

Nanoscale mapping of ZrSiO₄ phases in naturally shocked zircon using electron energy loss spectroscopy

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ABSTRACT

Coexistence in natural samples of zircon (ZrSiO₄) and reidite (a high-pressure polymorph of ZrSiO₄) is attributed to the effects of hypervelocity impact events. The grains and intergrowths in those minerals can be merely a few nanometers in size, which makes phase identification by standard methods of structure analysis difficult. However, analytical scanning transmission electron microscopy (STEM) utilizing electron energy loss spectroscopy (EELS) can provide important information on phase transition mechanisms and pressure-temperature conditions associated with the shock event at the nanoscale. Here we demonstrate that the valence as well as oxygen core-loss EELS can be employed for nanoscale mapping of zircon-reidite distributions in zircon-reidite aggregates. Moreover, other accompanying phases, e.g., baddeleyite, could also be identified and mapped by this method. We further compare the EELS maps with a 4D-STEM nanobeam precession electron diffraction data, and demonstrate the advantages of the EELS mapping, which provides spatial resolution down to the nanometer scale and is independent of crystal orientation.

Keywords: Zircon, reidite, EELS, phase mapping, Kara impact structure

INTRODUCTION

The natural occurrence of reidite, the high-pressure phase of zircon (ZrSiO₄), is regarded as one of the criteria for shock metamorphism (Reddy et al. 2015). However, the mechanism and exact *P-T-t* conditions of the zircon-reidite phase transition remain unclear (Erickson et al. 2017; Mihailova et al. 2019). In natural and synthetic samples, the presence of reidite has been confirmed by: X-ray powder diffraction (Kusaba et al. 1985; Knittle and Williams 1993; Glass et al. 2002); Raman and infrared spectroscopy (Gucsik et al. 2002, 2004; Wittmann et al. 2006); and electron diffraction patterns in transmission electron microscopy (TEM) (Leroux et al. 1999). In the last 10 years, the most common method to detect reidite has relied on a combination of scanning electron microscopy (SEM) backscattered electron (BSE) imaging, energy-dispersive spectroscopy (EDS) analyses, and electron backscattered diffraction (EBSD) (Reddy et al. 2015; Erickson et al. 2017; Xing et al. 2020; Plan et al. 2021; Zhao et al. 2024).

However, SEM imaging, EBSD mapping and Raman spectroscopy do not provide a spatial resolution high enough to access the complexity of the impact-produced Zr-phases, their reaction products and deformation microstructures (Kovaleva et al. 2021, 2023). Mineral grains and intergrowths in such samples can be as small as a few nanometers in size and thus below the resolution of SEM imaging (Kovaleva et al. 2021; Zhao et al. 2024) and Raman spectroscopy (Wittmann et al. 2006), which is

a serious limitation leading to imprecise interpretations of processes. This problem was highlighted by Wittmann et al. (2006), who suspected the presence of an orthorhombic-II (OII) polymorph of ZrO₂ in impact glasses, but could not conclusively confirm it with Raman spectroscopy.

On the other hand, electron energy loss spectroscopy (EELS) in combination with scanning TEM (STEM) is a powerful tool used to extract important information in materials science (Colliex 2022). EELS chemical mapping is routinely employed to visualize phase and element distribution down to the atomic scale (Mundet et al. 2021; Yeh et al. 2023). Moreover, EELS is capable of distinguishing allotropes of carbon based on their fine *K*-edge structure (Egerton et al. 1974). Generally, the low energy or valence region of EELS (VEELS) is utilized to measure band gaps in materials and to map the distribution of phases in hard and soft materials (Chen et al. 2024; Mu et al. 2016; Badr et al. 2023; Liang et al. 2018; Rossner et al. 2018). At the same time, the comprehensive interpretation of VEELS regions is very complicated, because they include various collective excitations (plasmons) as well as inner- and intraband transitions (Egerton 1996).

In this paper, we compare the VEELS and core-loss regions of reidite and zircon spectra and use them to visualize the nanoscale distribution of ZrSiO₄ phases in naturally shocked zircon grains from the Kara impact structure located in polar Russia. We document a simple and fast method of micro- and nanophase identification with very high spatial resolution using STEM. Using the obtained high-resolution data, we are improving constraints on the phase transition mechanism and impact conditions. The presented workflow can also be applied to any other geological samples with nanoscale textures and grain sizes for phase identification and mapping with very high spatial resolution.

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MATERIALS AND METHODS

Materials

The 65 km diameter, 70.3 Ma Kara impact structure (Trieloff et al. 1998) (Fig. 1a) is located in the Yugorsky Peninsula, Russia. The target rocks of the Kara structure are a folded terrigenous and carbonate Upper Proterozoic to Paleozoic sedimentary sequence, with the Ordovician–Carboniferous and early Permian sedimentary rocks topping the sequence (Masaitis 1999). Massive lithic breccias are overlain by allogenic suevites and impact melt rocks (tagamites). Suevites are composed of sandstone, limestone and slate, glass bombs, and veins, embedded in the fine-grained crushed matrix. Samples were collected from the crater-infill suevite breccias. Sample KK21-21 (GPS: 68°54′07.8″N 64°33′48.2″E) is a gray suevite in contact with the bedrock limestone (Fig. 1b). Sample KH21-08 (GPS: 68°56′50.9″N 64°50′26.8″E) is a brown suevite with large fluidal-textured glass fragments (Fig. 1c). Foil #7081 is from irregular zircon grain hosted by the fine-grained clastic matrix of a lithic fragment, likely a poorly sorted sandstone (sample KK21-21). The grain has characteristic subparallel short fractures and <1 μm thick reidite inclusions (bright in BSE) (Fig. 1d). Reidite forms two sets of parallel lamellae and an irregular network of thinner veinlets (Fig. 1d). Foil #7103 is from subhedral zircon embedded

into a lithic clast that resembles weakly foliated slate (sample KK21-08). Zircon has patchy flame-shaped inclusions of reidite in the core (BSE-bright) and radial fractures in the rim (Fig. 1e). Foil #7080 is from an aggregate at the boundary between a sandstone clast and fine-grained glass-rich matrix (sample KK21-21, Online Materials¹ Fig. S11). The aggregate has a porous texture and is composed of two fragments with radiating spherulitic internal texture (Fig. 1f).

Methods

TEM foils of approximately 15 × 10 μm were prepared by a lift-out method using a focused ion beam (FIB) Thermo Fisher Scientific (TFS) Helios G4UC instrument. TEM experiments were done using a TFS Themis Z 80-300 (S) TEM operated at 300 kV, and equipped with a TFS SuperX Energy-dispersive X-ray (EDX) detector, and a Gatan Imaging Filter (GIF) Continuum 1065. The individual reference EELS spectra and spectrum images (SIs) of studied specimens were collected with 0.15 and 1.5 eV dispersion using a Gatan Microscopy Suite (GMS or DigitalMicrograph) software. Using the Fourier-log ratios, the estimated specimen thickness in terms of the effective mean-free path (λ_{eff}) for inelastic scattering was ~0.7 and, based on the Kramers-Kronig analysis (Kronig 1926;

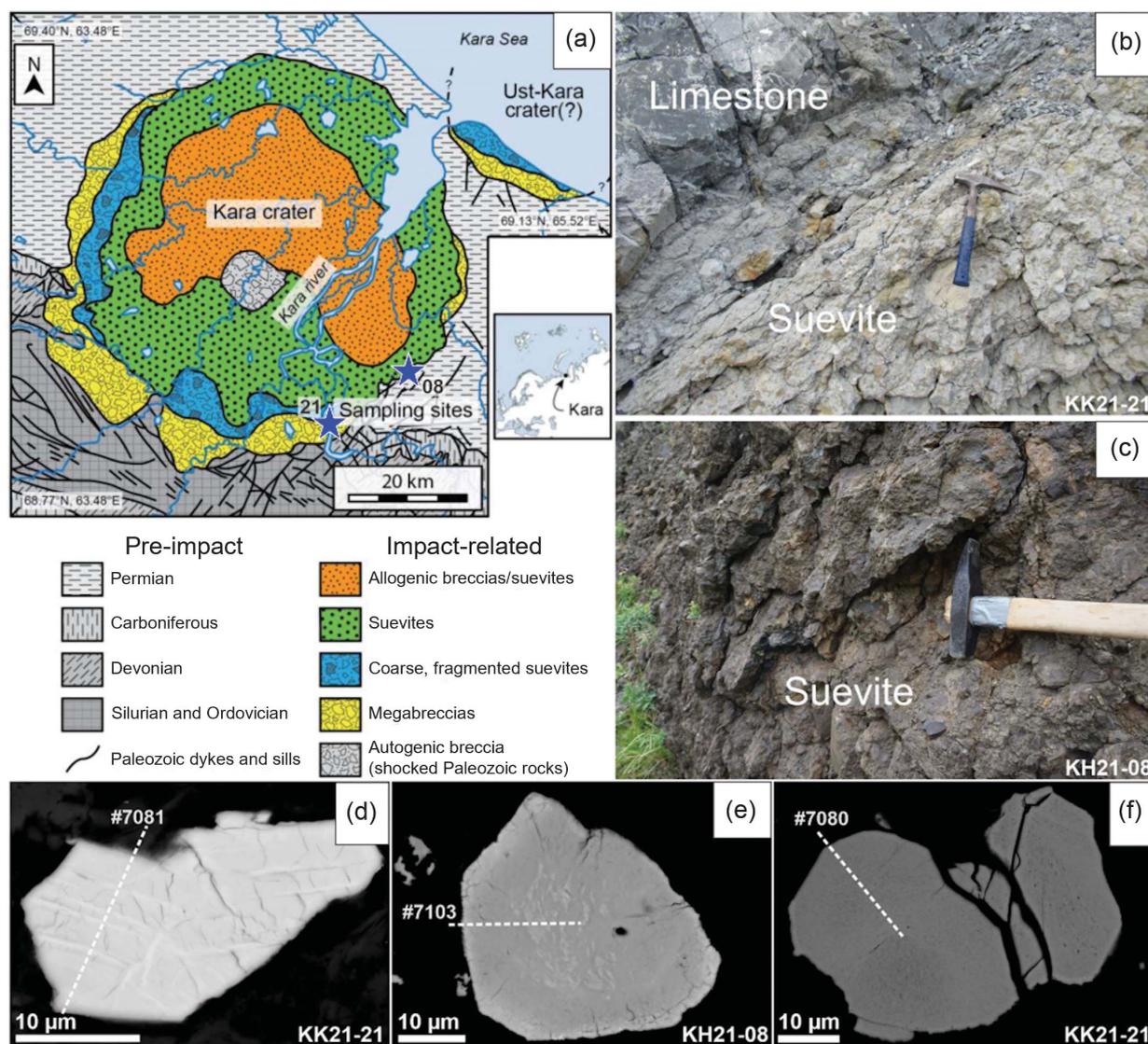


FIGURE 1. (a) Schematic geological map of Kara impact structure from Kenny et al. (2021). Blue stars indicate sampling locations from the crater rim. (b–c) Field photographs of the sampled outcrops. In (b), light yellow-gray suevite intercalates with the blocks of gray shattered limestone. (d–f) BSE images of studied ZrSiO₄ aggregates. The broken lines indicate the positions from which the TEM samples were cut.

Kramers 1927), corresponds to an approximately 70 nm FIB foil thickness. Collected reference EELS spectra of zircon and reidite with their characteristic features were subjected to zero-loss peak fitting with Gaussian-Lorentz functions, followed by deconvolution and background subtraction (Johnson and Spence 1974), using the GMS built-in functions. The reference spectra prepared this way were used for multiple linear least-squares (MLLS) fitting (Leapman and Swyt 1988; Yakovlev and Libera 2008) of individual EELS spectra from the corresponding SIs to visualize the distribution of phases in the TEM specimens. The GMS software features a built-in routine for performing MLLS fitting within a specified energy range (e.g., 10–60 eV). Electron diffraction data were collected and analyzed using a precession electron diffraction (PED) DigiSTAR device and the ASTAR software (NanoMEGAS SPRL, Belgium). The pole figures were prepared with an EDAX OIM Analysis software.

RESULTS

In the high-angle annular dark-field (HAADF) image (Fig. 2a), reidite forms sets of thin lamellae inside fractured zircon. The EDX chemical maps indicate the absence of other phases (Online Materials¹ Fig. SI2). Zircon and reidite were also identified using selected area electron diffraction (SAED). The SAED pattern of reidite in $[110]$ zone axis orientation shows multiple forbidden spots and broadening of high order with

asterism, which indicates a high density of defects and structural disorder of the reidite lamellae (Fig. 2b).

The positions and shapes of VEELS peaks of both phases are almost independent of the specimen orientation with respect to the incident electron beam (Fig. 2d; Online Materials¹ Fig. SI3). The VEELS of zircon is consistent with previously reported spectra (Jiang and Spence 2013). Zircon and reidite have distinct VEELS spectra (Fig. 2d), namely the largest peak at 25.5 eV associated with bulk plasmon in zircon is shifted by 2.3 to 27.6 eV in reidite. Other spectrum features are much less different. For example, a bump at around 8.0 eV is present in spectra from both phases, but is more pronounced in zircon. A small peak at 13.5 eV is similar in both phases. Also, zircon has a peak at 16.4 eV, while reidite has two at 15.3 and 17.8 eV (Table 1). The fine features between 30 and 40 eV in the VEELS (Fig. 2d) consist of two small peaks centered at ~ 33.5 and 35.5 eV in zircon, similar to those described by Jiang and Spence (2013) but absent in reidite. In zircon, the delayed maximum peak position is at about 41.1 eV with a shoulder at the right side, while in reidite it is at ca. 40 eV (Fig. 2d; Table 1).

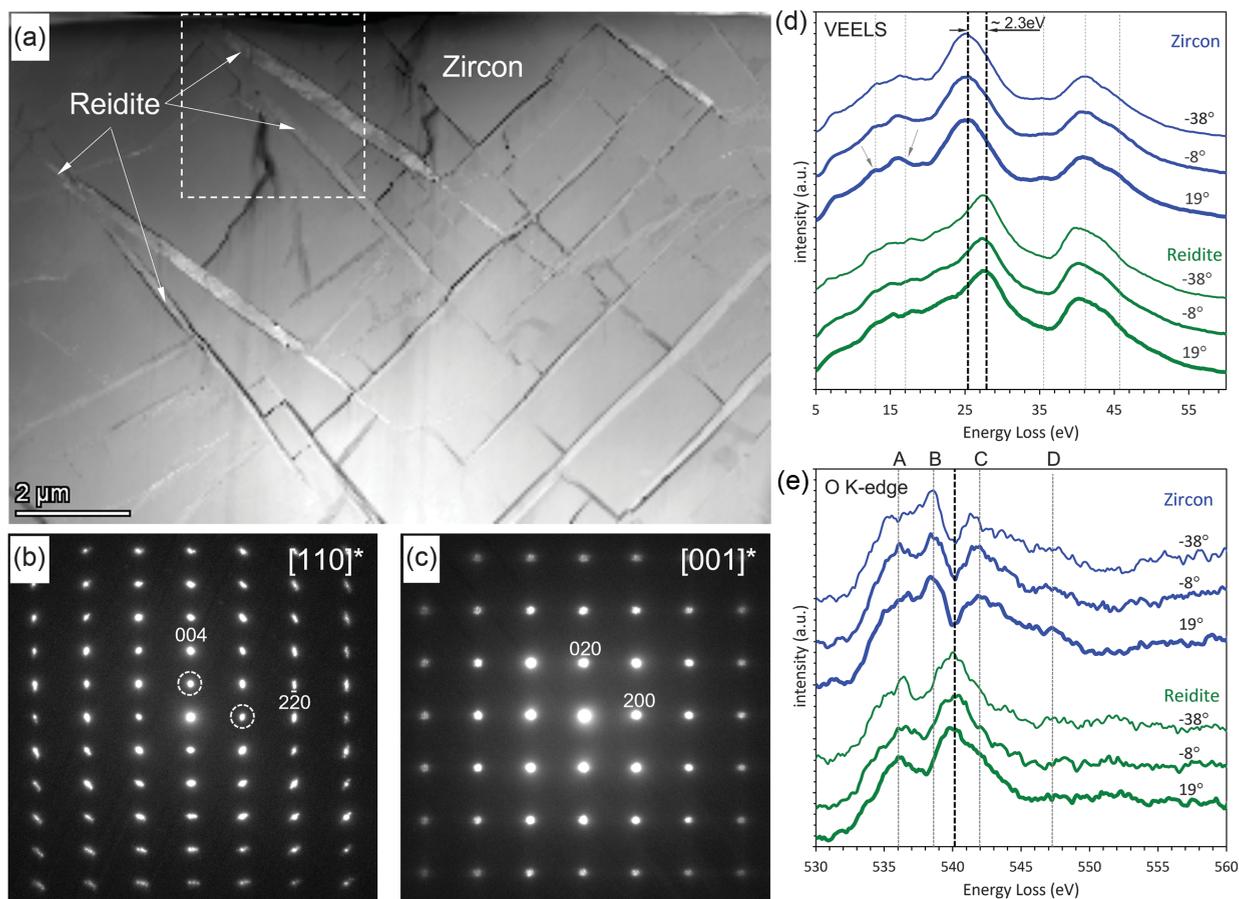


FIGURE 2. TEM images and analyses of FIB foil #7081. **(a)** HAADF image with identified lamellae of reidite (white arrows) and the fractured zircon. The dotted box shows the position of the phase map in Figure 3. **(b)** SAED pattern of reidite for the $[110]$ crystallographic direction; forbidden spots are marked with dashed circles; **(c)** SAED pattern of zircon for the $[001]$ direction. **(d)** Characteristic VEELS and **(e)** oxygen K -edge spectra of zircon (blue spectra) and reidite (green spectra) at different tilting angles (from bottom to top of each set represents the angles $+19$, -8 , and -38°). Black dashed lines show the positions of the main peaks of zircon and reidite shifted by ~ 2.3 eV. The pronounced peaks in the near-edge region (535–555 eV) of the O K -edge ELNES of zircon are marked A–D after Jiang and Spence (2009).

TABLE 1. The list of features in the VEELS spectra of zircon and reidite, along with their assignments, and the band gap energies (E_g)

Zircon				Reidite			
Bulk	Surface	Interband	Assessment	Bulk	Surface	Interband	Assessment
~7.4	~7.4	>7	O ₂ p→Zr ₄ d	~7.4	~7.4	>7	O ₂ p→Zr ₄ d
	12.2		Surface plasmon		12.2		Surface plasmon
13.5	14.6	14.0	O ₂ p→Si ₃ s	~13	14	>7	O ₂ p→Si ₃ s
16.4			Bulk plasmon	15.3			Bulk plasmon
				17.8			Bulk plasmon
	21.2		O ₂ p→Zr ₅ s O ₂ s-Op/Zr ₄ d		21.2		O ₂ p→Zr ₅ s O ₂ s-Op/Zr ₄ d
25.5			Bulk plasmon	27.6			Bulk plasmon
32.0–38.0			O ₂ s→Zr ₅ p	32.0–38.0			O ₂ s→Zr ₅ p
~41.1			Zr ₄ p→Zr ₄ d	~40.7			Zr ₄ p→Zr ₄ d

Note: The energy values are given in eV.

The oxygen *K*-edge spectra of the two phases also differ (Fig 2e). A single peak at about 540 eV in the oxygen *K*-edge of reidite is visible (Fig. 2e, black dashed line), while distinct peaks B and C occur in the oxygen *K*-edge of zircon. Also, there are two minor features between C and D peaks. The peaks A and B in the O *K*-edge EELS have been attributed to Zr-O interaction. Peak C in the O *K*-edge EELS appears due to the Si-O interaction, while peak D has both Si and Zr characteristics (Jiang and Spence 2009). The O *K*-edge ELNES of reidite shows a broad peak instead of two separate B and C peaks characteristics for zircon (Fig. 2e, thick dashed line), while peaks A and D in reidite both coincide with those in zircon. Interestingly, looking at the tilt series it becomes clear that the intensity of the A peak in zircon is dependent on the specimen tilt, while this effect is less pronounced in reidite. EELS spectra showing *K*- and *L*-edges of Si and *M*-edge of Zr are presented in Online Materials¹ Figure SI4. One can also notice differences between the spectra of zircon and reidite; however, the spectra are less suitable for mapping because of the poor signal-to-noise ratio at higher energy losses.

Using the EELS spectra, phase maps were produced (Fig. 3), where even thin (<30 nm) lamellae of reidite could be reliably

identified. The PED from the same area (Figs. 3e–3f) did not reveal the thin lamellae of reidite entirely because of the large tilt (~14°) from the [110] direction and shadowing from the surrounding zircon. Moreover, the size of the parallel electron beam (2–3 nm) used to collect electron diffraction data is larger than the size of the converged beam used in STEM experiments (~0.1 nm). Crystallographic relationships between zircon and reidite show a complex dependency of reidite from parent zircon: *c*-axes of reidite form four clusters, roughly coincident with <110>_z and <100>_z. Reidite axes show significant rotation (asterism).

Following the above-tested approach, we also mapped the more complex mineralogical composition of foil #7103, which consists of a mixture of granular zircon and irregular grains of reidite, with particles less than 1 μm in size and with a high defect density (Fig. 4; Online Materials¹ Fig. SI5). In addition, this sample contained fractures outlined by nanometer-sized *m*-ZrO₂ (baddeleyite, monoclinic), which could be identified using high-resolution HAADF imaging followed by fast Fourier transformation (FFT) analysis (Fig. 4b) and by comparing its EELS spectra with those published earlier (McComb et al. 1992;

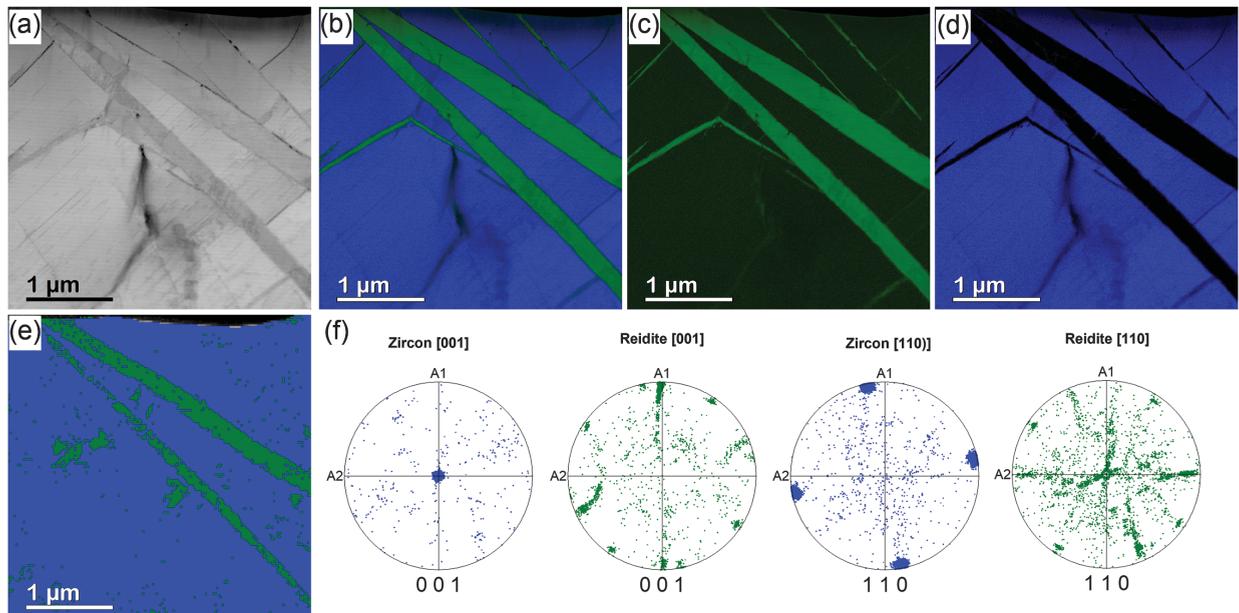


FIGURE 3. The detail marked as a dotted box in part of FIB foil #7081 shown in Figure 2a. (a) STEM image; (b) combined EELS map showing distribution of (c) reidite (green) and (d) zircon (blue); (e) combined image showing phase distribution as identified using PED; (f) corresponding pole figures of reidite and zircon for [001] and [110] orientations.

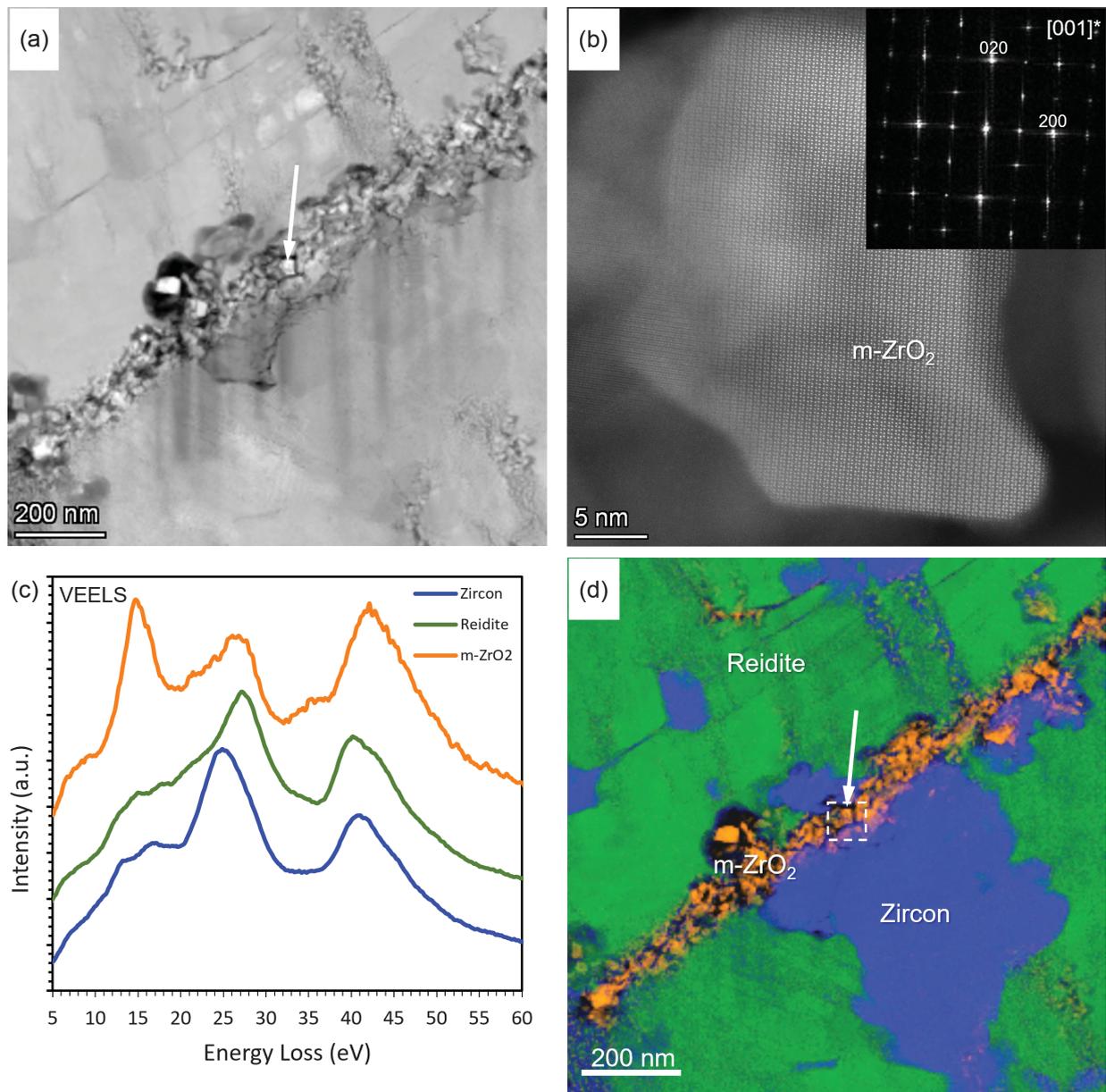


FIGURE 4. TEM results from FIB foil #7103. (a) HAADF image. A white arrow points at a single grain of $m\text{-ZrO}_2$ shown in (b); (b) HR-STEM image of $m\text{-ZrO}_2$ along [001]; (c) VEELS spectra of zircon, reidite, and the $m\text{-ZrO}_2$ phases collected over the imaged area shown in (a); (d) composite image showing the distribution of the three phases: reidite (green), zircon (blue), and $m\text{-ZrO}_2$ (orange).

Dash et al. 2004). In contrast to foil #7081, the textures and phases in foil #7103 are practically undistinguishable in the HAADF images alone. Thus, using representative spectra (Fig. 4c) of the different phases as explained above, we performed a MLLS fit to allow differentiation (Fig. 4d; Online Materials¹ Fig. S15).

Two examples of small reidite and $m\text{-ZrO}_2$ inclusions in zircon are shown in Figures 5a and 6a, respectively. The inclusions were detected by VEELS mapping (Figs. 5b–5d, 6b, and 6c). Irregular nanometer-sized grains of reidite with sharp ragged boundaries are included in dense granular zircon matrix. The reidite inclusion in Figure 5 has sharp (Fig. 5e) and blurred (Fig. 5f)

interfaces with surrounding zircon. The interface looks locally blurred due to the overlapping of both phases. It is worth noting that the true 2D shape of reidite inclusion in Figure 5a cannot be traced from the HAADF image. In contrast, the EELS maps allow for straightforward detection of such an inclusion even down to a nanometer scale.

In Figure 7, we show an image and corresponding VEELS and HL EELS spectrum of a lamellar grain with ZrSiO_4 composition but unusual radiating texture. Thin radiating lamellae 100–200 nm thick and several micrometers long have lobate boundaries, tapering tips, and pore space at the boundaries of adjacent lamellae. They are slightly misoriented from each other ($<10^\circ$). The

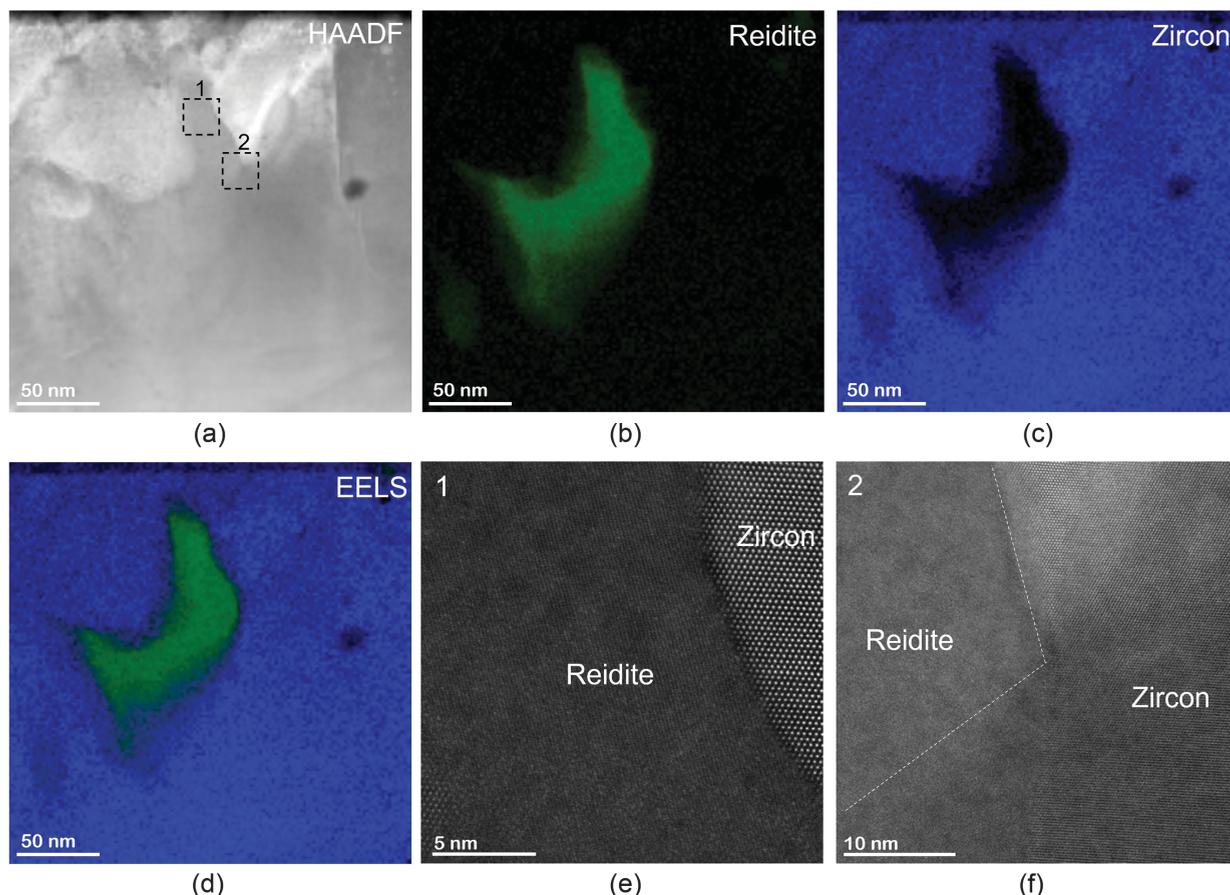


FIGURE 5. #7103 TEM specimen. (a) HAADF image of a reidite grain inside zircon matrix; corresponding phase maps of reidite (b) and zircon (c), respectively; (d) composed image of (b) and (c) maps; zoomed areas showing examples of a sharp (e) and blurred (f) reidite/zircon interfaces.

SAED was inconclusive to assign it to one of the two ZrSiO_4 polymorphs (Fig. 7a, inset). Therefore, we compared its VEELS (Fig. 7b) and O *K*-edge (Fig. 7c) spectra with those from the reference zircon (synthesized zircon), and shocked zircon and reidite materials. One can see that the VEELS and O *K*-edge spectra from the unknown aggregate differ from the spectra of zircon and reidite. While the positions of the main features in VEELS are similar to VEELS of zircon, the fine structure of the O *K*-edge is less visible. A more detailed study of crystalline structure using high-resolution STEM imaging (Fig. 7d) has revealed numerous voids of different sizes, overlapping areas between adjacent radial lamellae, and a high density of structural defects. We assume that the presence of these lattice distortions may be a reason for the lack of fine spectral features.

DISCUSSION

Interpretations of observed textures

Using the proposed TEM-based mapping method, we demonstrated intricate features of two different textural relationships between zircon and reidite, which greatly facilitate interpretations of phase transition and mineral reactions.

The texture of foil #7081 shows defined planar reidite lamellae cutting the zircon grain in four directions with sharp

boundaries without intermediate amorphous layers or other phases (Figs. 2–3; Online Materials¹ Fig. S12). The zircon blocks have the same orientation (Figs. 3f–3g), suggesting that the zircon was a single grain that pre-existed the impact event and the reidite formation. The transverse relationships of oriented reidite lamellae in blocky zircon indicate shock-induced zircon-reidite transition (e.g., Reddy et al. 2015; Erickson et al. 2017). The high dislocation density in the reidite and cracks in the host zircon indicate changes in strain and volume associated with transformation (the reidite is approximately 10% denser). All four possible orientations of reidite (Erickson et al. 2017) are visible only in the EELS phase map (Fig. 2b), demonstrating the great advantage of the proposed method.

A more complicated texture in foil #7103 shows a dense polygranular zircon mass (Fig. 4d; Online Materials¹ Fig. S15) with irregular inclusions of reidite (Fig. 5). The polygranular nature of zircon suggests that it is not original (preexisting), but recrystallized. Based on the empirical and experimental models, available in the literature, there are two scenarios by which this texture could have formed:

(1) As a result of zircon transformation to reidite and subsequent incomplete reversion from reidite back to zircon (Plan et al. 2021). Then the reidite inclusions (Online Materials¹ Fig. S15) can be considered relics and would indicate at the direction of phase

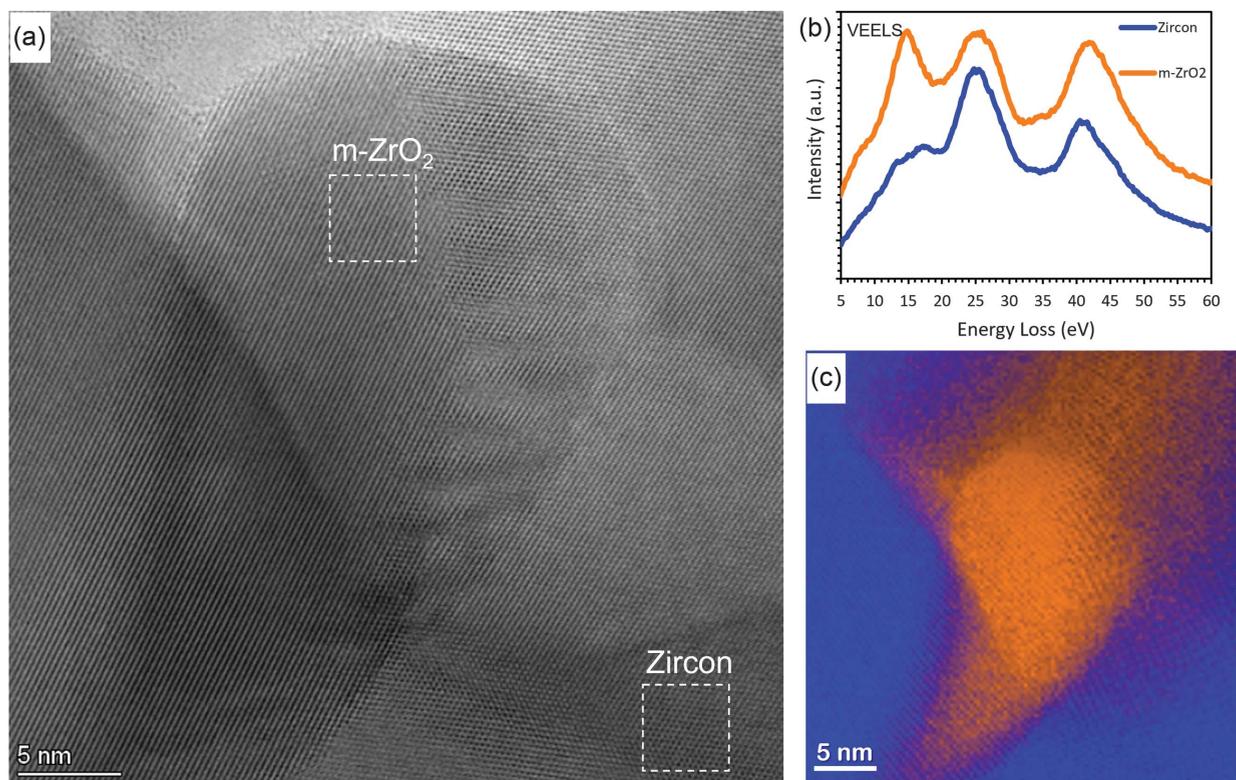


FIGURE 6. #7103 TEM specimen. (a) A bright-field high-resolution STEM image of $m\text{-ZrO}_2$ inclusion inside a small pore formed by zircon grains; (b) corresponding VEELS spectra collected from the positions marked with dashed squares in (a); (c) a corresponding phase distribution of $m\text{-ZrO}_2$ and zircon phases.

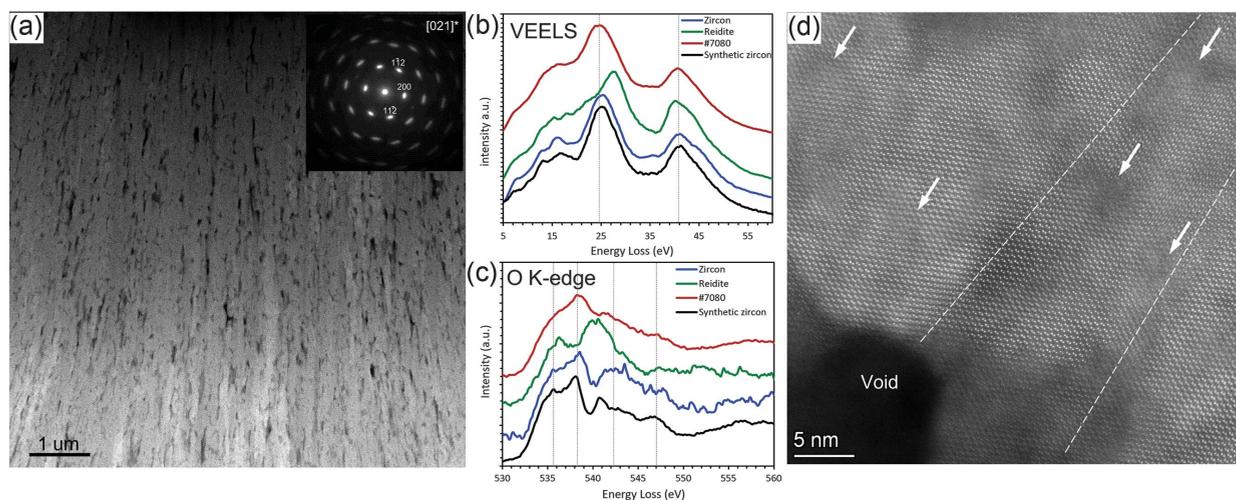


FIGURE 7. (a) HAADF image of TEM specimen #7080 and corresponding integrated nanobeam electron diffraction pattern (inset). Panels (b) and (c) correspond to VEELS and O K -edge spectra (red curves) as well as spectra of zircon (blue), reference zircon (black) and reidite (green) phases. The red curves are different from all three other phases, namely, it has no fine structure and only two broad main peaks visible. (d) HR STEM image of an area between two radiating lamellae. White dashed lines show the approximate overlapping area, which is terminated at a void; white arrows point out defects in the crystalline structure of zircon.

transition (reidite \rightarrow zircon). The scenario (1) would imply the following: (a) preexisting zircon experienced high-pressure transition to reidite, followed by (b) reidite transformation to polygranular zircon (Online Materials¹ Fig S15). Such texture

was previously described and interpreted this way by Plan et al. (2021). However, this scenario does not clarify why domains composed of recrystallized zircon with reidite inclusions do not occur in the shape of lamellae if they form after reidite. It is

possible that granular—not lamellar—reidite had developed in an amorphous zircon core, as suggested by Erickson et al. (2017).

(2) As a result of an incomplete transformation of zircon to reidite through an intermediate elastic high-pressure low-symmetry (HPLS) phase (Mihailova et al. 2019; Stangarone et al. 2019). The latter experimentally demonstrated that static compression of zircon at above 20 GPa creates a mixture of the HPLS $ZrSiO_4$ phase (intermediate phase between zircon and reidite) and proper reidite. When the pressure is released, the HPLS phase transforms back to zircon. Hence, quenched samples were composed of a mixture of reidite and zircon, similar to what we observed in sample #7103. Hence, our Kara sample could be a natural equivalent of the experimental product of Stangarone et al. (2019).

In addition, small grains of m - ZrO_2 occur in arrays along fractures in foil #7103 (Figs. 4d; Online Materials¹ Fig. S15). No silica or cubic zirconia were observed to support the high-temperature incongruent melting of $ZrSiO_4$, neither are there inclusions of ZrO_2 enclosed inside the zircon granules, opposite to the texture described by Kovaleva et al. (2021). Baddeleyite is most likely formed the latest, at the expense of $ZrSiO_4$ phases as a result of post-impact hydrothermally driven processes (Lewerentz et al. 2019).

Zircon aggregate sample #7080 with spherulitic shape and radiating lamellae of significantly damaged crystalline zircon (Fig. 7a) may be a result of:

(1) Complete (bulk) melting during shock loading (e.g., Takagi et al. 2022), followed by cooling to form radial lamellae. The position of zircon aggregate (sample KK21-21) at the boundary between a glass-rich domain and a lithic breccia domain (Online Materials¹ Fig. S11) potentially supports this interpretation, as zircon could have crystallized from local melt upon contacting a colder breccia. However, no ZrO_2 or SiO_2 inclusions were found that often accompany high-temperature reactions involving zircon melting (e.g., Kovaleva et al. 2021, 2023; Wittmann et al. 2006). It would also be difficult to bulk-melt such a comparatively large volume of zircon.

(2) Transformation to HPLS phase and back to zircon upon pressure release (Stangarone et al. 2019). The lamellar structure of this sample resembles the lamellar nature of reidite inclusions in zircon, or aggregates that are almost completely transformed to reidite [Fig. 3; see multiple examples in Plan et al. (2021)]. Hence, the observed sample might represent an incomplete transition to massive reidite, which reverted back to zircon from the non-quenchable HPLS phase.

A variety of phases and their morphologies and structural damage found in shocked zircon aggregates suggests that we deal with various fixed snapshots of phase transformations determined by thermodynamic conditions and constrained by the local crystallo-chemical environment.

Broader geochemical and cosmochemical significance

Here, reidite is reported from the Kara impact structure, adding to the list of impact craters with this rare high-pressure polymorph. The presence of reidite indicates that the maximum shock pressure was at least 30 GPa (Kusaba et al. 1985). We also demonstrate the high-resolution appearance of texture resulting

from the zircon-reidite and reidite-zircon phase transitions (Plan et al. 2021). More generally, our work provides a case study of how damaged and/or nano-sized phases can be resolved and studied with the level of details, unavailable for any other analytical methods. The presented workflow can be utilized for many other terrestrial and extraterrestrial samples, e.g., low volume samples obtained in sample return missions. Using the example of zircon-reidite aggregates, we show how phase distributions occurred due to mineral reactions and phase transitions can be studied at the atomic resolution.

IMPLICATIONS

Our data demonstrate that zircon and reidite are indeed characterized by distinct spectra in both low- and core-loss regions. This difference can be used to map their distribution in TEM specimens at very high spatial resolution and precision. Moreover, the fact that only low-loss spectra had to be collected for the maps provides another great advantage, as this means much shorter acquisition time (milliseconds vs. tens or hundreds of milliseconds for standard high-loss spectra per pixel or ED pattern per image in 4D-STEM experiments). This allows for minimizing the electron dose and thus the electron irradiation damage. Therefore, the presented EELS technique could also be applied to a variety of electron-beam-sensitive materials. Moreover, the procession time for PED data is much longer and crucially depends on the accuracy of the crystallographic information files that are used. Thus, our novel approach has great potential for phase mapping with unsurpassed spatial resolution for a series of geoscience applications.

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