

Phase transformation of ferric-iron-rich silicate in Earth's mid-mantle

MINGDA LV^{1,†}, SHENGCAI ZHU^{2,*}, JIACHAO LIU¹, YI HU^{3,4}, FENG ZHU^{5,6,‡}, XIAOJING LAI^{3,4,7},
DONGZHOU ZHANG^{3,§}, BIN CHEN³, PRZEMYSŁAW DERA³, JIE LI^{5,||}, AND SUSANNAH M. DORFMAN^{1,*}

¹Department of Earth and Environmental Sciences, Michigan State University, East Lansing, Michigan 48824, U.S.A.

²School of Materials, Sun Yat-Sen University, Guangzhou 510275, China

³Hawai'i Institute of Geophysics and Planetology, University of Hawai'i at Mānoa, Honolulu 96822, Hawaii, U.S.A.

⁴Department of Geology and Geophysics, School of Ocean and Earth Science and Technology, University of Hawai'i at Mānoa, Honolulu, Hawaii 96822, U.S.A.

⁵Department of Earth and Environmental Sciences, University of Michigan, Ann Arbor, Michigan 48109, U.S.A.

⁶State Key Laboratory of Geological Processes and Mineral Resources, School of Earth Sciences, China University of Geosciences, Wuhan 430074, China

⁷Gemological Institute, China University of Geosciences, Wuhan, Hubei, China

ABSTRACT

The incorporation of ferric iron in mantle silicates stabilizes different crystal structures and changes phase transition conditions, thus impacting seismic wave speeds and discontinuities. Recent experiments of $\text{MgSiO}_3\text{-Fe}_2\text{O}_3$ mixtures indicate the coexistence of fully oxidized iron-rich ($\text{Mg}_{0.5}\text{Fe}_{0.5}^{3+}$)($\text{Fe}_{0.5}^{3+}\text{Si}_{0.5}$) O_3 with Fe-poor silicate (wadsleyite or bridgmanite) and stishovite at 15 to 27 GPa and 1773 to 2000 K, conditions relevant to subducted lithosphere in the Earth's transition zone and uppermost lower mantle. X-ray diffraction (XRD) shows that ($\text{Mg}_{0.5}\text{Fe}_{0.5}^{3+}$)($\text{Fe}_{0.5}^{3+}\text{Si}_{0.5}$) O_3 recovered from these conditions adopts the $R3c$ LiNbO_3 -type structure, which transforms to the bridgmanite structure again between 18.3 and 24.7 GPa at 300 K. XRD data are used to obtain the equation of state of the LiNbO_3 -type phase up to 18.3 GPa. Combined with multi-anvil experiments, these observations suggest that the stable phase of ($\text{Mg}_{0.5}\text{Fe}_{0.5}^{3+}$)($\text{Fe}_{0.5}^{3+}\text{Si}_{0.5}$) O_3 is bridgmanite at 15–27 GPa, which transforms on decompression to LiNbO_3 -type structure. Our calculation revealed that ordering of the ferric ion reduces the kinetic energy barrier of the transition between ($\text{Mg}_{0.5}\text{Fe}_{0.5}^{3+}$)($\text{Fe}_{0.5}^{3+}\text{Si}_{0.5}$) O_3 LiNbO_3 structure and bridgmanite relative to the MgSiO_3 akimotoite-bridgmanite system. A dense Fe^{3+} -rich bridgmanite structure is thus stable at substantially shallower depths than MgSiO_3 bridgmanite and would promote subduction.

Keywords: LiNbO_3 , bridgmanite, ferric iron, transition zone, lower mantle, phase transition, equation of state