

Compositional dependence of H₂O solubility along the joins NaAlSi₃O₈-KAlSi₃O₈, NaAlSi₃O₈-LiAlSi₃O₈, and KAlSi₃O₈-LiAlSi₃O₈

C. ROMANO,¹ D.B. DINGWELL,¹ H. BEHRENS,² AND D. DOLFI³

¹Bayerisches Geoinstitut, Universität Bayreuth, D-95440 Bayreuth, Germany

²Institut für Mineralogie, Universität Hannover, Welfengarten 1, 30167 Hanover, Germany

³Dipartimento di Scienze della Terra, Università degli studi di Roma "La Sapienza," P. le Aldo Moro 1, 00185, Rome, Italy

ABSTRACT

The composition-dependence of the solubility of H₂O in silicate melts along the binary joins NaAlSi₃O₈-KAlSi₃O₈, NaAlSi₃O₈-LiAlSi₃O₈, and KAlSi₃O₈-LiAlSi₃O₈ has been determined at 2 kbar and 1040 °C. The study involved 1 atm dry oxide fusion, hydrothermal saturation, isobaric rapid quench and macroscopic analysis for H₂O using Karl Fischer titration (KFT) and microscopic investigation of homogeneity using infrared absorption spectroscopy, respectively.

The solubility of H₂O in these melts increases in the order KAlSi₃O₈ (5.12 wt%) < NaAlSi₃O₈ (6.03 wt%) < LiAlSi₃O₈ (7.32 wt%). The total relative weight percent variation in solubility is 43%, which corresponds to a relative mole percent variation of 23%. Along the joins, small but systematic deviations from additivity are observed at the 1σ uncertainty level. These deviations are positive on the NaAlSi₃O₈-KAlSi₃O₈ join but negative on the LiAlSi₃O₈-bearing joins.

The present results confirm the early suggestion of Voigt et al. (1981) that solubility variation along the NaAlSi₃O₈-KAlSi₃O₈ join is nonlinear and disagree in detail with the equimolal approximation of Burnham (1975, 1981) and Burnham and Davis (1974). Models of H₂O solubility must take into account not only the nonequimolal solubility of H₂O in feldspathic melts but also nonlinear binary variations as well. We interpret the relative solubilities of the end-member compositions in terms of the relative stability of the tectosilicate melt structure and the nonlinearities along the joins in terms of the next-nearest neighbor distributions of Al and Si in these melts.

INTRODUCTION

H₂O is one of the most abundant volatile species in magmas, and its presence strongly influences physical and chemical melt properties as well as crystal-melt phase relations. A few weight percent of dissolved H₂O drastically decreases the liquidus and solidus temperature of a melt and strongly affects the phase relations and the compositions of the minerals crystallizing from the evolving magma (Yoder 1965; Kushiro et al. 1968). The viscosity can be modified by more than five orders of magnitude (Shaw 1963; Dingwell 1987). Additionally, in the presence of H₂O, the electrical conductivity is enhanced (Lebedev and Khitarov 1964), and the diffusivities of some crystal-forming elements are increased (Watson 1979). Although partial molar volume estimates for H₂O vary (see review by Lange 1994), numerous studies have shown that the presence of H₂O decreases the density of silicate melts and glasses (e.g., Acocella et al. 1984; Silver et al. 1990; Romano et al. 1994).

Because of these consequences of the incorporation of H₂O into silicate melts, significant attempts at modeling the systematics of the dependence of the solubility of H₂O in silicate melts on melt composition, through the devel-

opment of thermodynamic models, have been performed (Burnham 1975; Burnham and Davis 1974; Spera 1974; Stolper 1982b; Silver and Stolper 1989; Silver et al. 1990) and have met with some success. One of the best constraints on thermodynamic modeling of H₂O solution in silicate melts is highly precise and accurate solubility data. Surprisingly, much too few data exist (see reviews by McMillan and Holloway 1987; McMillan 1994), despite the long history of investigation of H₂O in silicate melts (Goranson 1931, 1936, 1938; Yoder et al. 1957; Yoder 1965; Burnham and Jahns 1962; Burnham and Davis 1971; Orlova 1962; Kadik and Lebedev 1968; Boettcher and Wyllie 1969; Oxtoby and Hamilton 1978a, 1978b; Voigt et al. 1981; Day and Fenn 1982; Dingwell et al. 1984; Hamilton and Oxtoby 1986; Holtz et al. 1992, 1993, 1995). Investigations of the microscopic interactions of H₂O in melts have also been pursued in an effort to develop structural models capable of generalizing the results of solubility studies in thermodynamic terms (Silver and Stolper 1985, 1989; Mysen and Virgo 1986; McMillan and Remmele 1986; Farnan et al. 1987; Kohn et al. 1989, 1992; Kümmerlen et al. 1993).

Inspired by evidence from recent studies pointing to significant compositional dependence of the solubility of

H₂O in felsic melts (Holtz et al. 1992, 1993, 1995), and in an attempt to resolve controversy surrounding previous determinations, we turned our attention to melts of the feldspar stoichiometry (K,Na,Li)AlSi₃O₈.

In this study, we present a series of experiments specifically devoted to the determination of the compositional dependence of H₂O solubility in the systems NaAlSi₃O₈-KAlSi₃O₈, NaAlSi₃O₈-LiAlSi₃O₈, and KAlSi₃O₈-LiAlSi₃O₈ at 1040 °C and 2 kbar.

EXPERIMENTAL METHODS

Synthesis

The anhydrous starting glasses used in the present study were generated by the direct fusion of 100 g (decarbonated equivalent) batches of powder mixes of the alkali carbonates, SiO₂, and Al₂O₃. The powders of the starting materials were dried at 120 °C overnight prior to weighing, then mixed in plastic bottles by agitation for approximately 5–10 min. The powder mixes were fused in air in thinwalled 75 cm³ platinum crucibles for approximately 2 h at 1650 °C. The products of the initial fusion were bubble-rich and inhomogeneous. To remove bubbles, react fully, and homogenize the starting glasses, the melted batches were transferred to a second furnace equipped with a high temperature viscometer and stirred at 50 or 100 rpm for hours to days at 1650 °C. The melts were periodically sampled by removal of the spindle to check for the presence of bubbles and crystals. The products of these second fusions were cooled slowly in the box furnace to <800 °C, and then removed and allowed to cool slowly to room temperature to prevent fracturing during quenching.

The composition of the starting glasses was determined by ICP-AES analyses. The results of the analysis are given in Table 1 as the proportions (in weight percent) of the oxides. The analyzed compositions deviate from the ostensible compositions by <0.05 in the Al/alkali ratio and <1 wt% in SiO₂ content.

Glass cylinders (diameter: 0.3 cm; length: 0.3–0.4 cm) were drilled by water-cooled diamond coring tools, cleaned ultrasonically in acetone, and then dried at 110 °C to remove any residue of acetone. Glass samples (~0.1 g) were then loaded together with ~0.01 g of doubly distilled water in capsules formed from platinum tubing (outer diameter: 0.32 cm; length: 1.5 cm, wall thickness 0.015 cm) and sealed with an arc-welder.

The amount of added water was chosen to be about 3–4% higher than the expected solubility value. This value was high enough to prevent the effect of air enclosed in the capsule during preparation but low enough to minimize possible changes in the anhydrous stoichiometry of the melt resulting from incongruent dissolution of melt components in water.

The capsules were checked for possible leakage by testing for weight loss after drying in an oven at 100 °C for at least 1 h. After an hour at 100 °C, the added water was considered to be distributed randomly (space-filling) in

TABLE 1. Analyzed anhydrous compositions (wt% oxide)

Comp.	SiO ₂ (±0.6)	Al ₂ O ₃ (±0.4)	Na ₂ O (±0.5)	K ₂ O (±0.5)	Li ₂ O (±0.5)	Total
Na ₁₀₀	68.1	19.1	11.8			99.0
	<i>68.7</i>	<i>19.4</i>	<i>11.8</i>			
Na ₉₀ K ₁₀	69.1	19.0	10.9	1.8		101.0
	<i>68.3</i>	<i>19.3</i>	<i>10.6</i>	<i>1.8</i>		
Na ₇₀ K ₃₀	67.8	19.0	8.0	5.4		100.2
	<i>67.5</i>	<i>19.1</i>	<i>8.1</i>	<i>5.3</i>		
Na ₅₀ K ₅₀	67.6	18.9	5.6	8.7		100.8
	<i>66.7</i>	<i>18.9</i>	<i>5.7</i>	<i>8.7</i>		
Na ₃₀ K ₇₀	66.4	18.8	3.3	12.0		100.5
	<i>65.9</i>	<i>18.6</i>	<i>3.4</i>	<i>12.1</i>		
Na ₁₀ K ₉₀	65.3	18.8	1.2	15.6		100.9
	<i>65.1</i>	<i>18.4</i>	<i>1.1</i>	<i>15.3</i>		
K ₁₀₀	65.4	18.5		16.5		100.4
	<i>64.8</i>	<i>18.3</i>		<i>16.9</i>		
K ₉₀ Li ₁₀	64.6	18.8		15.8	0.5	99.7
	<i>65.6</i>	<i>18.5</i>		<i>15.4</i>	<i>0.5</i>	
K ₇₀ Li ₃₀	66.3	19.2		12.6	1.6	99.7
	<i>67.0</i>	<i>19.0</i>		<i>12.3</i>	<i>1.7</i>	
K ₅₀ Li ₅₀	68.2	20.0		9.3	2.9	100.4
	<i>68.7</i>	<i>19.4</i>		<i>9.0</i>	<i>2.9</i>	
K ₃₀ Li ₇₀	70.3	20.6		5.8	4.2	100.9
	<i>70.5</i>	<i>19.9</i>		<i>5.5</i>	<i>4.1</i>	
K ₁₀ Li ₉₀	72.2	20.9		1.9	5.5	100.5
	<i>72.3</i>	<i>20.4</i>		<i>1.9</i>	<i>5.4</i>	
Li ₁₀₀	73.7	20.7			6.2	100.6
	<i>73.2</i>	<i>20.7</i>			<i>6.1</i>	
Li ₉₀ Na ₁₀	73.3	20.8	1.1		5.1	100.3
	<i>72.8</i>	<i>20.6</i>	<i>1.3</i>		<i>5.4</i>	
Li ₇₀ Na ₃₀	72.2	20.5	3.6		3.5	99.8
	<i>71.8</i>	<i>20.3</i>	<i>3.7</i>		<i>4.2</i>	
Li ₅₀ Na ₅₀	70.8	20.2	6.1		2.3	99.4
	<i>70.9</i>	<i>20.1</i>	<i>6.1</i>		<i>2.9</i>	
Li ₃₀ Na ₇₀	69.6	20.4	8.4		1.1	99.5
	<i>70.0</i>	<i>19.8</i>	<i>8.4</i>		<i>1.7</i>	
Li ₁₀ Na ₉₀	68.2	19.5	10.7		0.6	99.0
	<i>69.2</i>	<i>19.6</i>	<i>10.7</i>		<i>0.6</i>	

Note: Na₁₀₀ = NaAlSi₃O₈ (albite), K₁₀₀ = KAlSi₃O₈ (orthoclase). Compositions determined by ICP analysis (nominal compositions in italics).

the platinum capsule, and thus around the glass sample. The sealed capsules were placed in TZM vessels (argon pressure medium), pressurized to 2 kbar and heated to 1040 °C where they were held for a time sufficient to allow complete homogenization of H₂O dissolved in the melt by diffusion through the sample (experimental durations ranged from 70 to 216 h. See Table 2). Temperature was measured with a Ni-NiCr thermocouple (accuracy ±15 °C) and the pressure was measured with a strain-gauge manometer (accuracy ±0.05 kbar). After the high pressure–high temperature dwells, the samples were quenched rapidly and isobarically in the TZM vessels by dropping the sample into the cold part of the vessel (estimated quench rate 200 °C/s). Special care was taken during the quench to maintain isobaric conditions by opening the vessel to the pressure line (2±0.05 kbar).

Synthesis of hydrous glasses was also performed in the Institut für Mineralogie, Universität Hannover, for some compositions along the NaAlSi₃O₈-KAlSi₃O₈ join. Dry glasses for Ab and Or end-members were prepared according to the procedure described in Behrens (1995). For intermediate compositions, the starting materials were hydrous glasses with approximately 2 wt% of H₂O. The

TABLE 2. Experimental results

Comp.	Sample	t (h)	H ₂ O _{ex} * (wt%)	H ₂ O (wt%)	H ₂ O** (mol fr.)	Comp.	Sample	t (h)	H ₂ O _{ex} * (wt%)	H ₂ O (wt%)	H ₂ O** (mole fr.)	Comp.	Sample	t (h)	H ₂ O _{ex} * (wt%)	H ₂ O (wt%)	H ₂ O** (mole fr.)
K ₁₀₀	341-1	84	8.19	5.04	0.170	K ₅₀ Na ₅₀	Han-1†	70	6.30	5.61	0.182	K ₃₀ Li ₇₀	335-1	88	9.39	6.40	0.195
K ₁₀₀	184-1	74	6.39	5.20	0.175	K ₄₀ Na ₆₀	Han-1†	70	10.30	5.84	0.188	K ₃₀ Li ₇₀	335-1	88	9.39	6.43	0.196
K ₁₀₀	Han-1†‡	70	8.50	5.09	0.172	K ₄₀ Na ₆₀	Han-1†	70	7.60	5.86	0.188	K ₁₀ Li ₉₀	336-1	92	9.17	6.96	0.206
K ₉₀ Na ₁₀	95-1	72	9.50	5.26	0.176	K ₃₀ Na ₇₀	Rq-1	82	9.90	6.03	0.192	K ₁₀ Li ₉₀	336-1	92	9.17	6.94	0.205
K ₉₀ Na ₁₀	95-4	72	9.50	5.28	0.176	K ₃₀ Na ₇₀	Rq-1	82	9.90	5.96	0.190	Li ₁₀₀	214-2	88	15.12	7.52	0.217
K ₈₀ Na ₂₀	Han-1†	70	7.70	5.36	0.178	K ₃₀ Na ₇₀	413-1	84	11.93	5.83	0.187	Li ₁₀₀	214-2	88	15.12	7.37	0.214
K ₇₀ Na ₃₀	326-1	100	9.06	5.58	0.183	K ₁₀ Na ₉₀	394-1	144	8.90	6.05	0.191	Li ₁₀₀	214-1	88	15.12	7.19	0.209
K ₇₀ Na ₃₀	326-1	100	9.06	5.70	0.186	K ₁₀ Na ₉₀	394-3	144	8.90	6.00	0.189	Li ₁₀₀	214-4	88	15.12	7.19	0.209
K ₇₀ Na ₃₀	103-1	96	9.36	5.45	0.180	K ₁₀ Na ₉₀	412-1	84	10.00	5.99	0.189	Na ₉₀ Li ₁₀	339-1	82	8.81	6.11	0.191
K ₇₀ Na ₃₀	103-4	96	9.36	5.46	0.180	K ₁₀ Na ₉₀	412-1	84	10.00	6.00	0.189	Na ₉₀ Li ₁₀	339-3	82	8.81	6.13	0.191
K ₇₀ Na ₃₀	359-1	70	8.89	5.41	0.178	Na ₁₀₀	325-1	72	8.50	5.99	0.188	Na ₇₀ Li ₃₀	200-1	70	12.05	6.38	0.196
K ₇₀ Na ₃₀	103-4	96	9.36	5.51	0.181	Na ₁₀₀	Rq-1	216	9.30	5.99	0.188	Na ₇₀ Li ₃₀	200-1	70	12.05	6.20	0.191
K ₇₀ Na ₃₀	415-1	84	10.25	5.48	0.180	Na ₁₀₀	176-1	108	9.28	6.10	0.191	Na ₇₀ Li ₃₀	200-3	70	12.05	6.27	0.193
K ₇₀ Na ₃₀	415-1	84	10.25	5.53	0.182	Na ₁₀₀	176-1	108	9.28	6.05	0.190	Na ₅₀ Li ₅₀	334-1	82	9.39	6.38	0.194
K ₇₀ Na ₃₀	415-1	84	10.25	5.50	0.181	Na ₁₀₀	Han-1†‡	70	10.00	6.05	0.190	Na ₅₀ Li ₅₀	334-1	82	9.39	6.58	0.199
K ₇₀ Na ₃₀	98-1	104	10.55	5.48	0.180	K ₉₀ Li ₁₀	338-1	84	8.51	5.36	0.177	Na ₅₀ Li ₅₀	203-1	70	15.30	6.39	0.194
K ₅₀ Na ₅₀	149-2	72	9.20	5.80	0.188	K ₉₀ Li ₁₀	338-3	84	8.51	5.27	0.175	Na ₅₀ Li ₅₀	203-1	70	15.30	6.40	0.194
K ₅₀ Na ₅₀	149-1	72	9.20	5.83	0.188	K ₇₀ Li ₃₀	333-1	92	8.24	5.45	0.177	Na ₅₀ Li ₅₀	203-4	70	15.30	6.60	0.200
K ₅₀ Na ₅₀	149-4	72	9.20	5.83	0.188	K ₇₀ Li ₃₀	333-1	92	8.24	5.65	0.183	Na ₅₀ Li ₅₀	203-4	70	15.30	6.42	0.195
K ₅₀ Na ₅₀	358-1	72	9.68	5.56	0.181	K ₇₀ Li ₃₀	333-3	92	8.24	5.60	0.181	Na ₃₀ Li ₇₀	206-2	74	14.42	6.89	0.205
K ₅₀ Na ₅₀	358-1	72	9.68	5.58	0.181	K ₇₀ Li ₃₀	213-1	146	13.12	5.62	0.182	Na ₃₀ Li ₇₀	206-1	74	14.42	6.82	0.203
K ₅₀ Na ₅₀	414-1	84	10.00	5.58	0.181	K ₅₀ Li ₅₀	185-2	108	11.13	5.82	0.184	Na ₃₀ Li ₇₀	206-4	74	14.42	6.72	0.201
K ₅₀ Na ₅₀	414-1	84	10.00	5.55	0.181	K ₅₀ Li ₅₀	185-2	108	11.13	5.91	0.186	Na ₃₀ Li ₇₀	206-4	74	14.42	6.79	0.202
K ₅₀ Na ₅₀	414-1	84	10.00	5.72	0.186	K ₅₀ Li ₅₀	185-2	108	11.13	6.05	0.190	Na ₁₀ Li ₉₀	340-1	80	9.66	7.15	0.210
K ₅₀ Na ₅₀	Han-1†	70	11.40	5.77	0.187	K ₅₀ Li ₅₀	185-1	108	11.10	6.04	0.190	Na ₁₀ Li ₉₀	340-3	80	9.66	7.12	0.210

Note: Compositions are of anhydrous glasses. Sample numbers are representative of single experiments. Each experiment has been measured by KFT for H₂O determination at different times to observe possible variation of solubility values due to absorption processes, dissolution processes, or both (see text); 1 = first determination, few hours to few days from the quench of the experiments, 2 = 3 months after the quench, 3 = 6 months after the quench, and 4 = 9 months after the quench.

* Excess of H₂O in the capsule during the synthesis of the wet glasses.

** Mole fraction of H₂O calculated on 2 O atom basis.

† Experiments performed in the Institut für Mineralogie, Universität Hannover, Germany.

‡ The value is referred to an average of several experiments.

hydrous glasses were prepared from gels by fusion in an internally heated gas pressure vessel for more than 20 h at 1100 °C and 2 kbar (H₂O undersaturated conditions).

The hydrous synthesis was performed in a horizontal gas pressure vessel (argon as pressure medium) at $P = 2$ kbar (accuracy ± 0.02 kbar) and $T = 1100$ °C (accuracy ± 10 °C).

Composition of the starting glasses was determined by microprobe analyses. The results of the analysis are given in Table 3 as the proportions (in weight percent) of the oxides.

Karl Fischer titration (KFT)

The H₂O contents of the quenched glasses were determined by KFT. This analytical procedure has been es-

TABLE 3. Analyzed anhydrous compositions (wt% oxide)

Comp.	SiO ₂	Al ₂ O ₃	Na ₂ O	K ₂ O	Total
Na ₁₀₀	69.19	18.83	11.96	0.02	100.00
Na ₆₀ K ₄₀	66.54	18.78	7.57	7.12	100.01
Na ₅₀ K ₅₀	67.00	18.59	5.86	8.55	100.00
Na ₂₀ K ₈₀	65.65	18.52	2.24	13.59	100.00
K ₁₀₀	66.02	18.18	0.02	15.76	99.98

Note: Na₁₀₀ = NaAlSi₃O₈ (albite), K₁₀₀ = KAlSi₃O₈ (orthoclase). Compositions determined by microprobe analysis (CAMECA CAMEBAX MB Hannover) with a defocused beam (20–30 μm). Five to ten measurements were collected for each composition.

tablished to be a powerful method for determination of H₂O contents in mineral and glasses in several recent studies (Turek et al. 1976; Westrich 1987; Holtz et al. 1992, 1993, 1995; Behrens 1995; Behrens et al. 1996). The method is based on a titration involving the reaction of H₂O, in presence of specific reagents, to generate the I⁻ ion whose concentration is determined coulometrically. The coulometer used for this study is a Mitsubishi CA 05. This technique has been successfully applied to determine relatively high H₂O contents in silicate materials (up to 5 wt%, e.g., Turek et al. 1976) as well as low H₂O contents (down to 0.1 wt% H₂O, e.g., Westrich 1987). An advantage of the KFT is that only small amounts of glass are necessary to obtain reliable measurements (typically 10–20 mg for substances containing ~6 wt% H₂O). A schematic illustration of the apparatus is provided in Figure 1. The reader is referred to Behrens (1995) for further details.

In this study single chips have been used for analysis with the added precaution against sputtering that single specimens were wrapped in platinum foil. For several glasses that did not undergo sputtering, either with or without the foil, KFT yielded essentially identical results.

Measurements were performed both within a few hours after the experimental quench and after periods of several months in air. Again, no detectable difference was observed in the determined H₂O contents, confirming that

single chips of the investigated samples do not lose or gain H₂O as a function of time (Table 2).

The glasses were heated progressively from 20 to 1300 °C using a radio frequency induction furnace (Fa. Linn, HTG 1000/1,3). The temperature was measured with a Pt-Pt₉₀Rh₁₀ thermocouple, 3 mm below the sample. The chosen heating procedure depended on the anticipated H₂O concentration in the samples. The heating program was chosen to give a relatively low heating rate (50 °C/min) below 100 °C to remove and analyze the H₂O adsorbed on the surface of the sample; subsequently, a higher heating rate was employed (100 °C/min) up to the onset of the maximum extraction rate and was followed by a lower heating rate (50 °C/min) until the final temperature was reached. The heating rate must be kept reasonably low during the dehydration of the glass to avoid sputtering of the sample, as well as to allow a complete liberation of H₂O within the time of the measurement. An example of a heating program for a glass of albitic composition is: 50 °C/min from 20 to 100 °C, 100 °C/min from 100 to 400 °C, 50 °C/min from 400 to 800 °C, and finally 100 °C/min from 800 to 1300 °C. The liberated H₂O is transported by a dried Ar gas stream in the titration cell and collected there. Titration starts 4–6 min after the start of the heating experiment. The duration of the titration is 7–10 min.

The analytical precision of the H₂O analysis depends mostly on the duration of the titration. For a typical measurement time of 12 min and a sample mass of 0.01 g the maximum uncertainty is about 0.15 wt%. Duplicate analyses on the same samples indicate that accuracy may be higher than this (Holtz et al. 1992, 1993, 1995; Behrens 1995).

An important point in any bulk extraction technique, including the KFT determination, is the complete, reproducible extraction of H₂O during the dehydration process. Behrens (1995) has demonstrated that a residual H₂O content of 0.1 ± 0.05 wt%, regardless of the duration of the titration or of the composition of the starting material, remains in all glasses after the extractions under the conditions used in the present study. The H₂O solubility data determined by KFT have accordingly been corrected for this H₂O.

RESULTS

The quenched experimental products consisted of bubble- and crystal-free limpid glasses. The homogeneity of the glasses was investigated in two ways. Samples were cut into two slices, perpendicular to the capsule elongation. One slice was cut again into two or three pieces (depending on the total length of the original material), each was analyzed individually by KFT. The other half of the samples was then analyzed with an infrared spectrometer (Bruker IFS 120 HR). Several measurements were performed along perpendicular and parallel sections and showed no differences in the background-subtracted peak area at ~ 4500 cm⁻¹ (OH groups) and at ~ 5230 cm⁻¹ (molecular H₂O), indicating that H₂O was homo-

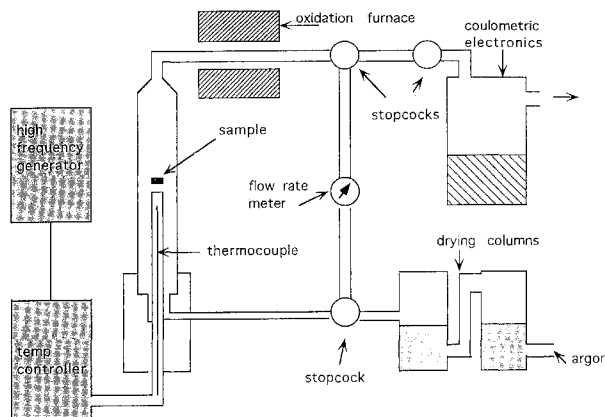


FIGURE 1. Scheme of the apparatus used for the Karl Fischer titration. Argon is used as the gas medium to transport the water liberated in the induction furnace to the titration cell. The gas is dried by two columns containing granulated phosphorus pentoxide (Fa. Baker, Granusic A).

geneously distributed throughout the samples and that the experimental durations were sufficient to produce complete hydration and equilibration of the samples.

The reproducibility of the hydration experiments was tested by multiple syntheses and analyses. The maximum difference between two analyzed H₂O contents has been found to be 0.3 wt% H₂O (Table 2), corresponding to a maximum variation of 6.0% relative to the solubility value. The multiple determinations of H₂O content of separate experimental products for the most intensively investigated composition here (K₇₀Na₃₀) yields a mean of 5.51 wt% and a standard deviation of ± 0.08 wt% ($n = 10$).

The solubility data for H₂O in melts along the joins of alkali feldspar stoichiometry NaAlSi₃O₈-KAlSi₃O₈, NaAlSi₃O₈-LiAlSi₃O₈, and LiAlSi₃O₈-KAlSi₃O₈ determined at 2 kbar and 1040 °C are given in Table 2. The H₂O solubility values (expressed as weight percent H₂O) are plotted in Figure 2. The H₂O solubility values expressed as mole percent (two O atom basis) are shown in Figure 3. H₂O solubility is dependent on anhydrous melt composition; the total H₂O content decreases in the order LiAlSi₃O₈ (7.32 wt% H₂O) > NaAlSi₃O₈ (6.03 wt% H₂O) > KAlSi₃O₈ (5.12 wt% H₂O).

Hydration experiments were also performed in the Institut für Mineralogie, Universität Hannover, for some compositions on the NaAlSi₃O₈-KAlSi₃O₈ join (Table 2; Figs. 2a and 3a). Both sets of data show excellent agreement. This equivalence of data, obtained for different materials synthesized in different laboratories and with different starting materials, confirms the validity of the KFT as method for highly accurate and reproducible H₂O solubility determinations.

The present data indicate a 23% relative increase in the molar solubility from KAlSi₃O₈ to LiAlSi₃O₈. The variation in solubility along the binary joins NaAlSi₃O₈-

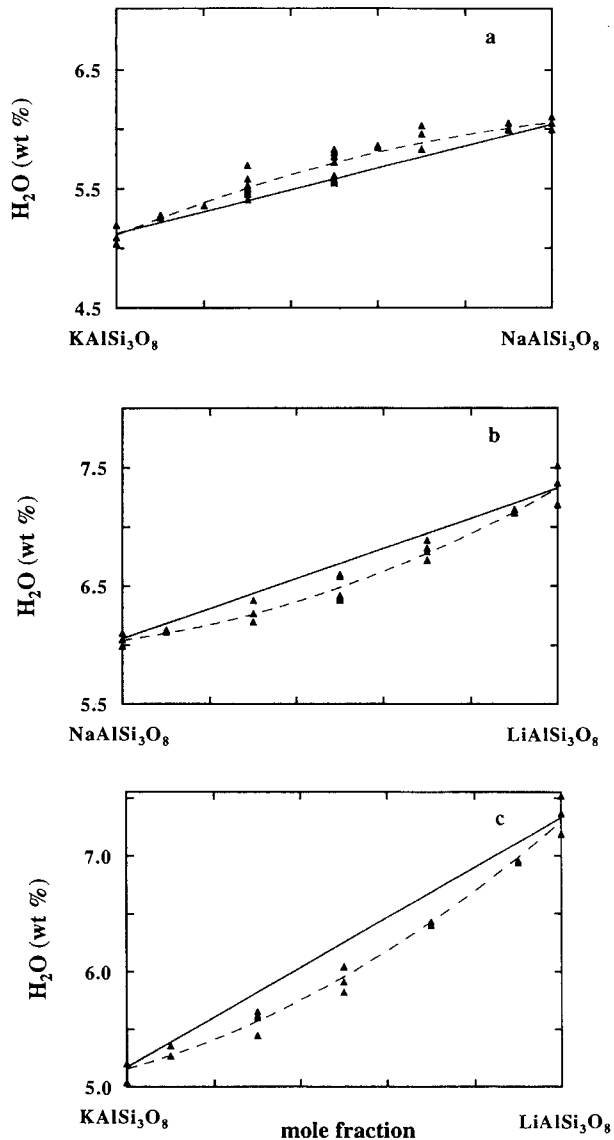


FIGURE 2. Composition dependence of H₂O solubility in (a) NaAlSi₃O₈-KAlSi₃O₈, (b) NaAlSi₃O₈-LiAlSi₃O₈, and (c) LiAlSi₃O₈-KAlSi₃O₈. The solubilities are expressed in weight percent.

KAlSi₃O₈, NaAlSi₃O₈-LiAlSi₃O₈, and KAlSi₃O₈-LiAlSi₃O₈ appears to be a nonlinear function of the molar composition. The deviation from linearity along the Na-K join is positive, whereas that along the NaAlSi₃O₈-LiAlSi₃O₈ and KAlSi₃O₈-LiAlSi₃O₈ joins is negative.

DISCUSSION

Comparison with previous work

Although the acquisition of H₂O solubility data for aluminosilicate melts of silicic composition has been the focus of several studies, the present data are directly comparable with relatively few of them. Comparable H₂O

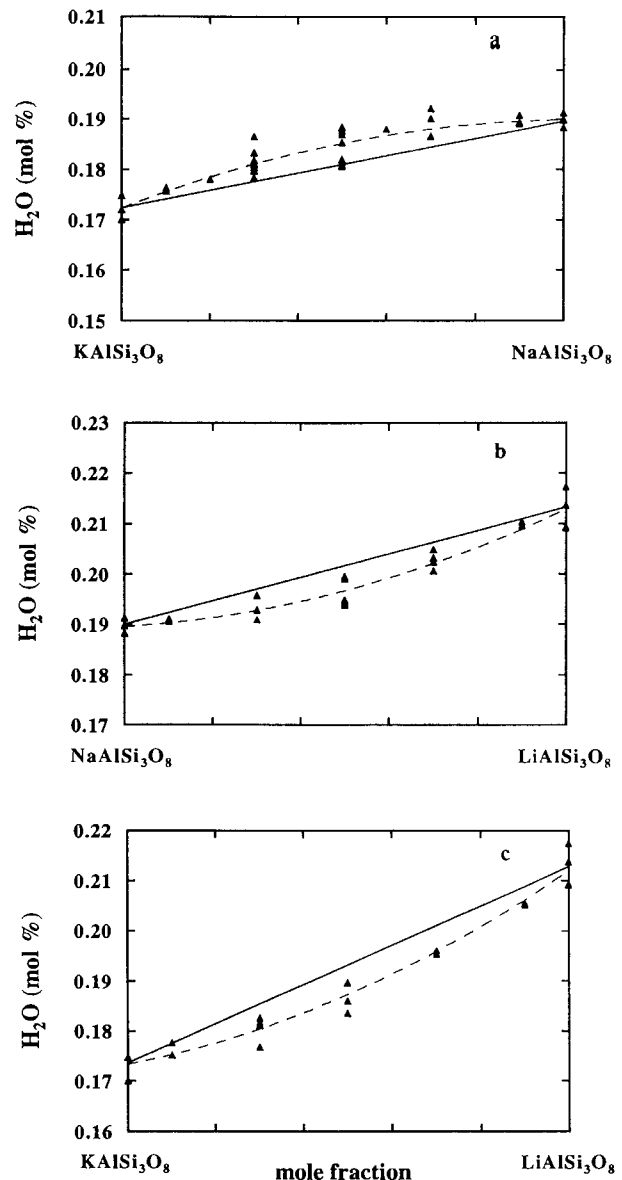


FIGURE 3. Molar H₂O solubilities (two O atom basis). (a) NaAlSi₃O₈-KAlSi₃O₈, (b) NaAlSi₃O₈-LiAlSi₃O₈, and (c) LiAlSi₃O₈-KAlSi₃O₈.

solubility measurements include those previously conducted on end-members and intermediate compositions of the NaAlSi₃O₈-KAlSi₃O₈ system.

The albite composition melt has been intensively investigated in the past 30 years (Goranson 1936; Tuttle and Bowen 1958; Burnham and Jahns 1962; Kadik and Lebedev 1968; Oxtoby and Hamilton 1978a, 1978b; Sorapure 1981; Voigt et al. 1981; Day and Fenn 1982; Blamart et al. 1986; Hamilton and Oxtoby 1986; Richet et al. 1986; Paillat et al. 1992; Silver et al. 1990; see reviews by Dingwell 1986; McMillan and Holloway 1987; McMillan 1994). The H₂O solubility data obtained in the

present study for both albite and orthoclase compositions show marked differences from data reported in previous investigations. In the case of the albitic composition, most previous studies reported higher solubility values (Hamilton and Oxtoby 1986: 1000 °C, H₂O = 6.91 wt%; 1100 °C, H₂O = 6.52 wt%; Orlova 1962: 1125 °C, H₂O = 6.45 wt%). This difference might be partly explained by the difference in the physical state of the starting materials. The starting materials used in most previous studies were fine- or coarse-grained glass powders. The use of powder as a starting material can lead to an overestimation of the H₂O content because of the presence of inclusions or bubbles in the experimental products, corresponding to undissolved H₂O at the pressure and temperature conditions of synthesis. A retrograde dissolution of H₂O into the melt during the quench (for the low pressure case of negative temperature dependence of H₂O, Paillat et al. 1992; Holtz et al. 1995) could also be responsible for the observed differences.

The retrograde solubility of H₂O in albitic melt at low pressure, when applied to the solubility data of Kadik and Lebedev (1968; 1200 °C, H₂O = 6.6 wt%), Goranson (1938; 1200 °C, H₂O = 6.32 wt%), and Paillat et al. (1992; 1400 °C, H₂O = 6.01 wt%), also generates (for the *T* conditions of our experiments) higher solubilities than those reported by the present study. For the case of the Goranson (1938) and Kadik and Lebedev (1968) data, the use of powder or crushed glass as starting material could have yielded apparently higher solubility values.

Day and Fenn (1982) obtained lower solubilities than those reported in this study (2.5 kbar, 1000 °C, 5.3–5.9 wt%). The discrepancy could be due in this case to the method employed for the determination of H₂O solubility (weight loss), which tends to underestimate the total H₂O content dissolved in the melt. The criteria for the detection of H₂O saturation used by the authors are the presence of two-phase bubbles in the glass (quench bubbles) and the presence of large bubbles, the latter interpreted as precursors of the indentation process (dimples on the surface). The presence of those bubbles was interpreted to indicate the attainment of two-phase equilibrium (fluid-melt). However, bubbles in the interior of glasses are not an unequivocal indication of saturation of H₂O at high pressure. In fact, in the low pressure experiments, inclusion bubbles formed by using powder as starting material could be responsible for the two-phase equilibrium observed, even in undersaturated conditions (Behrens 1995). In the case of the Day and Fenn (1982) study, the duration of the experiments probably was not long enough to remove such bubbles from the original material, thus explaining the low values reported by the authors.

The present data for orthoclase composition melt indicate, in contrast, a slightly higher solubility of H₂O compared with the data of Oxtoby and Hamilton (1978a; 4.4 wt% H₂O, *P* = 2 kbar, *T* = 1000 °C). This discrepancy could be due to the analytical procedure (a weight-loss technique) adopted by those authors to determine the H₂O

content of the glasses. In the weight-loss technique, H₂O concentration is obtained by comparison of the weight of H₂O added to the charge before the experiment with the weight of the postexperimental hydrous glass after the charge has been pierced and heated to 110 °C. This technique tends to underestimate the H₂O solubility, especially for high H₂O contents, because of the loss of H₂O during drying at 110 °C. The preferential dissolution of H₂O in albite-rich melts vs. orthoclase-rich melts has been consistently observed (Oxtoby and Hamilton 1978b; McMillan and Holloway 1987; Holtz et al. 1992).

The variation of the solubility of H₂O along the join NaAlSi₃O₈-KAlSi₃O₈ itself has been more controversial. Early in the development of the equimolar solubility model of Burnham (1975) and Burnham and Davis (1974), the solubilities of H₂O in albite and orthoclase melts were assumed to be equimolar. According to Voigt et al. (1981), however, the molar H₂O solubility remains almost constant between Ab and Ab₂₅Or₇₅ compositions and then drops to almost half the value obtained for albite melt in the case of orthoclase melts. Such a large drop in solubility near orthoclase composition is not seen in the present study.

There are several possible explanations for the difference between the present data and those of Voigt et al. (1981). Those authors considered the presence of dimples on the surface of the quenched melts as an indication of excess H₂O in their experiments. However, the dimple technique cannot be applied easily to determine H₂O solubility. H₂O dimples can form on the surface of the melts even if there is insufficient H₂O to saturate the whole melt, leading to an erroneous interpretation of the saturation conditions of the experiment. Moreover, at pressures below 2 kbar, the L = L + V technique requires the careful addition of small amounts of H₂O to experimental charges and the execution of H₂O-undersaturated experiments on melts for which equilibrium is difficult to achieve.

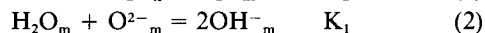
No H₂O solubility studies have been conducted on Li-bearing melts, with compositions similar to those investigated in this study. Burnham and Jahns (1962) reported solubility data for the Harding pegmatite, which contains significant concentrations of Li and F; however, the peraluminous character of the pegmatite (normative corundum = 2.4 wt%), as well as the high silica content (quartz normative = 45%), precludes a direct comparison between those results and the data presented here.

Structural interpretation

The experimental results presented here provide important constraints on the solubility mechanisms of H₂O in feldspathic melts. The changes in H₂O solubility reported indicate: (1) an increase of H₂O solubility as a function of the cation charge-balancing, in the order K < Na < Li and (2) a nonlinear variation of solubility along the NaAlSi₃O₈-LiAlSi₃O₈, KAlSi₃O₈-LiAlSi₃O₈, and NaAlSi₃O₈-KAlSi₃O₈ joins. The deviation is positive for the NaAlSi₃O₈-KAlSi₃O₈ join and negative for the

NaAlSi₃O₈-LiAlSi₃O₈ and KAlSi₃O₈-LiAlSi₃O₈ joins. These data suggest that the effect of exchanging the alkali cation in aluminosilicate melt strongly influences the H₂O solubility and hence the H₂O dissolution mechanism.

It has been established from several spectroscopic studies (Stolper 1982a, 1982b; Silver and Stolper 1985, 1989; Silver et al. 1990) that water dissolves in aluminosilicate melts both as molecular H₂O and as OH groups, according to the equilibria (Stolper 1982a, 1982b)



where m = melt and fl = fluid.

The equilibrium constants K_1 and K_2 have been demonstrated to be a function of T , P , and composition (Stolper 1982a, 1982b, 1989; Silver and Stolper 1985, 1989; Dingwell and Webb 1990; Silver et al. 1990; Zhang et al. 1991; Romano et al. 1995). Differences in H₂O solubility reported in our investigation can thus in general reflect changes in the OH or H₂O concentration in the melt.

Changes in H₂O solubility might be expected to arise from variable reactivities of O atoms with H₂O molecules. Thus in any model of significant chemical interaction between dissolved H₂O and the solvent silicate melt matrix, one factor governing the dissolution of H₂O and therefore the total amount of H₂O dissolved in the melt is the chemical bonding of O atoms in the aluminosilicate structure. Changes in O bonding environments related to the exchange of alkalis can be expected to be reflected in the relative stability of the aluminosilicate framework with differing alkali cations serving as network stabilizing cations.

The stabilization of the T-O-T framework linkage has long been inferred to be a function of the strength of the bond between the bridging O atom and the charge-balancing alkali cation (see discussion in Pichavant et al. 1992). On the basis of calorimetrically obtained enthalpies of solution and molecular orbital calculations, Geisinger et al. (1985) and Navrotsky et al. (1985) observed a positive correlation between the extent of perturbation of the T-O bond length and the z/r^2 of the charge-balancing cation at a given Al/(Al + Si). The perturbation, expressed as a narrowing of the T-O-T angle and a lengthening of the bridging T-O bonds, increases in the order K < Na < Li. Strengthening of the metal-oxygen (M-O) bond with increasing atomic number of M along a row of the periodic table is assessed to be the primary cause of weakening and lengthening of the T-O bonds (Navrotsky et al. 1985).

EXAFS studies along the NaAlSi₃O₈-KAlSi₃O₈ join (Brown et al. 1988) have demonstrated that the K-O bond is longer (3.06 Å) and has a lower field strength than the Na-O bond in the liquid (bond length = 2.6 Å). Recent neutron diffraction studies (Zotov et al. 1995) report that the average T-O₁ distances (where for O₁ is intended the next-nearest-O atom to the Si or Al) increase while the average O-O₁ distances decrease along the KAlSi₃O₈-NaAlSi₃O₈ join, which also suggests a strengthening of

the M-O bond and a weakening of T-O bond length with increasing Na content in that system.

Observations of the macroscopic physical and chemical properties of these systems (including phase equilibria, heat of solution, viscosity, and immiscible melt chemistry) support these arguments (Hess and Wood 1982; Urbain et al. 1982; DeYoreo et al. 1990; Pichavant et al. 1992).

According to all these data, the T-O-T framework linkage is more destabilized in presence of Li than in presence of Na and K, and this is consistent with a higher reactivity toward hydration in aluminosilicate melts. The relation between degree of stabilization of MAIO₂ component and the polarizing power of charge-balancing cation in the liquid would therefore appear to be a reasonable explanation for the observed variation in solubility for the end-members compositions NaAlSi₃O₈, KAlSi₃O₈, and LiAlSi₃O₈.

The observed deviations from linearity of the solubility of H₂O along the binary join NaAlSi₃O₈-KAlSi₃O₈, LiAlSi₃O₈-KAlSi₃O₈, and NaAlSi₃O₈-LiAlSi₃O₈ also require explanation.

The nonlinear variation of the physical properties of glasses when a given cation is partially substituted by a second one, generally referred to as the mixed alkali effect, has been known in silicate glasses for many years (Isard 1969; Hayward 1976; Day 1976; Dietzel 1983) and is still an active area of research (see review in Scholze 1988). Most observations of the mixed alkali effect concern transport properties (e.g., viscosity, electrical conductivity, cationic diffusivities). Nonlinear behavior has also been observed for properties in the KAlSi₃O₈-NaAlSi₃O₈ system including density, refractive index, atomic polarizability, electrical conductivity, and enthalpy of solution (Klonkowski 1983; Hayward 1976; Hervig and Navrotsky 1984; Hovis 1984; Rogez et al. 1983; Fraser et al. 1983, 1985; Fraser and Bottinga 1985; Rammensee and Fraser 1982, 1987). Also, similar deviations from linearity of the electrical conductivity along joins in the system Na₂O-Li₂O-Al₂O₃-SiO₂ have been reported by Lapp and Shelby (1987).

Despite these investigations, a model that coherently explains the mixed alkali effect has yet to be presented. Nevertheless, some useful information can be derived from spectroscopic studies of these glasses. Recent neutron diffraction studies performed on NaAlSi₃O₈-KAlSi₃O₈ glasses (Zotov et al. 1995) have shown a nonlinear decrease of the width of the first sharp diffraction peak (FSDP) which the authors attribute to a decrease of the medium-range order (MRO) with increasing mole fraction of K₂O. The change in MRO with increasing Na/(Na + K) has been related to increasing Si-Al order in the glassy network upon substitution of Na for K.

Raman spectroscopic studies of KAlSi₃O₈ and NaAlSi₃O₈ glasses (McMillan et al. 1982) indicate that Na has stronger preference for the formation of Q⁴ (2Al) rather than Q⁴ (1Al) groups. The increase of the ²⁹Si MAS NMR peak width in feldspar glasses with increasing Na/

(Na + K) ratio can also be interpreted (Oestrike et al. 1987) in terms of an increase of the relative abundances of different Q⁴ (nAl) species.

The relative abundance of the Si-O-Si, Si-O-Al, and Al-O-Al groups can affect the stability of the overall structure. It has been demonstrated on the basis of thermodynamic measurements of heats of solutions as well as molecular orbital calculations (Navrotsky et al. 1985; DeJong and Brown 1980), that Si-O-Al groups should be more stable than Si-O-Si and Al-O-Al because of polarization of the bridged O atom, and hence reduced O competition between the tetrahedral cations.

Moreover, it has also been observed that in the Al-O-Al groups, the average T-O bond length increases and the T-O-T angle decreases compared to Si-O-Si and Si-O-Al groups (Al-O has been predicted to be 1.72 Å, whereas Si-O is known to be 1.63 Å). Given the inverse relationship between bond length and strength, the Al-O-Al bonds are weaker than Si-O-Si and Si-O-Al bonds and thus potentially more reactive with H₂O.

From the comparison of Figures 2a and 4 (from Zotov et al. 1995) a direct proportionality between the solubility of H₂O and the Si-Al correlation length order can be inferred. We therefore suggest that in the system KAlSi₃O₈-NaAlSi₃O₈ the solubility of H₂O is directly related to the stability of the structure by the stability of MAI₂O₂ component (responsible for the observed variation for the end-members) and by the Si-Al distribution along the system (responsible for the deviation from linearity observed along the join).

The solubility of H₂O in Li-bearing systems shows a negative deviation from linearity, which is somehow difficult to interpret in the absence of direct structural data on these melts.

In this context we must point out that the speciation of H₂O in melts along the NaAlSi₃O₈-KAlSi₃O₈ join, when corrected for fictive temperature effects (Romano et al. 1995), appears to be independent of the identity of the alkali but is a simple function of total H₂O content. Therefore it appears that the higher solubility in albitic glasses is both due to an increase of OH and H₂O in comparison with orthoclastic composition. Moreover it has also been demonstrated (Dingwell and Webb 1990; Romano et al. 1995; Nowak and Behrens 1995) that, even though the molecular H₂O is still a stable species in high *T* melt, the amount of OH groups is significantly higher than previously expected up to 70–80 wt% of the total dissolved H₂O in the melt.

We can therefore assume that the progress of Reaction 2 is governed by the chemical stability of the structure and that the molecular H₂O plays a less important role in influencing the total amount of volatile dissolved in these melts. It has been suggested (Pichavant et al. 1992) that molecular H₂O enters the structure by occupying the free volume available. Pichavant et al. (1992) proposed that this free volume could be inversely related to the size of the cations (or directly related to their field strength), and that, consequently, the relative amount of

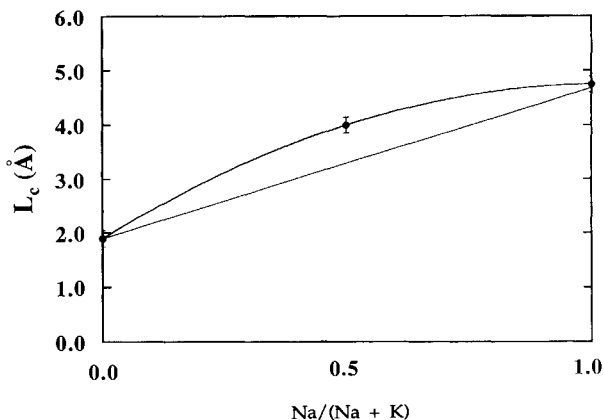


FIGURE 4. The width of the first sharp diffraction peak (FSDP), expressed as correlation length (L_c), as a function of composition (redrawn from Zotov et al. 1995).

H₂O_{mol} clustering around those cations would increase in the order K < Na < Li, in agreement with the observed trend in H₂O solubility here reported. Although the hypothesis of molecular H₂O filling holes in the structure is consistent with the trends observed, such a mechanism implies that the partial molar volume of H₂O in the melt should be close to zero, resulting in a corresponding increase in the density of the melt. Accocella et al. (1984) have demonstrated, however, that the partial molar volume of molecular H₂O and water as OH is 12 cm³/mol for total water contents up to 5 wt%, beyond which value the volume of molecular water increases (and the density decreases), whereas that of water as OH decreases (with increasing density). Such a partial molar volume for H₂O, similar to the partial molar volume per O for other oxide components (Lange 1994), indicates that the solubility of H₂O_{mol} in the melt should not be viewed as the case of passively filling the available free volume or "holes" in the structure.

Pichavant et al. (1992) also proposed (on the basis of enthalpies of hydration of gaseous alkali ions in water at 25 °C) that hydration of the clusters is promoted in the order K < Na < Li yielding the higher solubility in Li-bearing system in comparison with Na- and K-bearing systems. The solvation energies of alkali ions in H₂O could provide a reasonable explanation for the contribution of molecular water to the total water solubility along the present joins.

REFERENCES CITED

- Accocella, J., Tomozawa, M., and Watson, E.B. (1984) The nature of dissolved water in silicate glasses and its effect on various properties. *Journal of Non-Crystalline Solids*, 65, 355–372.
- Behrens, H. (1995) Determination of water solubility in high-viscosity melts: An experimental study on NaAlSi₃O₈ and KAlSi₃O₈ melts. *European Journal of Mineralogy*, 7, 905–920.
- Behrens, H., Romano, C., Nowak, M., Holtz, F., and Dingwell, D.B. (1996) Near-infrared spectroscopic determination of water species in glasses of the system MAI₃O₈ (M = Li, Na, K): An interlaboratory study. *Chemical Geology*, in press.

- Blamart, D., Graham, C.M., and Sheppard, S.M.F. (1986) The system NaAlSi₃O₈ (melt)-H₂O-sea water: Solubility of water and hydrogen isotope fractionation factor at 3, 5 and 8 kbar. International Symposium of Experimental Mineralogy and Geochemistry, Abstracts, 23–24.
- Boettcher, A.L., and Wyllie, P.J. (1969) Phase relationships in the system NaAlSi₃O₈-SiO₂-H₂O to 35 kilobar pressure. American Journal of Science, 267, 875–909.
- Brown, G.E., Ponader, C.W., Waychunas, G.A., and Jackson, W.E. (1988) EXAFS studies of cations environments in silicate melts and glasses. Chemical Geology, 70, 86.
- Burnham, C.W. (1975) Thermodynamics of melting in experimental silicate-volatile systems. Fortschritte der Mineralogie, 52, 101–118.
- (1981) The nature of multicomponent aluminosilicate melts. Physics and Chemistry of the Earth, 13/14, 197–227.
- Burnham, C.W., and Jahns, R.H. (1962) A method for determining the solubility of water in silicate melts. American Journal of Science, 260, 721–725.
- Burnham, C.W., and Davis, N.F. (1971) The role of water in silicate melts: I. *P-T* relations in the system NaAlSi₃O₈-H₂O to 1 kilobar and 1100 °C. American Journal of Science, 270, 54–79.
- (1974) The role of water in silicate melts: II. Thermodynamic and phase relations in the system NaAlSi₃O₈-H₂O to 10 kilobar, 700 ° to 1100 °C. American Journal of Science, 274, 902–940.
- Day, D.E. (1976) Mixed alkali glasses: Their properties and uses. Journal of Non-Crystalline Solids, 21, 343–372.
- Day, H.W., and Fenn, P.M. (1982) Estimating the *P-T-X*-H₂O conditions during crystallization of low calcium granites. Journal of Geology, 90, 485–507.
- DeJong, B.H.V.S., and Brown, J.E., Jr. (1980) Polymerization of silicate and aluminate tetrahedra in glasses, melts, and aqueous solutions: I. Electronic structure of H₆Si₆O₇⁻, H₆AlSi₆O₇⁻, and H₆Al₂O₇⁻. Geochimica et Cosmochimica Acta, 44, 491–511.
- DeYoreo, J.J., Navrotsky, A., and Dingwell, D.B. (1990) Energetics of the charge-coupled substitution Si-NaT in the glasses SiO₂-NaTO₂, T = (T Al, Fe, Ga, B). Journal of the American Ceramic Society, 73, 2068–2072.
- Dietzel, A. (1983) On the so-called mixed alkali effect. Physics and Chemistry of Glasses, 24, 172–180.
- Dingwell, D.B. (1986) Volatile solubilities in silicate melts. In Mineralogical Association of Canada Short Course Handbook, 12, 93–129.
- (1987) Melt viscosities in the system NaAlSi₃O₈-H₂O-F₂O₁. In Geochemical Society Special Publication, 1, 423–433.
- Dingwell, D.B., Harris, D.M., and Scarfe, C.M. (1984) The solubility of H₂O in melts in the system SiO₂-Al₂O₃-Na₂O-K₂O. Journal of Geology, 92, 387–395.
- Dingwell, D.B., and Webb, S.L. (1990) Relaxation in silicate melts. European Journal of Mineralogy, 12, 427–449.
- Farnan, I., Kohn, S.C., and Dupree, R. (1987) A study of the structural role of water in hydrous silica glass using cross-polarization magic angle spinning NMR. Geochimica et Cosmochimica Acta, 51, 2869–2873.
- Fraser, D.G., Rammensee, W., and Jones, R.H. (1983) The mixing properties of melts in the system NaAlSi₃O₈-KAlSi₃O₈ as determined by Knudsen Cell Mass Spectrometry. Bulletin of Mineralogy, 106, 111–117.
- Fraser, D.G., and Bottinga, Y. (1985) The mixing properties of melts and glasses in the system NaAlSi₃O₈-KAlSi₃O₈: Comparison of experimental data obtained by Knudsen Cell Mass Spectrometry and solution calorimetry. Geochimica et Cosmochimica Acta, 49, 1377–1381.
- Fraser, D.G., Rammensee, W., and Hardwick, A. (1985) Determination of the mixing properties of molten silicates by Knudsen Cell Mass Spectrometry: II. The systems (Na-K)AlSi₃O₁₀ and (Na-K)AlSi₃O₁₂. Geochimica et Cosmochimica Acta, 49, 349–359.
- Geisinger, K.L., Gibb, G.V., and Navrotsky, A. (1985) A molecular orbital study of bond length and angle variations in framework structures. Physics and Chemistry of Minerals, 11, 266–283.
- Goranson, R.W. (1931) The solubility of water in granitic magma. American Journal of Science, 22, 481–502.
- (1936) Silicate-water systems: The solubility of water in albite melt. Transactions of the American Geophysical Union, 17, 257–259.
- (1938) Silicate water systems: Phase equilibria in the NaAlSi₃O₈-H₂O and KAlSi₃O₈-H₂O systems at high temperature and pressure. American Journal of Science, 35a, 71–91.
- Hamilton, D.L., and Oxtoby, S. (1986) Solubility of water in albite melt determined by the weight-loss method. Journal of Geology, 94, 626–630.
- Hayward, P.J. (1976) The mixed alkali effect in aluminosilicate glasses: Part 1. The join SiO₂-(Na,K)AlSi₃O₈. Physics and Chemistry of Glasses, 17, 55–61.
- Hervig, R.I., and Navrotsky, A. (1984) Thermochemical study of glasses in the system NaAlSi₃O₈-KAlSi₃O₈-Si₂O₆ and the join Na_{1.6}Al_{1.6}Si_{2.4}O₈-K_{1.6}Al_{1.6}Si_{2.4}O₈. Geochimica et Cosmochimica Acta, 48, 513–522.
- Hess, P.C., and Wood, M.I. (1982) Al coordination in metaluminous and peralkaline silicate melts. Contributions to Mineralogy and Petrology, 81, 103–112.
- Holtz, F., Behrens, H., Dingwell, D.B., and Taylor, R.P. (1992) Water solubility in aluminosilicate melts of haplogranitic compositions at 2 kbar. Chemical Geology, 196, 289–302.
- Holtz, F., Dingwell, D.B., and Behrens, H. (1993) The effects of fluorine, boron and phosphorus on the solubility of water in haplogranitic melts compared to natural silicate melts. Contributions to Mineralogy and Petrology, 113, 492–501.
- Holtz, F., Behrens, H., Dingwell, D.B., and Johannes, W. (1995) H₂O solubility in haplogranitic melts: Compositional, pressure, and temperature dependence. American Mineralogist, 80, 94–108.
- Hovis, G.L. (1984) A hydrofluoric acid solution calorimetric investigation of glasses in the systems NaAlSi₃O₈-KAlSi₃O₈ and NaAlSi₃O₈-Si₂O₆. Geochimica et Cosmochimica Acta, 48, 523–525.
- Isard, J.O. (1969) The mixed alkali effect in glass. Journal of Non-Crystalline Solids, 1, 235–261.
- Kadik, A.A., and Lebedev, Y. (1968) Temperature dependence of the solubility of water in albite melt at high pressures. Geochemistry International, 5, 1172–1181.
- Klonkowski, A. (1983) Bond character and properties of the mixed alkali system Na₂O-K₂O-Al₂O₃-SiO₂. Journal of Non-Crystalline Solids, 57, 339–353.
- Kohn, S.C., Dupree, R., and Smith, M.E. (1989) A multinuclear magnetic resonance study of the structure of hydrous albite glasses. Geochimica et Cosmochimica Acta, 53, 2925–2935.
- Kohn, S.C., Dupree, R., and Mortuza, M.G. (1992) The interaction between water and aluminosilicate magmas. Chemical Geology, 96, 399–409.
- Kümmerlen, J., Schaller, T., Sebald, A., and Keppler, H. (1993) The structural role of water in silicate glasses: ¹H and ²⁹Si NMR evidence. Bulletin of Magnetic Resonance, 14, 278–281.
- Kushiro, I., Yoder, H.S., and Nishikawa, M. (1968) Effect of water on the melting of enstatite. Geological Society of America Bulletin, 79, 1685–1692.
- Lange, R.A. (1994) The effect of H₂O, CO₂ and F on the density and viscosity of silicate melts. In Mineralogical Society of America Reviews in Mineralogy, 30, 331–369.
- Lapp, J.C., and Shelby, J.E. (1987) The mixed alkali effect in lithium-sodium aluminosilicate glasses. Journal of Non-Crystalline Solids, 95, 889–896.
- Lebedev, E.G., and Khitarov, N.I. (1964) The dependence of electrical conductivity of granite melt and the beginning of granite melting on high pressure of water. Geokhimiya, 3, 195–201.
- McMillan, P.F. (1994) Water solubility and speciation models. In Mineralogical Society of America Reviews in Mineralogy, 30, 131–156.
- McMillan, P.F., Piriou, B., and Navrotsky, A. (1982) A Raman spectroscopy study of glasses along the joins silica-calcium aluminate and silica-potassium aluminate. Geochimica et Cosmochimica Acta, 46, 2021–2037.
- McMillan, P.F., and Remmele, R.L., Jr. (1986) Hydroxyl sites in SiO₂ glass: A note on infrared and Raman spectra. American Mineralogist, 71, 772–778.
- McMillan, P.F., and Holloway, J.R. (1987) Water solubility in aluminosilicate melts. Contributions to Mineralogy and Petrology, 97, 320–332.
- Mysen, B.O., and Virgo, D. (1986) Volatile in silicate melts at high pres-

- sure and temperature: Interaction between OH⁻ group and Si⁴⁺, Al³⁺, Ca²⁺, Na⁺, and H⁺. *Chemical Geology*, 57, 303–331.
- Navrotsky, A., Geisinger, K.L., McMillan, P., and Gibbs, G.V. (1985) The tetrahedral framework in glasses and melts: Influence from molecular orbital calculations and implications for structure, thermodynamics, and physical properties. *Physics and Chemistry of Minerals*, 11, 284–298.
- Nowak, M., and Behrens, H. (1995) The speciation of water in haplogranitic glasses and melts determined by in situ near-infrared spectroscopy. *Geochimica et Cosmochimica Acta*, 59, 3445–3450.
- Oestrike, R., Yang, W.H., Kirkpatrick, R.J., Hervig, R.L., Navrotsky, A., and Montez, B. (1987) High-resolution ²³Na, ²⁷Al, and ²⁹Si NMR spectroscopy of framework aluminosilicate glasses. *Geochimica et Cosmochimica Acta*, 51, 2199–2209.
- Orlova, G.P. (1962) The solubility of water in albite melts under pressure. *International Geological Reviews*, 6, 254–258.
- Oxtoby, S., and Hamilton, D.L. (1978a) The discrete association of water with Na₂O and SiO₂ in NaAl silicate melts. *Contributions to Mineralogy and Petrology*, 66, 185–188.
- (1978b) Solubility of water in melts of Na₂O-Al₂O₃-SiO₂ and K₂O-Al₂O₃-SiO₂. *Natural and Environmental Research Council Current Research Report Series D*, 11, 33–36.
- Paillat, O., Elphick, S.C., and Brown, W.L. (1992) The solubility of water in NaAlSi₃O₈ melts: A re-examination of Ab-H₂O phase relationships and critical behaviour at high pressures. *Contributions to Mineralogy and Petrology*, 112, 490–500.
- Pichavant, M., Holtz, F., and McMillan, P. (1992) Phase relations and compositional dependence of H₂O solubility in quartz-feldspar melts. *Chemical Geology*, 96, 303–319.
- Rammensee, W., and Fraser, D.G. (1982) Determination of activities in silicate melts by Knudsen Cell Mass Spectrometry: I. The system NaAlSi₃O₈-KAlSi₃O₈. *Geochimica et Cosmochimica Acta*, 46, 2269–2278.
- (1987) The effect of changing Si/Al ratio on the mixing of melts in the system NaAlSiO₄-KAlSiO₄-SiO₂. *Chemical Geology*, 62, 103–110.
- Richet, P., Roux, J., and Pineau, F. (1986) Hydrogen isotope fractionation in the system H₂O-liquid NaAlSi₃O₈: New data and comments on D/H fractionation in hydrothermal experiments. *Earth Planetary Science Letters*, 78, 115–120.
- Rogez, J., Chastel, R., Bergman, C., Brousse, C., Castanet, R., and Mathieu, J.C. (1983) Etude thermodynamique du système albite-orthose par calorimétrie de dissolution et effusion de Knudsen couplée à un spectromètre masse. *Bulletin de Minéralogie*, 106, 119–128.
- Romano, C., Dingwell, D.B., and Sterner, S.M. (1994) Kinetics of quenching of hydrous feldspathic melts: Quantification using synthetic fluid inclusions. *American Mineralogist*, 79, 1125–1134.
- Romano, C., Dingwell, D.B., and Behrens, H. (1995) The temperature-dependence of the speciation of water in NaAlSi₃O₈-KAlSi₃O₈ melts using fictive temperatures derived from synthetic fluid-inclusions. *Contributions to Mineralogy and Petrology*, 122, 1–10.
- Scholze, H. (1988) *Glas-Struktur und Eigenschaften*, 407 p. Springer, Berlin.
- Shaw, H.R. (1963) Obsidian H₂O viscosities at 1000 and 2000 bars in the temperature range 799–900 °C. *Journal of Geophysical Research*, 68, 6337–6343.
- Silver, L.A., and Stolper, E.M. (1985) A thermodynamic model for hydrous silicate melts. *Journal of Geology*, 93, 161–177.
- (1989) Water in albitic glasses. *Journal of Petrology*, 30, 667–709.
- Silver, L.A., Ihinger, P.D., and Stolper, E.M. (1990) The influence of bulk composition on the speciation of water in silicate glasses. *Contributions to Mineralogy and Petrology*, 104, 142–162.
- Sorapure, R. (1981) The solubility of water in albite melt determined from the nucleation of vesicles following isothermal decompression. *Natural and Environmental Research Council Current Research Report Series D*, 5, 7–9.
- Spera, F.J. (1974) A thermodynamic basis for predicting water solubilities in silicates melts and implications for the low velocity zone. *Contributions to Mineralogy and Petrology*, 80, 276–284.
- Stolper, E.M. (1982a) Water in silicate glasses: An infrared spectroscopic study. *Contributions to Mineralogy and Petrology*, 81, 1–17.
- (1982b) The speciation of water in silicate melts. *Geochimica et Cosmochimica Acta*, 46, 2609–2620.
- (1989) Temperature dependence of the speciation of water in rhyolitic melts and glasses. *American Mineralogist*, 74, 1247–1257.
- Turek, A., Riddle, C., Cozens, B.J., and Tetley, N.W. (1976) Determination of chemical water in rock analyses by Karl Fisher Titration. *Chemical Geology*, 17, 261–267.
- Tuttle, O.F., and Bowen, N.L. (1958) Origin of granite in the light of experimental studies in the system NaAlSi₃O₈-KAlSi₃O₈-SiO₂-H₂O. *Geological Society of America Memoir*, 74, 153.
- Urbain, G., Bottinga, Y., and Richet, P. (1982) Viscosity of liquid silica, silicates and aluminosilicate melts. *Geochimica et Cosmochimica Acta*, 46, 1061–1072.
- Voigt, D.E., Bodnar, R.J., and Blencoe, J.G. (1981) Water solubility in melts of alkali feldspar composition at 5 kbar, 950 °C. *Eos*, 62, 428.
- Watson, E.B. (1979) The effect of dissolved water on cesium diffusion in molten granite. *Eos*, 60, 402.
- Westrich, H.R. (1987) Determination of water in volcanic glasses by Karl-Fischer Titration. *Chemical Geology*, 63, 335–340.
- Yoder, H.S. (1965) Diopside-anorthite-water at five and ten kbar and its bearing on explosive volcanism. *Carnegie Institution of Washington Year Book*, 56, 206–217.
- Yoder, H.S., Stewart, D.B., and Smith, J.R. (1957) Ternary feldspars. *Carnegie Institution of Washington Year Book*, 56, 206–217.
- Zhang, Y., Stolper, E.M., and Wasserburg, G.J. (1991) Diffusion of water in rhyolitic glasses. *Geochimica et Cosmochimica Acta*, 55, 441–456.
- Zotov, N., Boysen, H., Romano, C., Dingwell, D.B., and Yanev, Y. (1995) Neutron diffraction study of feldspar glasses: Mixed alkali effect. *Journal of Non-Crystalline Solids*, 191, 124–131.