

## **Supplementary materials**

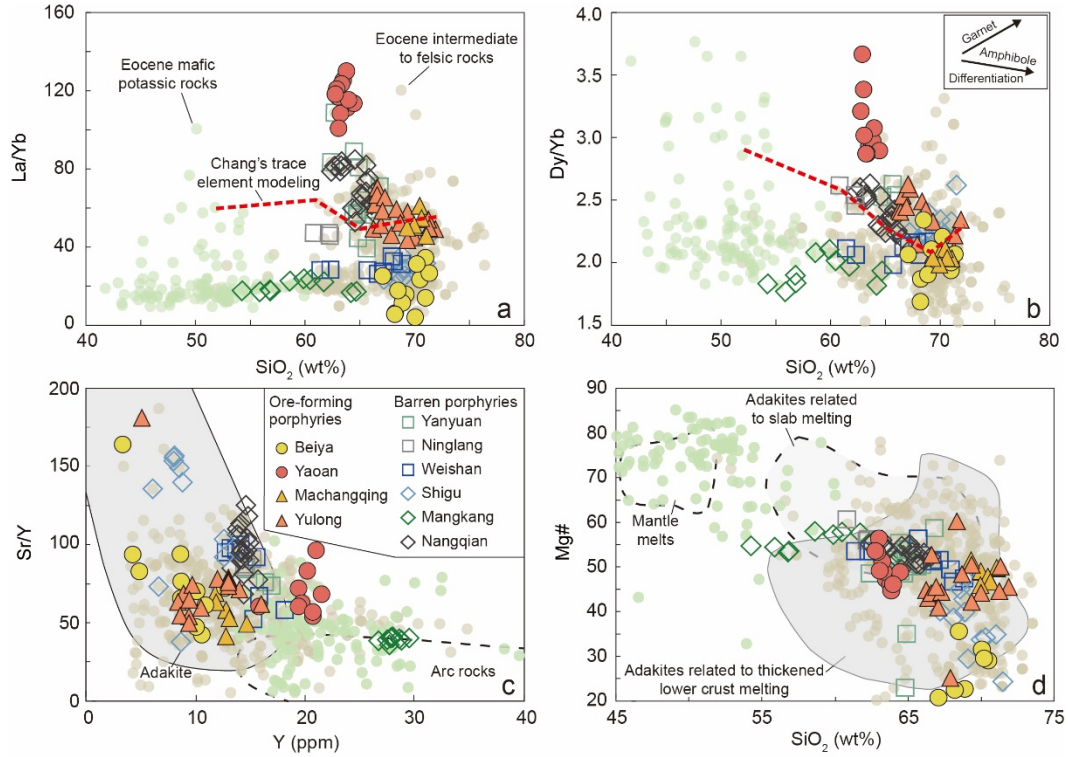
### **Analytical methods**

Apatite, amphibole, and plagioclase analyses were performed using an electron microprobe (JXA-8230) at the State Key Laboratory of Continental Dynamics, Northwest University, China. Samples were re-polished prior to electron probe microanalysis (EPMA) to remove any compositional modification induced by SEM electron-beam exposure, and subsequently carbon coated along with secondary standards to avoid variable light element X-ray attenuation. The instrument was operated at an accelerating voltage of 15 kV, beam current of 10 nA and beam diameter of 1  $\mu\text{m}$ . Where possible, apatite crystals were analysed with the c-axis parallel to the plane of the mount. This routine limits the potential for time dependent variability in halogen X-ray counts during analysis (Stock et al., 2016, 2018), while maintaining reasonable precision for low-concentration elements (i.e., Cl). Count times were 20–30s for major elements and 30–90s for minor elements (120s for Cl and SO<sub>2</sub> in apatite). Natural minerals and synthetic oxides were used as standards, including andradite for Si and Ca, rutile for Ti, corundum for Al, hematite for Fe, eskolaite for Cr, rhodonite for Mn, bunsenite for Ni, periclase for Mg, albite for Na, and K-feldspar for K. Matrix corrections were performed using the ZAF correction program supplied by the instrument manufacturer.

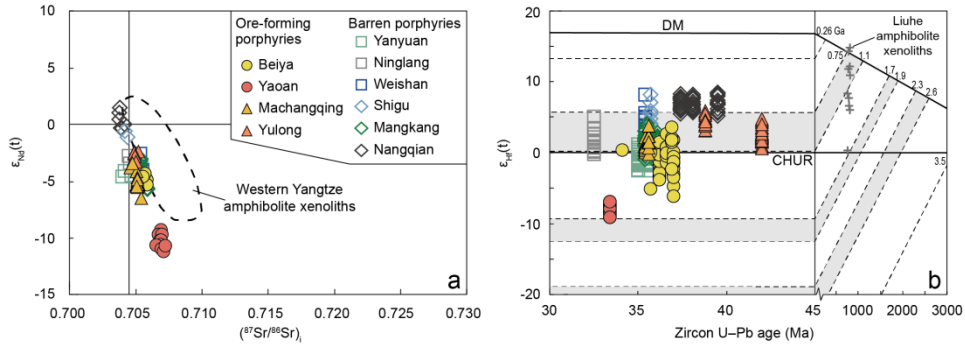
### **Rhyolite-MELTS fractionation modelling**

Isobaric fractional crystallisation models were run using Rhyolite-MELTS (Gualda et al., 2012) to constrain the conditions of magma storage. We used the melt composition from partial melting experiments on a synthetic as the starting composition for our Rhyolite-MELTS models (melt C-3136 of Qian and Hermann, 2013), as distinguished by the similar partial melting condition with the lower crust of Sanjiang metallgenic belt.

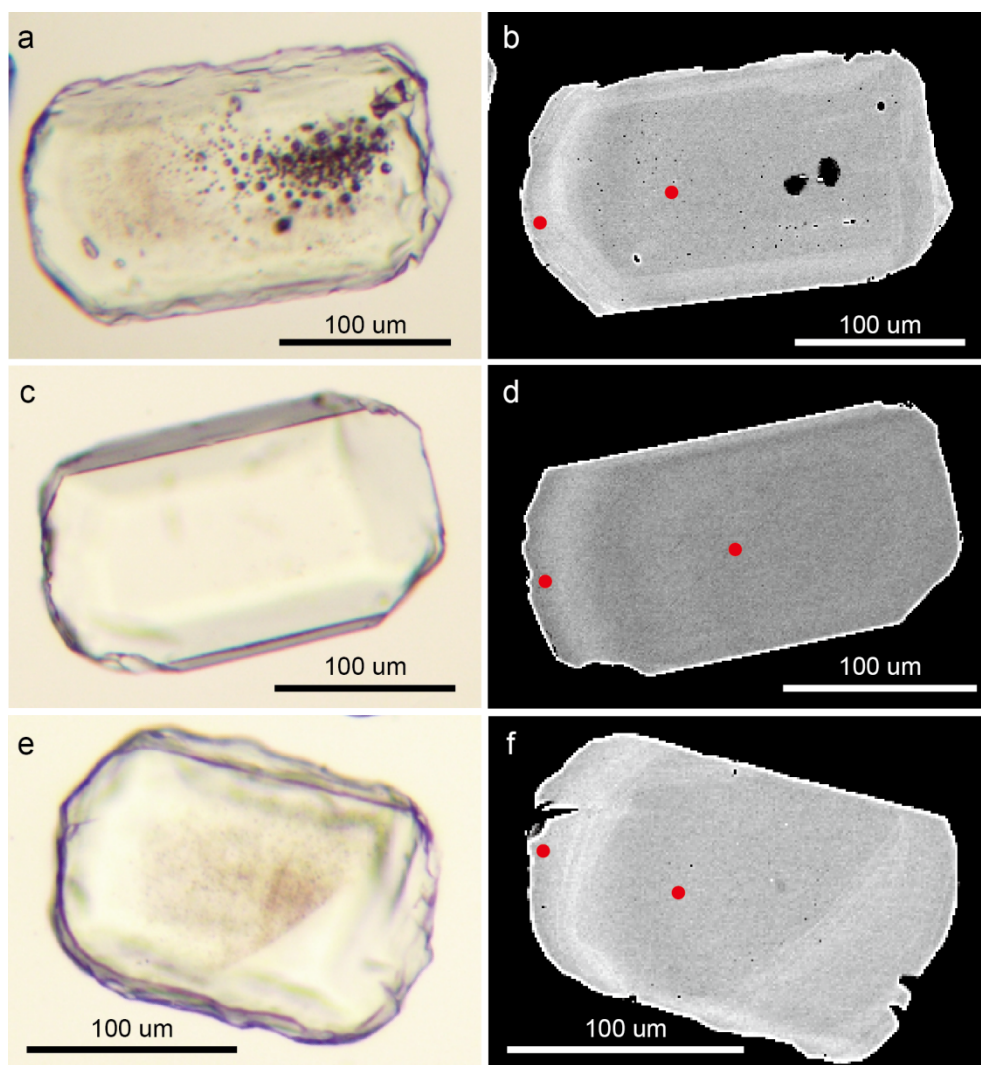
## Supplementary Figures



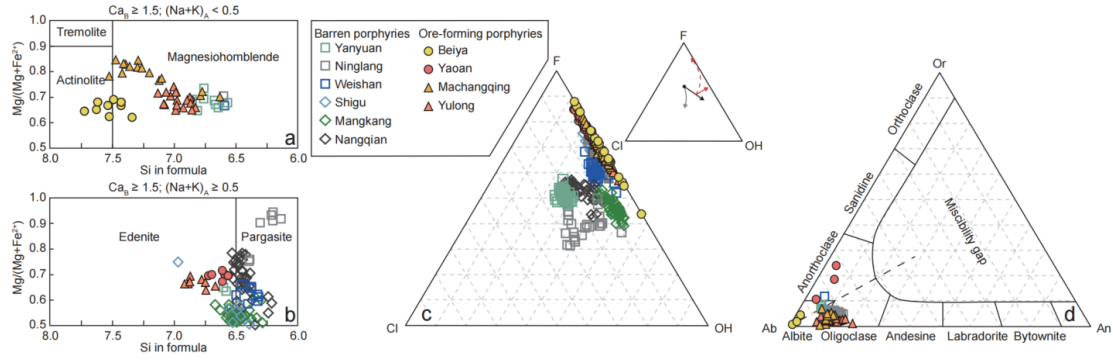
**Supplementary Figure 1.** Discrimination diagrams for Ore-forming and barren porphyries in the Sanjiang metallogenic belt. **a**, La/Yb ratios vs SiO<sub>2</sub> contents. **b**, Dy/Yb ratios vs SiO<sub>2</sub> contents. **c**, La/Yb ratios vs SiO<sub>2</sub> contents. **d**, Mg# vs SiO<sub>2</sub> contents. The red dashed lines in **a** and **b** represents the mafic rock fractionation modeling by Chang and Audétat (2023). The adakite and arc rocks fields are according to Defant and Drummond (1990). Fields indicating mantle melt, adakites related to slab melting and lower-crustal melting are from Condie (2005), Wang et al. (2006) and Zheng et al. (2012).



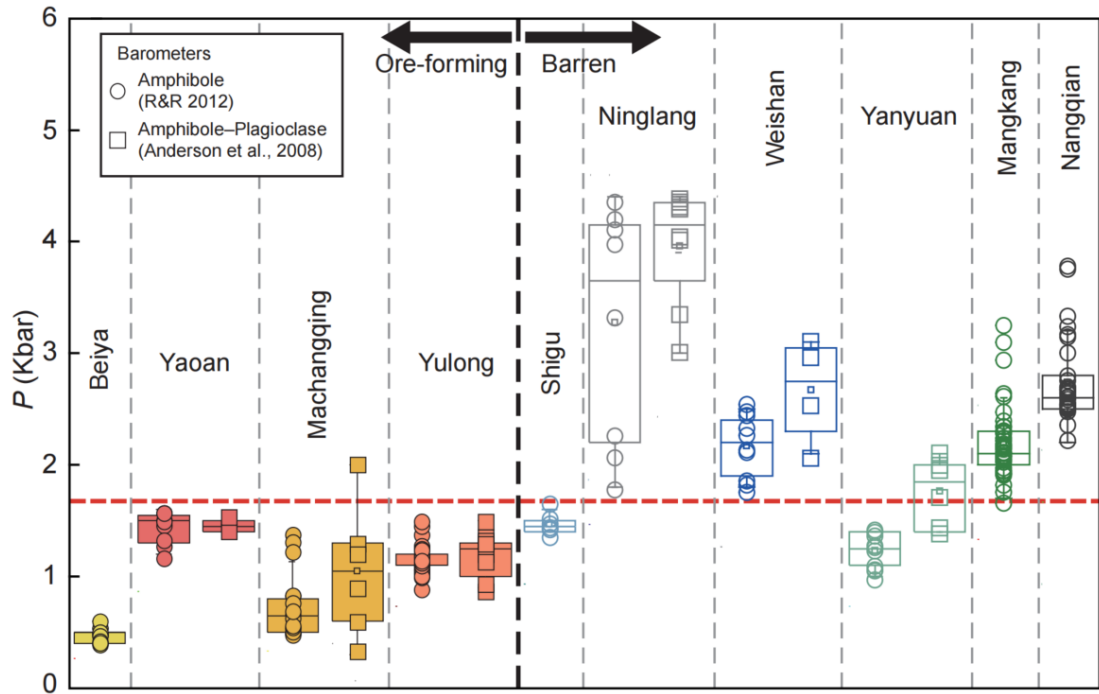
**Supplementary Figure 2.** Variation of **a**,  $(^{87}Sr/^{86}Sr)_i$  vs  $\epsilon_{Nd}(t)$  and **b**, initial  $\epsilon_{Hf}(t)$  isotope values vs zircon U–Pb ages. The field for western Yangtze amphibolite xenoliths is from Deng et al. (1998), Zhao et al. (2004) and Zhou et al. (2017). CHUR, chondrite uniform reservoir; DM, depleted mantle. The values used for constructing the depleted mantle (DM) and crustal evolution reference lines are from Griffin et al. (2000, 2002). The light gray fields represent episodes of major juvenile crustal growth in the Yangtze craton (Sun et al., 2009). Neoproterozoic amphibolite xenoliths in the Liuhe area (Hou et al., 2017; Zhou et al., 2017)



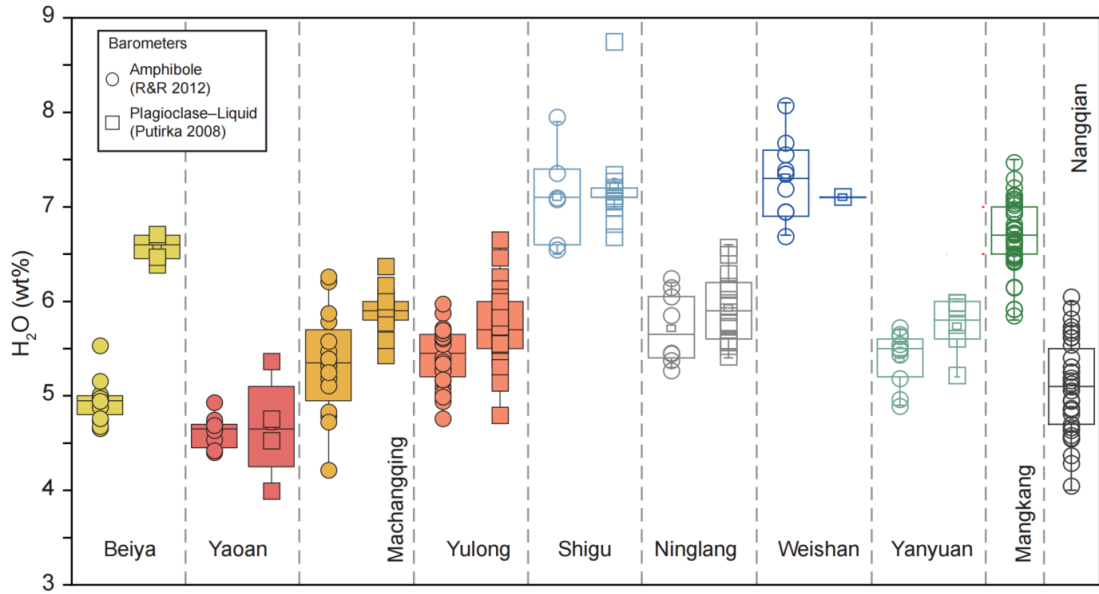
**Supplementary Figure 3.** (a, c, e) Photomicrograph and (b, d, f) backscattered electron images of apatites. The red spots show the positions of electron probe microanalysis (EPMA). All analyses of microphenocryst cores and rims are indistinguishable and show no evidence of re-equilibration with a volatile-saturated melt.



**Supplementary Figure 4. a and b**, Classification of amphibole (Leake et al., 1997). **c**, Ternary space of volatile compositions of apatites. **d**, Composition of feldspars in the An–Ab–Or ternary diagram (Smith, 1974). The representative ternary graph on the upper right illustrates theoretical apatite compositional trajectories for different crystallization scenarios. Apatite crystallization begins at the black point (see discussion of input parameters in the text). The continuous lines show apatite compositional evolution during H<sub>2</sub>O-undersaturation fractional crystallization with Dc/m Cl  $\approx$  0.6 (grey continuous line) and Dc/m Cl  $\approx$  0.9 (black continuous line). The red lines show apatite compositional evolution during H<sub>2</sub>O-saturation fractional crystallization under isobaric condition (0 wt% H<sub>2</sub>O loss; red continuous line) and polybaric condition (0.18 wt% H<sub>2</sub>O loss; red dotted line), after 40% crystallization under H<sub>2</sub>O-undersaturated condition. Amphiboles and apatites are from ore-forming and barren adakite-like porphyries in Sanjiang metallogenic belt.



**Supplementary Figure 5.** Box and whisker plots of comparison of different barometer results. The dots represent the results calculated by amphibole-only barometer (Ridolfi and Renzulli, 2012), and the squares represent the results calculated by amphibole-plagioclase barometer (Anderson et al., 2008). Boxes show first to third quartile range with bars showing extremes of data (excluding any outliers). Long and short lines in box show square and mean value, respectively. The points outside boxes represent their extremum values.



**Supplementary Figure 6.** Box and whisker plots of comparison of different hygrometer results. The dots represent the results calculated by amphibole-only hygrometer (Ridolfi and Renzulli, 2012), and the squares represent the results calculated by plagioclase–liquid hygrometer (Putirka et al., 2008). The box and whisker meanings are the same as those in Supplementary Figure 5.



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