Chemical bonding and electronic structures of the Al₂SiO₅ polymorphs, and alusite, sillimanite, and kyanite: X-ray photoelectron- and electron energy loss spectroscopy studies

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ABSTRACT

We have undertaken a detailed analysis of the X-ray photoelectron spectra obtained from the three polymorphs of Al_2SiO_5 ; andalusite, sillimanite, and kyanite. Comparison of the spectra was made based on the chemical bonding and structural differences in the Al- and Si-coordination within each polymorph. The spectra for Si(2p) for all three polymorphs are nearly identical, consistent with the fact that all the Si atoms are in 4-fold (tetrahedral) coordination, whereas the binding energies, peak shapes, and peak widths for Al(2p) vary depending on the type of polymorph. The upper-valence band for all three polymorphs is characterized by four main features derived from O(2p), Al(3s), Al(2p), Si(3s), and Si(3p), and the differences in their contributions are observed. The density of state of the Al_2SiO_5 polymorphs is relatively featureless compared to those observed from α - SiO_2 and α - Al_2O_3 , suggesting that the orbital overlaps span a greater range in energy. The observed band gap energy for Al_2SiO_5 (sillimanite) was $\sim 9.1 \, \text{eV}$, a value in between those for α - SiO_2 ($\sim 8.6 \, \text{eV}$) and α - Al_2O_3 ($\sim 9.6 \, \text{eV}$). The conduction band feature of Al_2SiO_5 was experimentally compared to those of α - SiO_2 and α - Al_2O_3 , and shown that it is indeed intermediate between the α - SiO_2 and α - Al_2O_3 phases.

Keywords: Al₂SiO₅ polymorphs, X-ray photoelectron spectroscopy, low electron energy loss spectroscopy, valence- and conduction band structures, and alusite, sillimanite, kyanite

INTRODUCTION

The three polymorphs of Al₂SiO₅, sillimanite, and alusite, and kyanite, are geologically important minerals, whose crystal structures, and physical and thermodynamic properties have been extensively investigated (Kerrick 1990 and references therein). Crystal structures of sillimanite, andalusite, and kyanite were refined by Burnham (1963a), Burnham and Buerger (1961), and Burnham (1963b), respectively, following the structure determinations of kyanite by St. Naray Szabo et al. (1929), and sillimanite and andalusite by Taylor and Hey (1931). Further refinements of the sillimanite and andalusite structures were carried out using single-crystal neutron diffraction methods (Finger and Prince 1972). Winter and Ghose (1979) determined the thermal expansion and the structural changes of all three polymorphs as a function of temperature up to 1000 °C. The crystal structures of all three polymorphs have also been refined at high pressure [andalusite, Ralph et al. (1984); kyanite, Yang et al. (1997a) and Comodi et al. (1997); and sillimanite, Yang et al. (1997b)].

Sillimanite and andalusite are orthorhombic with space groups Pnnm and Pbnm, respectively, whereas kyanite is triclinic with $P\overline{1}$. The crystal structures of the three polymorphs have one feature in common, namely, half of the Al atoms occur in sixfold (octahedral) coordination forming chains of edge-shared [AlO₆] octahedra parallel to the crystallographic \mathbf{c} -axis. The differences in their stability relations depend critically on the

Lattice dynamical studies of these minerals, including experimental measurements of phonon dispersion relations, density of states (DOS), and theoretical shell model calculations of their vibrational and thermodynamic properties, have been made by Rao et al. (1999). The electron density distributions in all three polymorphs have been determined experimentally from multipole refinements based on high resolution single crystal X-ray diffraction data collected at 100 K by a CCD detector (Dahaoui et al. 2001), and theoretically by ab initio quantum mechanical calculations using the LAPW method (Iglesias et al. 2001). The electric field gradient tensors at the ²⁷Al sites in all three polymorphs calculated from the experimentally and theoretically derived electron density distributions match the values obtained from single-crystal ²⁷Al NMR measurements (Raymond and Hafner 1970; Hafner et al. 1970; Hafner and Raymond 1967; Bryant et al. 1999).

The differences in the chemical bonding and electronic structures of the three Al₂SiO₅ polymorphs are fundamental to the understanding of their thermodynamic properties and

differences in the chemical bonding of the remaining Al atoms in each polymorph: Al is in fourfold (tetrahedral) coordination in sillimanite (stable at low P and high T), fivefold coordination (trigonal bipyramidal) in andalusite (stable at low P and low T), and sixfold coordination (octahedral) in kyanite (stable at high P and high T). The Si atom is in fourfold (tetrahedral) coordination in all three polymorphs. The reader is referred to the paper by Winter and Ghose (1979) for the details of the crystal structures of all three polymorphs.

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stability relations. X-ray photoelectron spectroscopy (XPS) is a technique to probe the chemical bonding of specific elements. However, application of the XPS technique to minerals, like the Al₂SiO₅ polymorphs, is difficult due to sample charging during the analysis. Aleshin et al. (1975) used photoemission spectroscopy to reveal the nature of chemical bonding of the Al₂SiO₅ polymorphs; however, as yet no systematic comparison of the chemical bonding for the three different polymorphs of Al₂SiO₅ is available. Detailed analyses of photoelectron spectra obtained from alumino-silicate glass have been reported (Yagi et al. 2001; Miura et al. 2000). In glass form, however, Al and Si polyhedra are all corner linked, therefore the present study from three different polymorphs can give further insight into the chemistry and electronic structure of bonding about edge- and face-linked polyhedrons. We have undertaken an X-ray photoelectron spectroscopic study of the three Al₂SiO₅ polymorphs to probe the differences in their chemical bonding, and compared the electronic structure of Al₂SiO₅ (sillimanite) with those of α-SiO₂ (quartz) and α-Al₂O₃ (corundum) using proper charge compensation techniques. Low electron energy loss spectroscopy (LEELS) was also used to study the conduction band structure of the sillimanite phase of Al₂SiO₅.

EXPERIMENTAL METHODS

Single-crystal samples of andalusite from Espiritu Santo, Brazil, sillimanite from Sri Lanka, and kyanite from Pizzo Forno, Switzerland were obtained from the Harvard Mineralogical Museum. Sample surfaces examined by XPS and LEELS were prepared by mechanically fracturing the samples in air, and mounted on a sample holder with indium solder. A typical duration for the fractured surface exposed in air (prior to evacuation in the sample load-lock chamber) was less than 30 min. The X-ray photoelectron spectroscopy experiments were performed using two different spectrometers: a modified SSL-300 spectrometer (Surface Science Laboratory Inc.) at the University of Washington, and a Quantum 2000 Scanning ESCA microprobe (Physical Electronics Inc.) at Pacific Northwest National Laboratories, Richland, Washington. In the first spectrometer, a focused monochromatic $AlK\alpha$ X-rays with an X-ray spot of ~600 μm in diameter was used to measure core level data. In the second spectrometer, a monochromatic $AlK\alpha$ X-ray beam (\sim 100 µm diameter) was rastered over an area of \sim 1.4 mm by \sim 0.2 mm on the specimen to collect valence band data. During analysis, charge neutralization was achieved by using a low energy (3-4 eV) electron flood gun. In the Quantum 2000 spectrometer, a low energy Ar-ion beam was applied in addition to the electron flood gun. All spectra were acquired using a pass energy of 50 eV with the SSL-300, and 23.5eV with the Quantum 2000 spectrometer. The binding energies were nominally calibrated using clean Cu and Au foils at Cu 2p_{3/2} = 932.72 eV, Cu 3p_{3/2} = 75.10, and Au $4f_{7/2}$ = 83.98 eV. In the analysis, the absolute value of the binding energy was calibrated by setting the C(1s) peak for the adventitious hydrocarbons accumulated on the surface at 284.60 eV.

For LEELS, the spectra were excited by unpolarized electrons with energies \sim 250 eV incident normal to the surface, and were measured with a single pass cylindrical mirror analyzer (CAM). The current density of the incident electron beam was minimized to prevent electron beam damage of the sample surfaces. A pulse counting mode was used to collect the data. To obtain a constant system resolution, the full width at half maximum (FWHM) of the incident beam was set at a value of about 0.7 eV. For LEELS measurements, only a sillimanite sample was examined, and the results are compared to those from α -Al₂O₃ (corundum) and α -SiO₂ (quartz) phases. The binding energies of O(1s) found in α -Al₂O₃ and α -SiO₂ are 531.21eV and 532.54eV, respectively, which are similar to those reported previously (531.2eV for α -Al₂O₃ and 532.8eV for α -SiO₂) (Barr 1994).

RESULTS AND DISCUSSION

Si(2p) core-level spectra

The Si atoms in all three polymorphs are tetrahedrally bonded to O atoms, and form Al-O-Si linkages. The Si(2p) core level

spectra were carefully measured from sillimanite, andalusite, and kyanite, and compared with that measured from $\alpha\text{-SiO}_2$ (Fig. 1). Analyses indicate that all the spectra reasonably fall into a single Gaussian peak with similar binding energy (E_b) at around 102.28 eV (see Table 1), suggesting that the electronic environment around Si atoms in all three polymorphs of Al_2SiO_5 and $\alpha\text{-quartz}$ ($\alpha\text{-SiO}_2$) is very similar, and that the second nearest-neighbor Al atoms do not significantly influence the electronic stucture of the tetrahedrally-bonded Si atoms.

Al(2p) core-level spectra

Figure 2 shows the Al(2p) spectra obtained from the three polymorphs. In kyanite, there are four symmetry-independent Al atoms present, all in octahedral coordination. The Al-O bond lengths vary within narrow limits (1.816 to 1.997 Å) (Winter and Ghose 1979); therefore, the Al(2p) peak is expected to be reasonably symmetric with a relatively narrow width. Curve fitting to a single Gaussian shape results in a peak maximum at 74.48 ± 0.05 eV with FWHM of 1.68 eV (Fig. 2). Since the Al atomic arrangement in α-Al₂O₃ is also characterized by octahedral coordination with the Al-O bond lengths similar to those found in kyanite, Al(2p) spectra from kyanite and α-Al₂O₃ may be compared (Fig. 3). As shown in Fig. 3, the two Al(2p) peaks are nearly identical in terms of the relative binding energy Note, however, that the Al(2p) peak from kyanite is slightly skewed in its low E_b tail, whereas a nearly perfect Gaussian peak was observed for α-Al₂O₃. This small difference may be due to the influence of the second nearest-neighbor tetrahedral Si atoms in kyanite, and/or greater variation of the Al-O bond distances in kyanite (1.816 to 1.997 Å; Winter and Ghose 1979) vs. α-Al₂O₃ (1.857Å and 1.969 for short and long Al-O bond lengths, respectively; Chiang and Xu 1994).

The Al(2p) peak for sillimanite appears at a similar position in binding energy; however, the peak is considerably broader, and

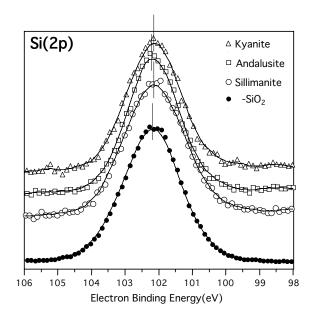


FIGURE 1. Si(2p) spectra obtained for sillimanite, and alusite, and kyanite, and comparison with Si(2p) of α -SiO₂.

Material	Si(2p)			Al(2p)			O(1s)		
	E _b (eV)	FWHM (eV)	Relative fraction %	E _b (eV)	FWHM (eV)	Relative fraction %	E _b (eV)	FWHM (eV)	Relative fraction %
Kyanite	102.27	1.81	100	74.48	1.68	100	531.39 530.59	1.97 1.98	80 20
Andalusite	102.31	1.80	100	74.07	1.70	100	531.69 530.84	1.98 1.93	20 80
Sillimanite	102.28	1.91	100	74.65 74.06	1.70 1.71	50 50	532.02 530.85	1.94 1.92	40 60
α -SiO ₂ α -Al ₂ O ₃	102.28	1.92	100	74.48	1.71	100	532.54 531.21	1.95 1.98	100 100

TABLE 1. Binding energy (E_b in eV), FWHM (eV) and relative fraction (%) for Si(2p), Al(2p), and O(1s) orbitals observed for the three polymorphs of Al-SiO-

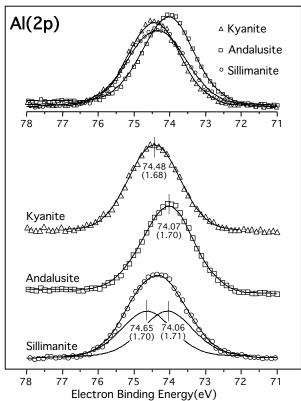


FIGURE 2. Overlay of Al(2p) spectra observed from three polymorphs, and the results from curve fitting. See text for details.

the distribution extends toward the lower E_b side, suggesting two kinds of Al atoms: one similar to that in kyanite (more ionically bonded) and another less ionically bonded (or more covalently bonded). Based on the crystal structure of sillimanite (Winter and Ghose 1979), the Al atoms are in octahedral and tetrahedral coordination with equal fractions. We therefore fitted the curve to two peaks with their ratio constrained to 1:1 (Fig. 2c), resulting in E_b (FWHM) values 74.65 (1.70) and 74.06 (1.71). The analysis indicates that ~50% of the octahedrally coordinated Al atoms in sillimanite are slightly more ionic than those found in kyanite based on their binding energy values (75.65 eV for sillimanite vs. 74.48 eV for kyanite), whereas the other ~50% in tetrahedral coordination are less ionic (E_b ~74.06).

In the andalusite structure, the octahedral Al-coordination is severely distorted: four closely bonded O atoms within a square planar configuration with two each at 1.827 and 1.891 Å, and two farther apical O atoms at 2.086 Å (298 K) (Winter

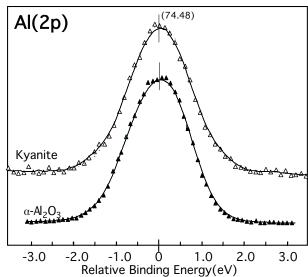


FIGURE 3. Comparison of Al(2p) for kyanite and α -Al₂O₃.

and Ghose 1979). The other Al atom in fivefold coordination is closely bonded to two O atoms at 1.814 Å and one at 1.816 Å forming a trigonal pyramid, while two further O atoms at 1.839 and 1.899 Å complete the trigonal bypyramid (298 K). In view of the structural complexity and hence, the complex nature of the chemical bonding, it does not seem justified to resolve the Al(2p) peak into contributions from the two types of Al atoms in distorted octahedral and the fivefold trigonal bipyramidal coordination. Therefore, we have fitted the Al(2p) spectra to a single peak. A reasonably good fit was found with a peak maximum at 74.06 eV with FWHM of 1.71 eV, as shown in Figure 2. These values are very similar to that for the tetrahedral Al atom in sillimanite (E_b 74.06) (Table 1). Furthermore, the overall Al(2p) peak, shifted toward the lower binding energy by ~0.4 eV compared to kyanite (E_b 74.48), suggests that the bonding for the Al atoms in andalusite is less ionic, compared to the octahedral Al atoms in either kyanite or sillimanite.

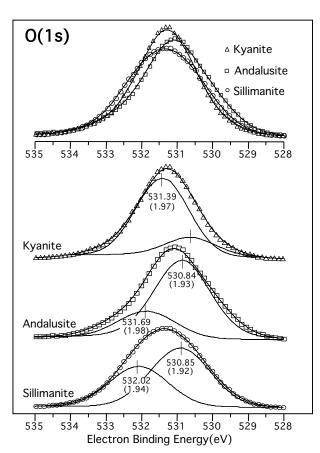
O(1s) core-level spectra

A set of O(1s) spectra obtained from three polymorphs is shown in Figure 4. The overall shapes of the O(1s) peaks observed in the three polymorphs are different. The chemical bonding of the O atoms with Al and Si in the Al₂SiO₅ polymorphs is complex. Four symmetry non-equivalent O atoms are present in both andalusite and sillimanite, whereas kyanite has 10 non-equivalent O atoms. Since the charge density difference for each

oxygen atom may not vary significantly, we do not expect to be able to resolve all these binding energy differences; however, the overall shapes of the O(1s) peaks observed in the three polymorphs are different.

In kyanite, of the 20 O atoms in the unit cell, 16 are involved in Si-O-Al linkage, and 4 in Al-O-Al type linkages. The O atoms in the latter type are expected to be slightly more ionic than the former. We have therefore fitted the spectra with two Gaussian peaks with the ratio of 4:1, and obtained O(1s) peaks corresponding to Si-O-Al linkage (80%) and Al-O-Al linkage (20%) at 531.39 eV and 530.59 eV, respectively (see Table 1). Although the full-width at half-maximum (FWHM) of O(1s) peak ranges from 1.92 to 1.98 eV based on the Gaussian curve fitting, the separation of the peak maxima for the two-types of O(1s) is large enough to distinguish one peak from another.

The bonding differences in the various O atoms in sillimanite and andalusite can be analyzed in terms of the total charge each receives from the surrounding cations. In sillimanite, 12 O atoms are either charge balanced or slightly over bonded, and 8 are slightly under bonded. In andalusite, 4 are under bonded and 16 are either charge balanced or slightly over bonded. Following these considerations, the curve fittings for sillimanite and andalusite were made in 4:1, 2:3, and 1:4 ratios, respectively. Curve fitted results are shown for each polymorph in Figure 4 and Table 1. Our analysis indicates that 60% of the total O atoms



 $\label{eq:FIGURE 4.0} \textbf{FIGURE 4.} \ \text{Overlay of O(1s) spectra observed from three polymorphs,} \\ \text{and the results from curve fitting. See text for details.}$

in sillimanite (E_b 530.85) are more ionic compared to the other 40% (E_b 532.02). This is consistent with the result obtained from Al(2p), in which one half of the total Al atoms is similar to those in kyanite, which is more ionic.

In andalusite, a major peak (>80% of the total) is centered around 530.84 eV, whereas <20% of the oxygen peak appeared with a peak maximum at 531.69 eV, suggesting that the majority of the O atoms are more ionic than those tetrahedrally bonded to Al in sillimanite (532.02). However, this result appears to be inconsistent with the Al(2p) value (74.07), which is comparable to that of tetrahedral Al in sillimanite (74.06). We currently do not have an explanation for this anomaly, and can only suggest that the explanation may lie in the structural complexity and hence, the complex nature of the chemical bonding of andalusite compared to that of kyanite and sillimanite.

Valence-level spectra

Valence-level X-ray photoemission spectra are useful to characterize the nature of the chemical bonding, and may be directly compared with the electron density of state (DOS) calculations. In materials like Al_2SiO_5 , an upper valence band (upper-VB) is characterized by a superposition of Al-Si-O hybridized orbitals and non-bonding O(2p) orbitals, whereas a lower valence band (lower-VB) is mainly due to O(2s) orbitals.

The upper-VB spectra obtained from the three different polymorphs of Al_2SiO_5 are shown in Figure 5. The valence band width is very similar for all three phases, around ~11.0 eV, and the top of the valence band positions are all ~3.9eV below the Fermi energy. Within the VB spectra, however, noticeable differences are observed from each polymorph, which are characterized by four main features labeled a, b, c, and d. The locations of these features in terms of E_b are similar, but the relative contributions vary for different polymorphs, obviously reflecting the types of bonding and coordination for Al, Si, and O atoms in the three polymorphs.

So far, the only DOS calculations available in the literature are those of Iglesias et al. (2001). These authors claimed that the cross-features of the DOS were similar for all atoms in all the three polymorphs. The total upper-VB DOS consists of mainly of O(2p), Al(3s), Al(2p), Si(3s), and Si(3p), in which the Si states appear at the lower energies, and the s-p splitting is more pronounced than for the corresponding Al states. All five orbitals are strongly mixed over the entire upper-VB regions. Note that Iglesias et al. (2001) assumed the top of the valence band as the Fermi-energy for the calculated partial density of states (PDOS), whereas the experimentally observed valence band starts around –4 eV in binding energy. In addition, the calculated binding energies are somewhat smaller than those observed experimentally, although the overall nature of the calculated and experimentally observed valence band structures is very similar.

Based on these calculations, the feature labeled a (around 5~6 eV in Fig. 5) consists of mainly O(2p) contribution with little overlap from the Si and Al orbitals. Feature b around 9eV is the Si(2p and 3p) and Al(2p and 3s) overlapped with O(2p), whereas the s-orbitals from both Al and Si are not strongly involved. Feature c at around 11 eV is similar to b, but has more contribution from the Al(3s) orbitals; therefore, it varies for the three different polymorphs. In our measurements, this was seen

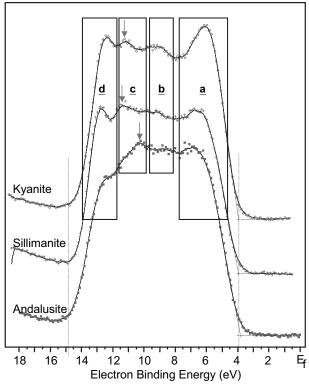


FIGURE 5. XPS valence band spectra from three polymorphs. See text for details about allows in the region c.

at different peak locations in the binding energy for sillimanite, kyanite, and and alusite [as represented by \downarrow (downward-pointing arrow) in Fig. 5]. The last feature labeled d originates mainly from Si(3s) orbitals, weakly overlapped with O(2p) orbitals. In andalusite, this feature d is much weaker than those found for kyanite and sillimanite, suggesting that the covalent nature of the Al-O type bonding is more dominant in this structure. This result is consistent with the fact that kyanite (which is stable at high pressure) has the highest coordination numbers for Al atoms in comparison to and alusite (6 and 5) and sillimanite (6 and 4), and hence an expected higher degree of ionic bonding. The higher ionic nature of Al in kyanite may be the reason why the feature "a" [O(2p) non-bonding] is more enhanced relative to sillimanite and and alusite. Note that a similar peak found in α-Al₂O₃ (feature "g" in Fig. 6) is also enhanced, which is consistent with the above interpretation To confirm this, however, more detailed computations comparing three polymorphs are needed.

It has been suggested that a shift of the leading edge of O(2p) in the valence band toward the Fermi energy is a direct measure of increased ionicity in some materials (Barr 1991). However, we were unable to clearly identify a shift based on the leading edge positions observed in this study.

COMPARISON WITH QUARTZ (α -SiO₂) AND CORUNDUM (α -AL₂O₃)

Electronic structures of quartz (α -SiO₂) and corundum (α -Al₂O₃) have been thoroughly investigated both experimentally and theoretically [for α -Al₂O₃, see Siegel et al. (2002); Shang-Di and Chiang (1998) and references cited therein, and for

α-SiO₂, Gupta (1985); Di Pomponio et al. (1995) and references cited therein]. In quartz, Si atoms are tetrahedrally coordinated, whereas Al atoms in corundum are all octahedrally coordinated. In Al₂SiO₅, however, Al assumes different types of coordination depending on the polymorph. Therefore, it is worth comparing the occupied valence band (VB) and unoccupied conduction band (CB) spectra of Al₂SiO₅ with its chemical components, namely, α -SiO₂ and α -Al₂O₃. For comparison, we have chosen the sillimanite phase to represent the Al₂SiO₅ polymorphs. In this structure, Al atoms are in both tetrahedral and octahedral coordination, whereas the Si atoms are tetrahedrally-coordinated. Figure 6 shows the XPS-VB spectra observed from α -SiO₂, sillimanite and α-Al₂O₃. Most noticeably, the upper-VB widths for the three materials are significantly different: ~13 eV, ~11.5 eV, and ~ 9 eV for α -SiO₂, sillimanite and α -Al₂O₃, respectively. For both α -SiO₂ and α -Al₂O₃, the upper-VB consists of mainly two features; a peak located at lower E_b side (labeled as e or g), and a peak at higher E_b side (labeled as f or h). Based on the DOS calculations, the features, e and g, are predominantly O(p)-non-bonding states with little contribution from O(s). These non-bonding states are spatially localized on the atomic sites and oriented transversally with respect to the Si-O and Al-O bonding direction for the quartz and corundum structures, respectively. The features at higher E_b side, f and h, mainly consist of strongly hybridized orbitals. The calculations indicate that these two features are well separated in both quartz and corundum, reflecting two distinct features found in calculated DOS (Iglesias et al. 2001). Observed DOS calculations for Al₂SiO₅ are relatively featureless compared to those observed for α -SiO₂ and α -Al₂O₃, suggesting that the orbital overlaps are much more spread out over the energy range, as previously described. It should also be pointed out that the valence band edges from the three materials occur nearly at the same binding energy $(3.80 \pm 0.10 \text{ eV})$.

The lower-VB spectra observed from the three materials are also shown in Figure 6. The DOS calculations reveal that the lower valence band is mainly O(2s) with small overlap from 2p-states associated with Si and/or Al. Thus, the differences in bonding and coordination are reflected in the peak locations and their shapes. The lower-VB spectra for sillimanite are similar to those of α -SiO₂ in terms of binding energies, but appear \sim 1 eV higher than that for α -Al₂O₃. Compared with O(1s) peaks, the lower VB peaks are considerably broader (\sim 4 eV of FWHM), and therefore, detailed structures cannot be resolved. Nevertheless, this difference implies that the bonding character for sillimanite is more covalent than that of α -Al₂O₃, but similar to that of α -SiO₂.

Low energy electron loss spectroscopy (LEELS) is a technique for studying unoccupied states since the core electrons are excited into the unoccupied states in the conduction band (Koma and Ludeke 1975; Araki et al. 1976). Because the core states have well-defined energies and angular momenta, LEELS can probe the variation in the angular-momentum-resolved density of the conduction band state as joint density of states (JDOS). Shown in Figure 7 is LEELS spectra obtained from $\alpha\text{-Al}_2O_3$, sillimanite, and $\alpha\text{-SiO}_2$; the spectra for $\alpha\text{-Al}_2O_3$ and $\alpha\text{-SiO}_2$ are excited from the Al(2p) and Si(2p) core levels, respectively, and for sillimanite from both Al(2p) and Si(2p) core levels. The LEELS spectra are then aligned by setting the threshold above 0 eV (Fermi energy) to an amount equal to the loss energy minus the corresponding

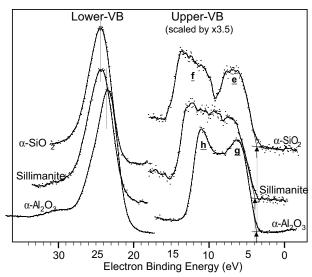


FIGURE 6. Comparison of the upper- and lower-valence bands from α -SiO₂, sillimanite, and α -Al₂O₃. The intensity of the upper-valence band was multiplied by 3.5.

core binding energy previously determined by XPS. The loss peak distribution above 0 eV represents JDOS of the conduction band. The onset energy of the JDOS corresponds to the energy difference between the bottom of the conduction band and the Fermi level of the material. The valence band edge occurs at 3.80 ± 0.10 eV for all three materials. Therefore, the band gap energies for α -Al₂O₃, sillimanite, and α -SiO₂ are determined to be $\sim 9.6 \text{ eV}(3.8 + 5.8)$, $\sim 9.1 \text{ eV}(3.8 + 5.3)$, and $\sim 8.6 \text{ eV}(3.8 + 5.8)$ 4.8), respectively. Published values for the band gap ranges from 5.67 to 11.5 eV for SiO₂ (Young-nian et al. and references therein) and from ~7 to ~10 eV for Al₂O₃ (French et al. 1994; Gillet and Ealet 1992). Explanations for the range include the formation of excitons, presence of defects and temperature dependent Urbach tails. The band gap energy for Al₂SiO₅, sillimanite, is intermediate to those of α-Al₂O₃ and α-SiO₂, which is reasonable based on the mixture of Si-O and Al-O bond configurations in these materials.

The JDOS spectra mainly consist of three primary regions labeled i (at $\sim 4-8$ eV), j ($\sim 10-15$ eV), and k ($\sim 20-26$ eV) above the Fermi energy. Differences in the relative height for each region are related to the transition probability between the core and unoccupied conduction levels. Although transitions that give rise to the core-loss edge are governed by the atomic-dipole selection rules for electronic transitions, the spin-orbit split for Al(2p) and Si(2p) cannot be resolved in the present case. The resultant JDOS deduced from the LEELS, therefore, consists of all s-, p-, and d-like structures of the unoccupied states of the material. The region i for the three materials is similar, but the peak from α-SiO₂ is shifted toward higher energy. In the regions j and k for α-SiO₂ is less pronounced compared to α-Al₂O₃. Sillimanite follows both features shown by α -SiO₂ and α -Al₂O₃, implying that the conduction band for sillimanite is indeed intermediate to those of α -SiO₂ and α -Al₂O₃ phases. Given the broad nature of the conduction band for sillimanite, it is expected that the conduction bands in andalusite and kyanite should be similar to that found in sillimanite.

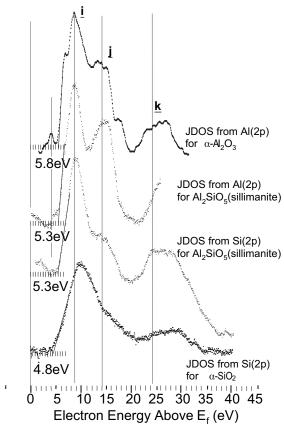


FIGURE 7. Comparison of the LEELS spectra obtained from α -SiO₂, sillimanite, and α -Al₂O₃ by setting the threshold above 0 eV (Fermi energy) to an amount equal to the loss energy minus the corresponding core binding energy previously determined by XPS.

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