 makes it a valuable tool for researchers studying a wide range of environments.
495 However, further investigations with extensive datasets and analysis are needed to

496 confirm the effectiveness of this approach for more different geological environments.
497 For example, researchers could refine calibration procedures or develop new methods
498 for predicting oxygen fugacity by collecting additional data (especially data with low
499 fO_2 values that were not included in this study) and improving the existing model. The
500 continuous development of this model could help researchers to understand the
501 changes of fO_2 more accurately in diverse magma systems, which is essential for
502 improving our understanding of crustal evolution.

503 The uncertainties of the machine learning model are significantly dependent on
504 the features used. In this study, the input features are trace elements in zircon and the
505 output feature is their corresponding fO_2 values. The bootstrap resampling Monte
506 Carlo calculation shows that uncertainties from the analytical error of the input
507 features is low (~ 0.01 log unit), which can largely be ignored compared to the model
508 error. In contrast, the uncertainty of the output feature is more significant because of
509 its relatively narrow range of possible values (between FMQ-4.9 and FMQ+3.0). The
510 uncertainty in the output feature mainly comes from the fact that data with fO_2 values
511 outside of this range have not been published so far, which leads to those data not
512 being learned by the computer. Nevertheless, considering that the fO_2 of most natural
513 magma systems is between FMQ-4.9 and FMQ+3 (Loucks et al. 2020), our model
514 can still solve most problems. In addition, as more published data becomes available,
515 the machine learning model will become increasingly more intelligent, and its
516 performance will continuously improve. Lastly, when inputting the trace-element data

517 into the model, we recommend using the average composition of several analyses on
518 zircon from the same sample to avoid errors from spurious analyses.

519 ***4.4. Online web application and offline software***

520 To make the oxybarometer more accessible and user friendly, a browser-based
521 application and a GUI software (named ZirconfO2) have been developed, which can
522 be used online (<https://shaohaozou-fo2-webapp-7xqvo0.streamlit.app/>) or offline (on
523 Windows), respectively. These applications have been published on the GitHub
524 platform (<https://github.com/shaohaozou/fo2>) and will be updated as additional
525 geochemical data of zircon becomes publicly available.

526 The web-based oxybarometer and the GUI software are easy to use and do not
527 need any programming knowledge. It should be noted that accessing the web-based
528 oxybarometer requires a Streamlit account. Following the introduction provided on
529 the webpage or within the software, users can upload their data in the same format as
530 the template file (this can be download from the website) and click the “Calculate”
531 and “Download your results” buttons in turn to complete the calculation and get the
532 results.

533 **5. Conclusions**

534 We calibrated a new oxybarometer based on data-driven machine learning algorithms
535 using the trace-element contents of zircon and their corresponding fO_2 constraints as
536 model inputs. Although the developed oxybarometer does not provide a specific

537 mathematical equation to illustrate the relationship between the trace-element content
538 of zircon and oxygen fugacity, a reliable machine learning-based relationship has been
539 recognized. In addition, this oxybarometer does not rely on any assumptions, and
540 shows a higher performance and lower error than traditional oxybarometer equations.
541 Feature importance analysis indicates that the machine learning model can well
542 identify fO_2 information in zircon and make fO_2 predictions. The successful
543 application of this oxybarometer to three different geologic situations demonstrates
544 that it has great potential to be widely applicable to geoscience. This work
545 demonstrates that machine learning is a promising tool that can be applied to
546 investigate other regression questions in the Earth sciences.

547 **Data Availability Statement**

548 The datasets, code and software for the machine learning models developed in this
549 study are available at <http://doi.org/10.5281/zenodo.7578390> and can be also found at
550 <https://github.com/shaohaozou/fO2>, which will be updated as more geochemical data
551 of zircon becomes publicly available. Further descriptions regarding the construction
552 of the machine learning models are available through the scikit-learn documentation
553 (<https://scikit-learn.org/stable/index.html>).

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815

816 **Figure captions**

817 **Figure 1.** Schematic diagram illustrating the data pre-processing and construction
818 workflow used for the development of machine learning models.

819 **Figure 2.** Compositional biplots for a) PC1–PC2 and b) PC2–PC3 showing the zircon
820 data points and principle component loadings. PC loadings for each element are
821 plotted as orange lines. The numbers in brackets in the axis titles denote the variance
822 accounted for by the PCs. (c) Scree plot for zircon trace-element chemistry combined
823 with the cumulative explanation of the total variance of PCs.

824 **Figure 3.** Probability density distribution of (a) R^2 and (b) RMSE values for
825 predictions made on the test dataset by the machine learning algorithms. The values
826 are based on 1000 Monte Carlo simulations.

827 **Figure 4.** Results of the propagated error calculated based on the trace-element
828 contents of zircon obtained via 1000 Monte Carlo simulations. (a) The fO_2 values
829 predicted using the machine learning-based oxybarometer, and (b) the probability
830 density distribution and kernel density estimation for residuals (from machine
831 learning-based oxybarometer) between the predicted and the measured fO_2 values. (c)
832 The fO_2 values calculated using the U–Ti–Ce equation, and (d) the probability density

833 distribution and kernel density estimation for residuals (from U–Ti–Ce equation)
834 between the calculated and measured fO_2 values.

835 **Figure 5.** (a) Feature importance bar plots generated by the extremely randomized
836 trees algorithm, which shows the relative importance of different elements when
837 predicting the fO_2 value of a magma. (b) SHAP summary plots showing the influence
838 of individual samples on the prediction. The SHAP value of the individual elements
839 are calculated across all samples, which are used to explore the impact of each
840 element on predicting the fO_2 values. The color of the data point (from red to blue)
841 illustrates the influence of a feature on the output of the prediction value (from high to
842 low).

843 **Figure 6.** Box-whisker plots illustrating the range of fO_2 values of arcs and MORBs
844 predicted by the machine learning-based oxybarometer. The original data are collated
845 from the GEOROC database and the utilized data (after filtering) are provided in
846 Table S4.

847 **Figure 7.** Box-whisker plots illustrating the range of fO_2 values of fertile and barren
848 rocks from porphyry systems using the machine learning-based oxybarometer.

849 **Figure 8.** (a) The fO_2 values are predicted using our machine learning-based
850 oxybarometer and zircons from S-type granites. (b) Sediments U contents in black
851 shales (Scott et al. 2008). (c) The blue line shows the best estimate of atmospheric

852 oxygen (pO_2 (% PAL)) best on the Ce anomaly in marine carbonates (in red line),
853 which are from Liu et al. (2021).

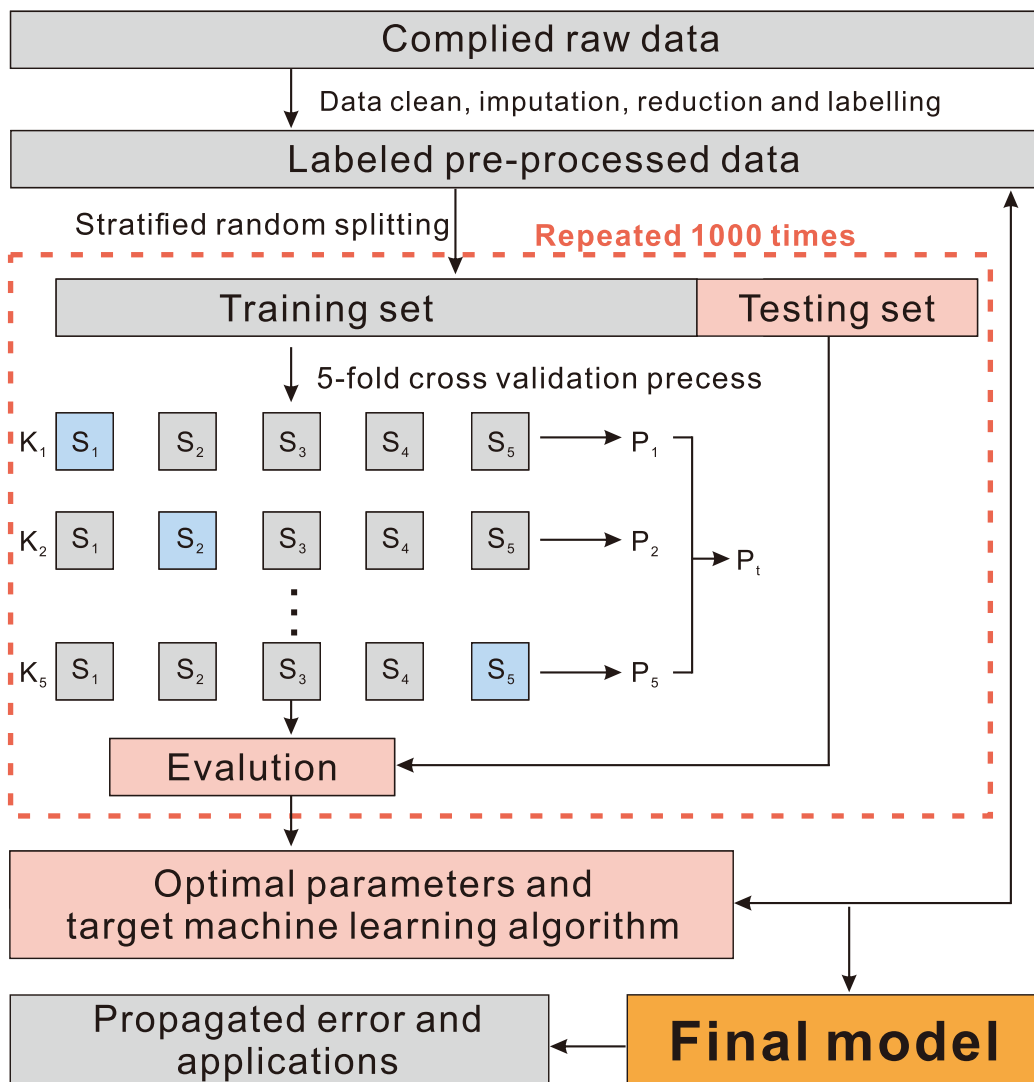


Figure 1

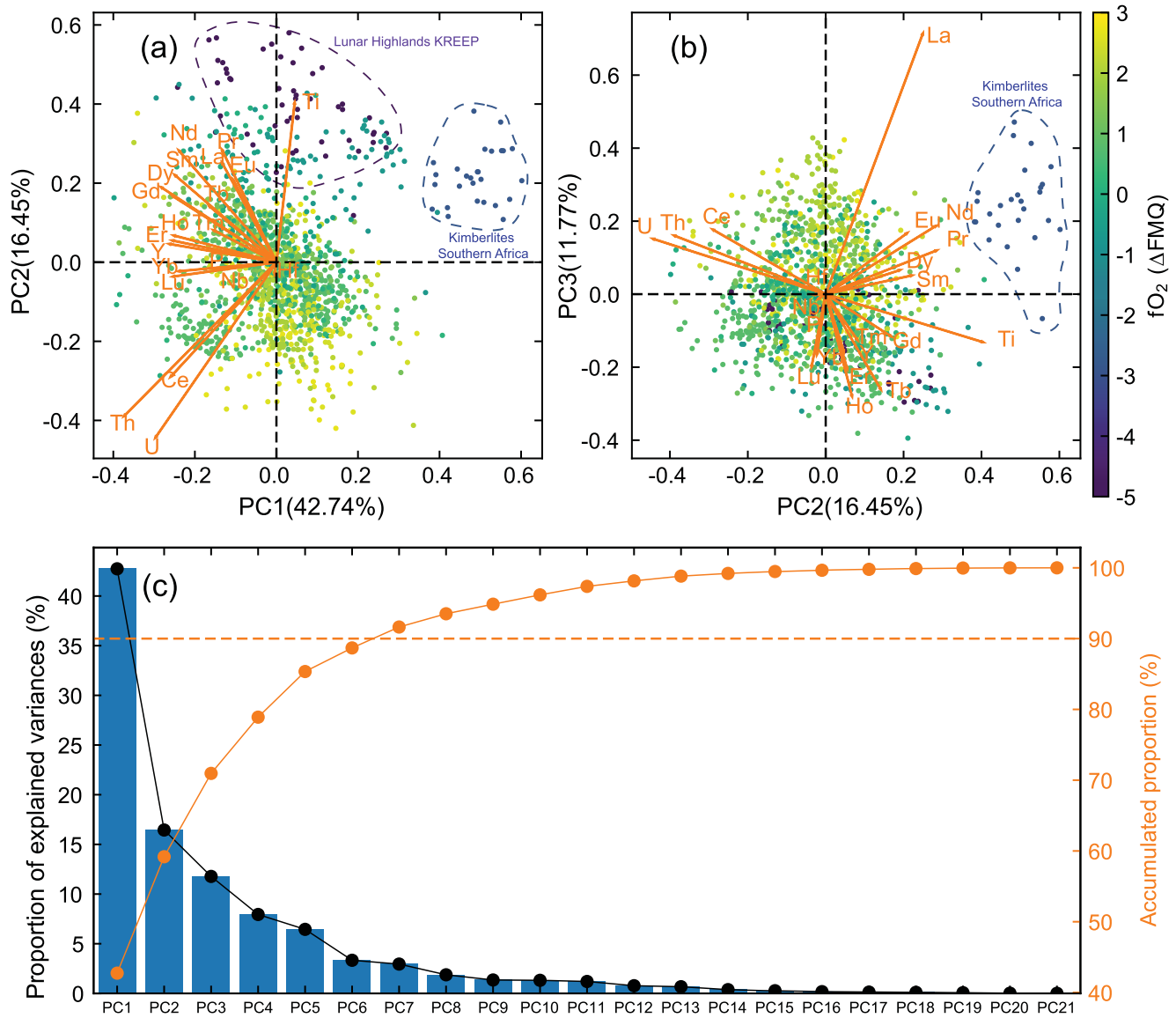


Figure 2

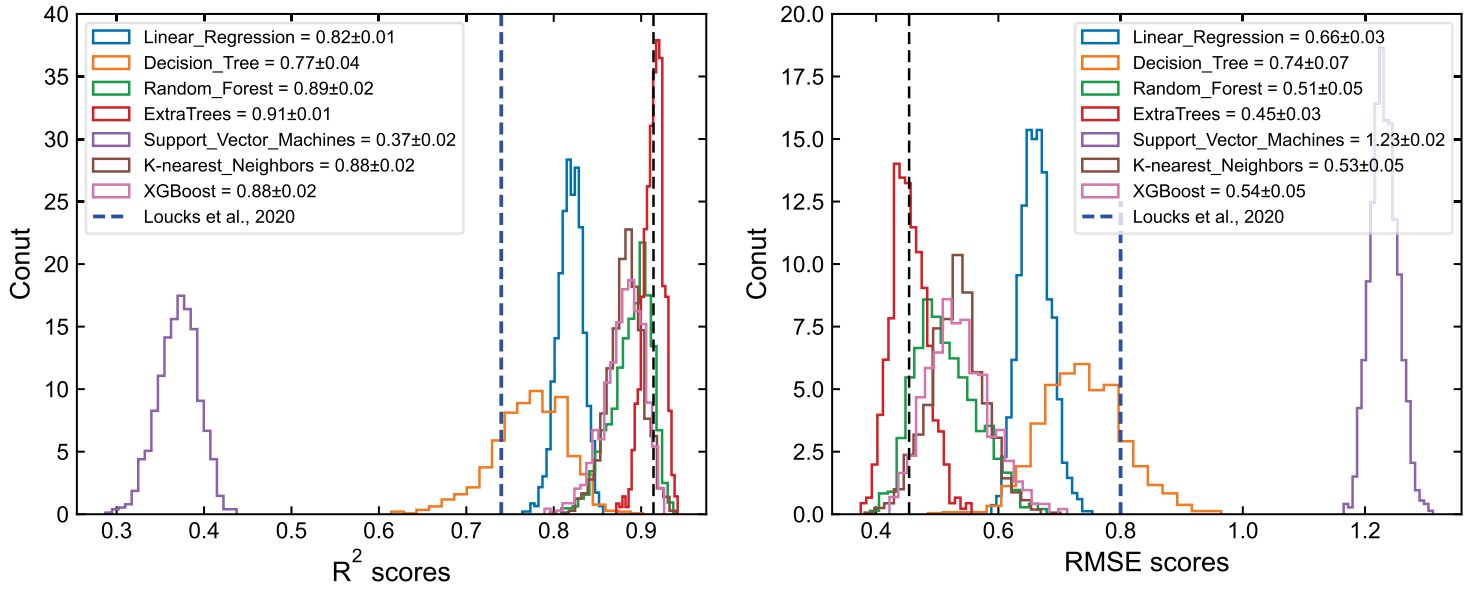


Figure 3

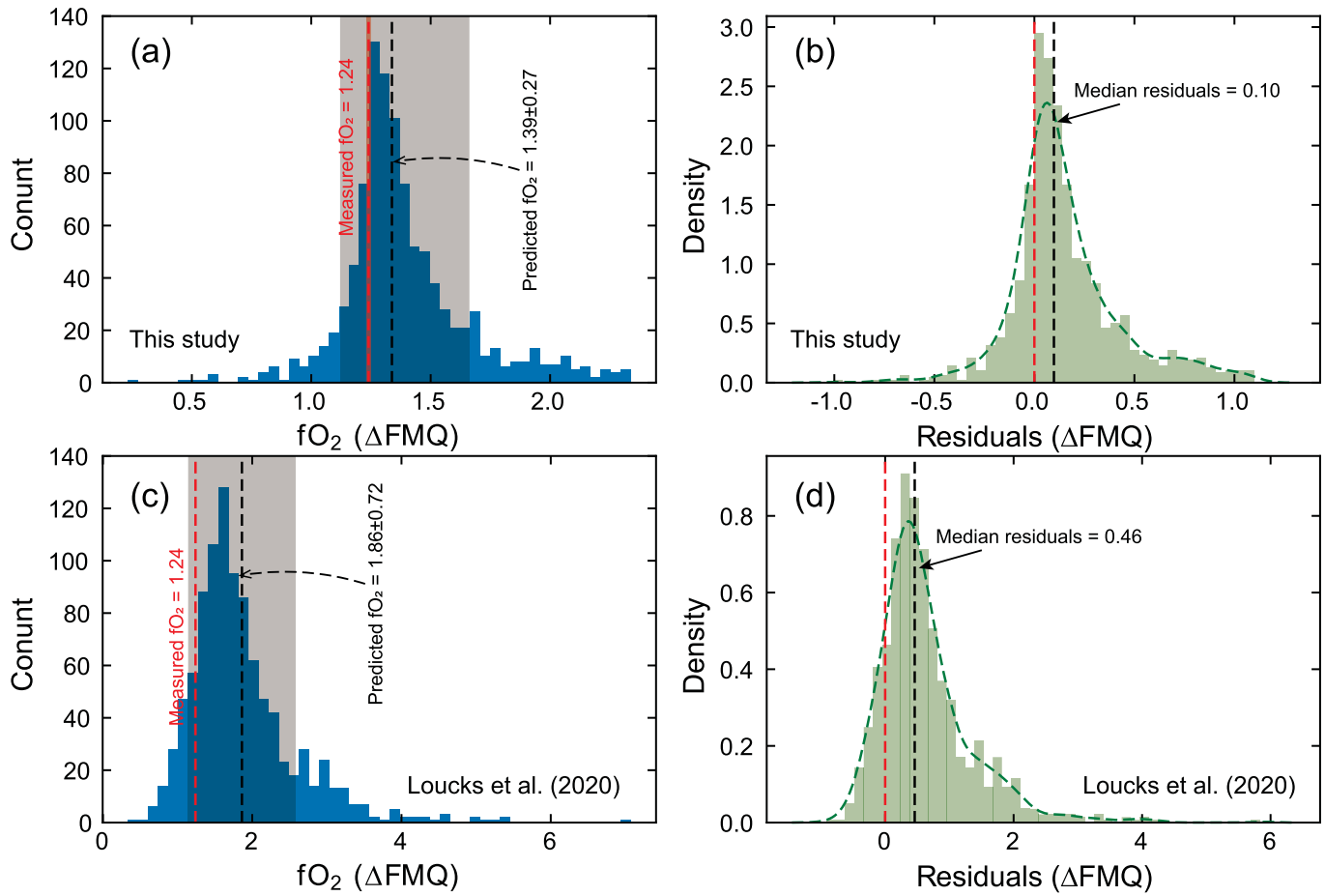


Figure 4

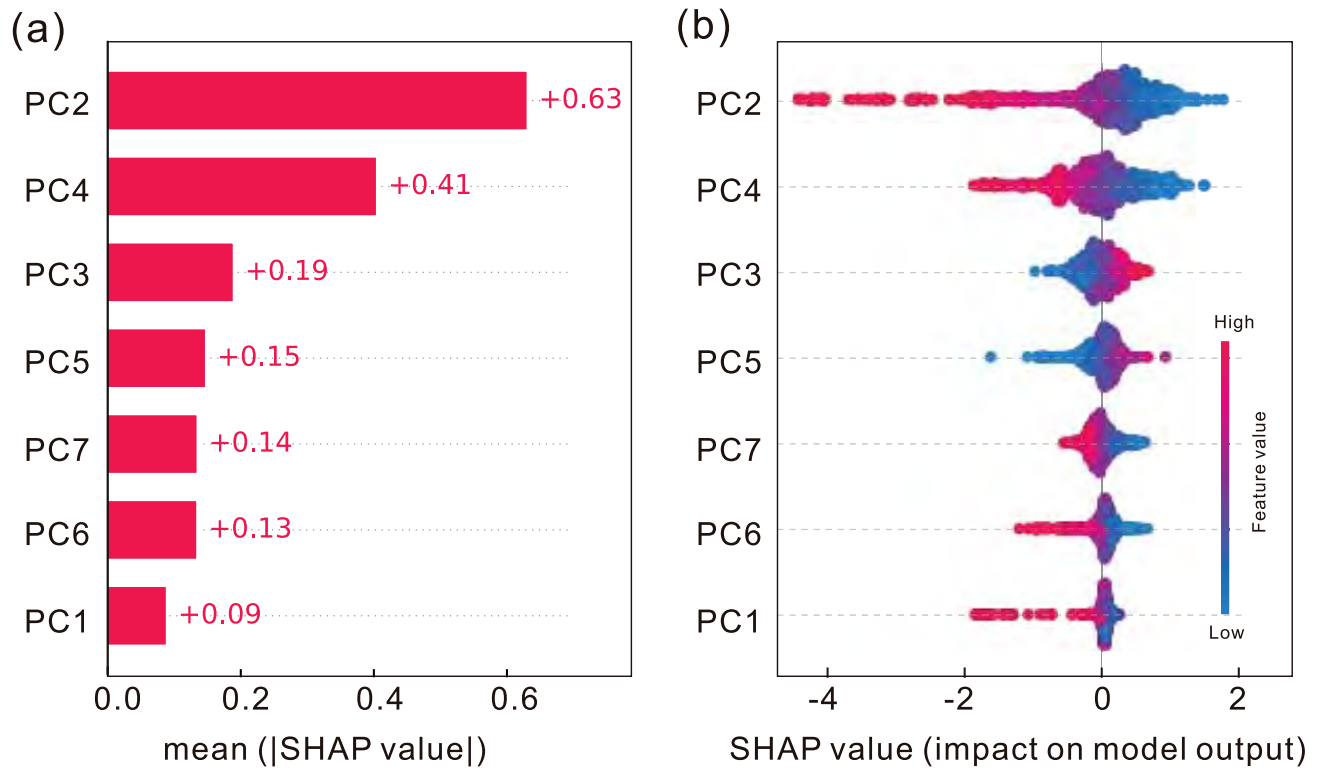


Figure 5

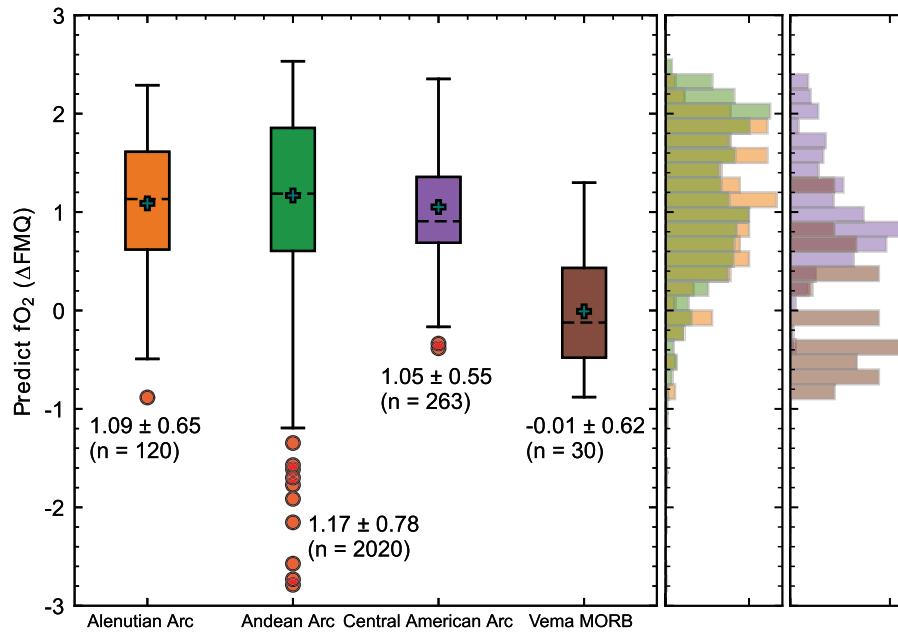


Figure 6

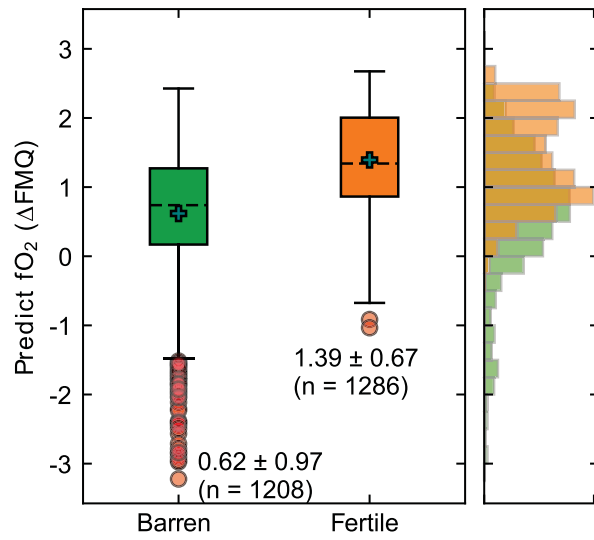


Figure 7

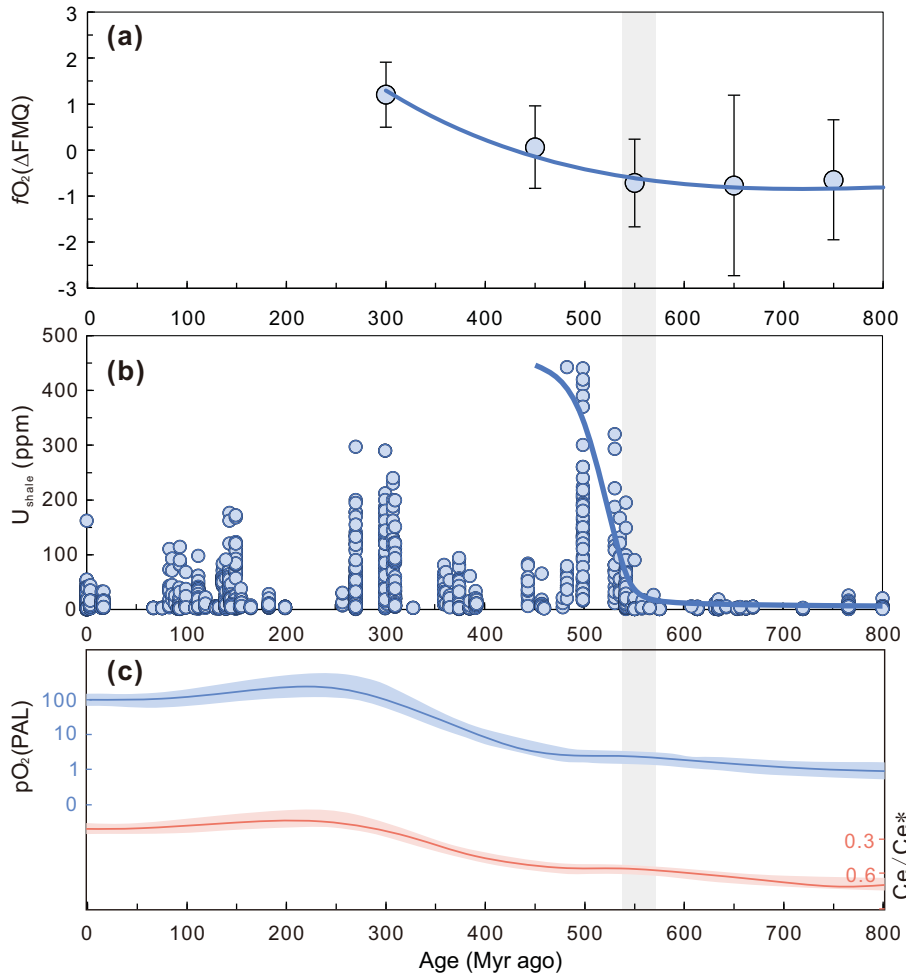


Figure 8