

Chemistry and genetic implications of tourmaline and Li-F-Cs micas from the Valdeflores area (Cáceres, Spain)

ALFONSO PESQUERA,^{1,*} JOSE TORRES-RUIZ,² PEDRO P. GIL-CRESPO,¹ AND NICOLAS VELILLA²

¹Departamento de Mineralogía y Petrología, Universidad del País Vasco 644, 48080 Bilbao, Spain

²Departamento de Mineralogía y Petrología, Universidad de Granada, E-18002 Granada, Spain

ABSTRACT

Pervasive metasomatism that involved the formation of tourmaline-rich rocks and influx of Li, F, and Cs into Ordovician psammo-pelitic metasediments occurred in the Valdeflores area (Cáceres, Spain). Numerous Li- and Sn-bearing, mineralized, greisen-type veins also can be observed here, in the vicinity of geochemically specialized granites. Tourmaline-rich rocks appear as: (1) massive, fine-grained, dark green to black rocks; and (2) fine-scale tourmaline-rich laminae, which alternate with quartz-rich layers parallel to the bedding.

Electron microprobe analyses indicate that the tourmaline lies mostly within the space defined by the exchange vectors: FeMg_{-1} (schorl), $\square\text{AlNa}_{-1}\text{Mg}_{-1}$ (foitite), $\text{AlOMg}_{-1}(\text{OH})_{-1}$ (olenite), and $\text{CaMgNa}_{-1}\text{Al}_{-1}$ (uvite). The $\text{Fe}/(\text{Fe}+\text{Mg})$ ratio ranges mainly from 0.87 to 0.54 and increases with Al in the Y-site. Analytical results and substitutional relations show an insignificant elbaite component. Mica in the tourmalinized rocks is very fine-grained (mostly $<50\ \mu\text{m}$). White mica ranges from lithian muscovite-phengite to lepidolite/zinnwaldite, containing up to 8.40 wt% F, 6.0 wt% Li_2O , and 10.73 wt% FeO. Dark mica shows a variable color and has compositions characterized by relatively high contents of Cs_2O (1.14–2.78 wt%) and F (1.94–8.08 wt%), with a deficit in K_2O (5.75–9.04 wt%). $\log(f_{\text{H}_2\text{O}}/f_{\text{HF}})$ of fluids in equilibrium with biotite in the tourmaline-rich rocks was 4.02–4.17 at $T \approx 400\ ^\circ\text{C}$. $\log(f_{\text{H}_2\text{O}}/f_{\text{HF}})$ values of fluids in equilibrium with topaz ($X_{\text{F}} \approx 0.8$) in country rock adjacent to contacts with veins, and in equilibrium with amblygonite-montebrazite ($X_{\text{amb}} = 0.2$) in the veins were about 4.30–4.60 and 6.4–6.7, respectively. These variations denote the existence of gradients in relative a_{HF} more than differences of temperature during metasomatism. The lack of tourmaline in the veins is interpreted to reflect the alkalinity and low Fe-Mg contents in the fluids, which precluded the formation of tourmaline. Consequently, most of the boron was expelled into metasediments where tourmaline was produced as a result.